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United States Department of the Interior

BUREAU OF LAND MANAGEMENT WASHINGTON, D.C. 20240

> February 5, 1987

Information Bulletin No. 87-106

All Field Officials To:

From: Director

r. U. HOX 25047 Derrer, CO 80225-0049 Subject: Technical Reference 3031-1 "Mineral Deposit Types and Their Characteristics" hadneted by the

Energy and mineral resources assessment are conducted by industry as well as by different government agencies for a variety of reasons. In BLM's case the driving force is a policy need related to the management of public lands through land-use planning.

In subsection .31 Mineral Occurrence Models of the 3031 Manual on Energy and Mineral Resources, the mineral specialist is encouraged to use models when assessing the potential for the occurrence of energy and mineral resources. Geologic models are key to understanding the potential for the occurrence of energy and mineral resources that a given area might have.

Over the years numerous models for different types of mineral deposits have been proposed, but their descriptions are dispersed in the professional literature. To facilitate the work of the minerals specialists and the use of models, we have initiated a project whose product is the subject Technical Reference. The authors have compiled, adapted and up-dated the description of 28 models for mineral deposits found mainly in western United States. This technical reference is organized in a three-ring binder so that changes, up-dates, and additions can easily be made.

The mineral specialists are encouraged to use this Technical Reference, add details pertinent to their respective area, and develop new models applicable to a specific area. If you consider that new models should be added for Bureauwide use, do not hesitate to communicate with Jean Juilland at FTS 653-2270.

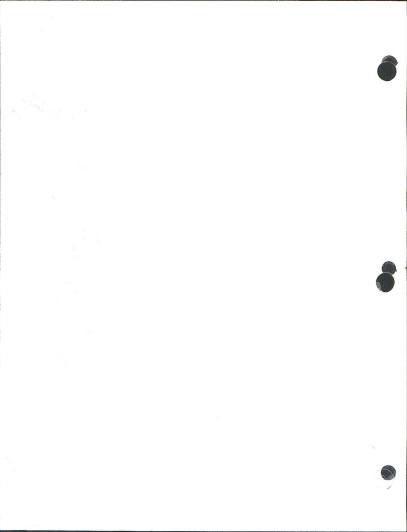
A limited number of copies have been printed and at least one copy is being distributed to each Field Office. Those offices receiving more than one copy should keep at least one copy in the library.

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Assistant Director, Energy and Mineral Resources 1 Attachment (sent under separate cover to addressees only):

1 - Technical Reference 3031-1 "Mineral Deposit Types and Their Characteristics"





PREFACE

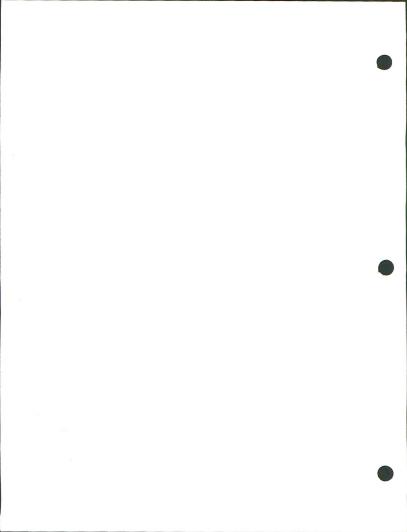
MINERAL DEPOSIT TYPES AND THEIR CHARACTERISTICS is a compilation and adaptation of existing information for Bureau of Land Management (BLM) use. The title might have been enhanced by including the term "mineral models" but the present title is more in keeping with the intent to make this presentation less complex and more basic. However, the two terms, "mineral types" and "mineral models" are used interchangeably in this report.

Mineral deposit characteristics, or attributes, or models are described as working hypotheses developed by observation, description, statistical data, or analogy of a phenomenon or process that cannot be observed directly, or that can be observed with great difficulty. The term "mineral deposit model" is currently popular for describing an age-old procedure of recognizing consistent geologic, geochemical and/or geophysical features that are considered to be related to, or the cause of, a certain type of mineral deposit. In its simplest form it has been used by prospectors and explorers for many years. In the last three decades improvement in analytical instruments and techniques has made it possible to partially understand the various interacting geologic, geochemical and geophysical processes that take place in the formation and emplacement of a certain type of mineral deposit. Also with increased dissemination of geologic information which is now available, the characteristics of deposits can be compared worldwide and analyzed for their critical anticula attributes.

The deposit types included in this report have been compiled from existing literature. They partially resemble a series of models published several years ago by the U.S. Geological Survey as Open Files Reports (82-795, 83-423, 83-623, 83-901, and 83-902). However, additional information from other publications and from personal files has been included. Selected models were expanded and revised to fit the type of deposits found mainly in western United States.

As our scientific knowledge advances, existing mineral types are improved and refined, and new types are developed, this report can be updated and changed as necessary. The user is strongly encouraged to develop his own models for a specific region if the need arises.

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MINERAL DEPOSIT TYPES AND THEIR CHARACTERISTICS

by D. G. Fisher and J. D. Juilland

November 1986

Technical Reference 3031-1

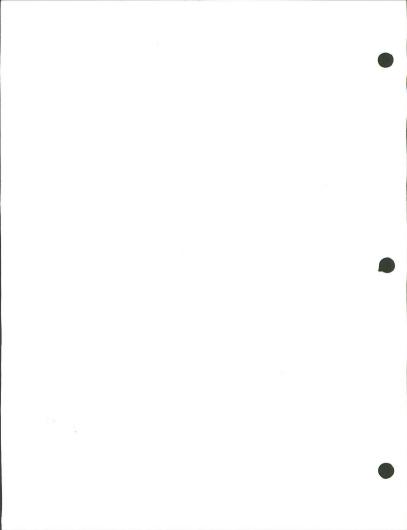
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U.S. Department of the Interior Bureau of Land Management Service Center Division of Resources Branch of Physical Resources

Denver, CO 80225-0047

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INTRODUCTION

Assessment of the potential for energy and mineral resources is a complex and difficult process which involves integration of many diverse factors. While some of these factors can be measured, others are judgemental and depend on the experience and knowledge of the mineral specialist making the assessment. To alleviate this intricate and trying process, the geologic profession has systematically arranged the information that describes the characteristics and essential attributes of a group of mineral deposits sufficiently similar that they suggest a related origin. This is the basic technique used in mineral modeling and it is an ever-evolving process that is constantly being refined. It is considered by most explorationists as an essential tool for understanding the potential that a given area might have for the occurrence of a mineral deposit.

There are several types of models, descriptive, genetic, probability of occurrence, grade-tonnage, and qualitative models, all based on increasing complexity. For BLM assessment work, the descriptive type is the most useful in most cases and is the one used in this report.

Because of different objectives, the use of mineral deposit models by the BIM is somewhat different. The areas which have to be assessed (target areas) have been already defined - the Resource Areas. The available geologic, geochemical, geophysical and mineral occurrence data are matched with the characteristics of existing models for various types of energy and mineral deposit. In addition, use of models can help the Bureau's mineral specialist determine what additional information may be needed for determining the mineral potential and/or for increasing the level of certainty as to whether a deposit is likely to occur.

The use of mineral deposit models technique helps the mineral specialist organize ideas and information while focusing on energy and mineral resources potential. It enables the mineral specialist to compare his/her observations with the collective knowledge of a wide group of experts. A major determinant of creditity is excellent documentation; using mineral deposit models, documents the assessment process thus allowing prompt revision whenever necessary and providing background support whenever the results are questioned. All these advantages will enhance the Bureau's ability to produce authoritative energy and mineral resources assessments.

The use of mineral deposit models by BLM would be primarily in the GEM resources assessment for input to planning, and other instances where the potential for energy and mineral resources meeds to be assessed.

In the GEM resources assessment the mineral deposit models technique is used as follows:

(1) Define area of interest. For input to planning the area of interest is usually the Resource Area.

(2) Literature search; obtain and organize geologic, geochemical, geophysical, mineral and other pertinent data. (3) Match available data with available models and select those models with which there is complete or predominant coincidence.

(4) Identify additional data if any are needed.

(5) Using appropriate models and available data, classify the land as to its potential for energy and mineral resources.

The 3031 Manual gives detail for the preparation of an assessment.

The deposit types presented in this report are arranged according to five basic modes of occurrence. These are felsic plutonic, submarine volcanic-hosted, vein and replacement (epigenetic), sediment-hosted, and sedimentary. The numbering system used reflects the group association.

The data that make up each type are subdivided into four main subparts and consist of an introductory section, a section on geologic features, a section on mineral characteristics, and a list of deposit examples and literary references. When evailable, a cross-section illustration of the type is also included.

In the introductory section of a deposit type, the common or scientific name in current use is given under **Deposit Type**. If there are important distinctions from the general type, a separate subtype is developed and named under Subtype and denoted by an A, B, etc. When there is another commonplace name for the type, this name is given under Synonym.

The GEOLOGIC FEATURES subpart contains Regional Setting where a description is given of the general tectonic position(s), related structural features, and relative age. Next, is Structural Features where the type and orientation of favorable fractures, faults, lineament patterns, and intrusive bodies are discussed. The last item of this subpart is Stratigraphic and Lithologic Characteristics which describes recognized favorable formations and rock types (host rocks).

Under the MINERAL CHARACTERISTICS subpart, Deposit Features describes in general the size, shape, grade and tonnage range, the ore, and associated minerals and commodities usually found. Next, listed under Alteration is a description of the types and patterns of hydrothermal alteration commonly found with the respective mineral deposit. Related stratigraphic, structural, and/or geochemical influences on the deposit are listed next under Ore Control. Any special weathering characteristics or secondary minerals that might serve as a prospecting guide are given under Weathering. Anomalous (either enriched or depleted) amounts of any pathfinder elements in or near a deposit and the zonal pattern are noted under Geochemical Expression. Under Geophysical Expression, any geophysical signature known to be associated with a particular type is described.

The last subpart is a list of Deposit Examples arranged by state with the corresponding literature reference. This is not an all inclusive list, but rather a list of some of the more typical deposits of this type. Also, some experts disagree on the proper classification of the deposits listed. The References chosen were the most up-to-date that still gave the best overall general description.

Deposit Type: PORPHYRY COPPER.

Subtype: This is for the general type. A few deposits can be differentiated into molybdenum-rich and/or gold-rich copper porphyries, but the observable differences are very subtle.

<u>Description</u>: A disseminated replacement deposit in which the copper minerals occur as discrete grains and closely spaced veinlets (stockwork) in a large volume of rock that is nearly always porphyritic.

GEOLOGIC FEATURES

Regional Setting: Continental margins or fragments of margins having abundant igneous intrusive rocks ranging principally from Mesozoic to Tertiary in age.

Structural Features: Most deposits are spatially or genetically related to intrusives that form as stocks, plugs, sills, or dikes. The intrusive action may be forceful, passive or permissive, but the larger deposits seem to favor passive intrusions. Breccia and brecciation pipes are common around some deposits.

Stratigraphic and Lithologic Characteristics: The intrusive rocks may range widely in composition, but quartz-bearing intermediate rocks are most common. Such types as quartz monzonite, granodiorite, and tonalite are favorable. At least one facies of the intrusive is porphyritic with a microgranular aphanitic quartz- feldspar groundmass. Carbonate rocks are particularly favorable when intruded by an intermediate to felsic porphyry. Siliceous metamorphic and other sedimentary rocks are next in order of preference.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: There is no predominate shape; some are flat or saucershaped whereas others are pipe-like with a circular or oval cross-section. Where figures are available, tonnages are in the 20 to 1100 million range. Copper content commonly ranges from 0.30 to 1.00 percent; molybdenum, 0.015 to 0.1 percent; gold, 0.10 to 0.65 g/ton; silver, 0 to 4.0 g/ton. Supergene sulfides may form a higher grade zone of secondary enrichment. Primary minerals are pyrite, chalcopyrite, and molybdenite. Less common is bornite, enargite, and scheelite. The gold and silver rarely form distinct minerals.

<u>Alteration</u>: One type of alteration or another is always found accompanying this type of mineralization. Five distinct types have been recognized. Propylitic alteration is characterized by lime-bearing minerals such as calcite and epidote and is generally found at the outer edges of the porphyry area. Argillic alteration is characterized by the presence of clay minerals (kaolinite or montmorillonite group) and by strong leaching of lime minerals giving a bleached look to the rocks. Potassic alteration is characterized by the assemblage of muscovite-biotite-potassium feldspar (no clay minerals are formed). Quartz-sericite alteration is marked by the absence of clay and potassium feldspars. Lime-silicate alteration occurs when carbonate sediments are replaced with lime-siltcate minerals at the intrusive contact. Not all deposits display all five types; in fact, some only exhibit one or two. Some deposits have been found to have undergone simultaneous alteration of different types at separate points within the same deposit. Others show repeated episodes of alteration over time.

Ore Controls: Mineralized closely spaced fractures and faults as the result of folding, and doming of the host rocks are necessary to allow access for the ore-forming solutions.

<u>Weathering</u>: The surrounding surface area may show widespread iron oxide staining depending upon the original amount of pyrite in the rocks and the degree of weathering. Frequently a leached capping forms over a deposit. If the area has undergone long periods of erosion, these features may be largely removed or buried beneath alluvium.

Geochemical Expression: Anomalous amounts of copper, molybdenum, gold, silver, and tungsten are found in the central part of a deposit. Lead, zinc, gold, silver, molybdenum, arsenic, antimony, tellurium, manganese, and rubidium may be found in the peripheral areas.

On a regional basis, stream sediment values that might be anomalous are:

Cu 50-100 ppm Pb 50-100 ppm Te 100-400 ppb

For a district-wide reconnaissance, anomalous values for stream sediments are:

Cu	100 ppm	Mn	1000 ppm
РЬ	100 ppm	zn	100 ppm
Te	400 ppb	Mo	5 ppm
Ag	0.5 ppm		
A11	50 nnh		

<u>Geophysical Expression</u>: Although magnetic methods have not been successful in directly finding deposits, it has been noted that some deposits are situated on the flanks of magnetic highs. Induced polorization has been used to detect pyrite halos that surround some porphyries. Other methods have been largely unsuccessful. When potassic alteration is present and exposed at the surface, gamma-ray spectrometry surveys could detect the potassium rich zone.

Associated Mineralization: Lead, zinc and silver veins are frequently found at the outer fringes of a porphyry system. In a few deposits, deep erosion has produced placer gold downstream from the main mineralization.





Deposit Examples:

Bond Creek (AK) Orange Hill (AK)

Ajo (AZ) Bagdad (AZ) Bisbee (AZ) Bluebird (AZ) Carpenter (AZ) Castle Dome (AZ) Christmas (AZ) Copper Basin (AZ) Copper Cities (AZ) Copper Creek (AZ) Dos Pobres (AZ) Florence (AZ) Helvita (AZ) Inspiration (AZ) Ithica Peak (AZ) Kalamazoo-San Manuel (AZ) Lakeshore (AZ) Metcalf (AZ) Miami (AZ) Mineral Butte (AZ) Pima-Mission (AZ) Ray (AZ) Red Mountain (AZ) Sacaton (AZ) Safford (AZ) Sanchez (AZ) San Juan (AZ) San Xavier (AZ) Sierrita-Esperanza (AZ) Silver Bell (AZ) Twin Buttes (AZ) Vekol (AZ)

Lights Creek (CA)

Catheart (MN)

Butte (MT) Heddleston (MT)

Bear (NV) Copper Canyon (NV) Ely (NV) MacArthur (NV) Yerington (NV)

References:

Singer and Mosier, 1983 Singer and Mosier, 1983 Dixon, 1966 Anderson et al., 1956 Bryand and Metz, 1966 Singer and Mosier, 1983 Singer and Mosier, 1983 Peterson, 1962 Koski and Cook, 1982 Johnston and Lowell, 1961 Simmons and Fowells, 1966 Singer and Mosier, 1983 Langon and Williams, 1982 Nason et al., 1982 Creasey and Quick, 1955 Olmstead and Johnson, 1966 Dings, 1952 Lowell, 1968 Hallof and Winniski, 1971 Moolick and Durek, 1966 Peterson, 1962 Singer and Mosier, 1983 Kinnison, 1966 Metz and Rose, 1966 Corn, 1975 Cummings, 1982 Robinson and Cook, 1966 Singer and Mosier, 1983 Blake, 1971 King, 1982 West and Aiken, 1982 Graybeal, 1982 Barter and Kelly, 1982 Chaffee, 1977 Storey, 1978 Singer and Mosier, 1983 Meyer, 1968 Singer and Mosier, 1983 Singer and Mosier, 1983

Nash and Theodore, 1971 Bauer et al., 1966 Singer and Mosier, 1983 Wilson, 1963

#2.01

Deposit Examples: Hillsborough (NM) Santa Rita (NM) Tyrone (NM) Bingham (UT) American River (WA) Buckindy (WA) Camp Creek (WA) Crescent (WA) Glacier Peak (WA) Goat Haven (WA) Gold Mountain (WA) Manzama (WA) Margaret (WA) Middle Fork (WA) Mineral Creek (WA) Miners Queen (WA) Monument (WA) North Fork (WA) Red Mountain (WA) Ross (WA) Round Mountain (WA) Silver Creek (WA) Sunrise (WA) Van Epps (WA)

Cloud Home Peak (WY) Deer Creek (WY) Eagle Creek (WY) Kirwin (WY) Needle Creek (WY)

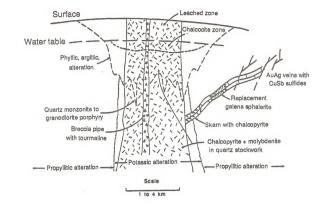
References:

Dunn, 1982 Rose and Baltosser, 1966 Kolessar, 1982

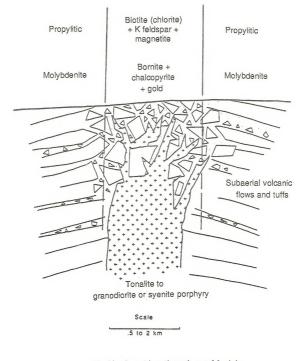
James, 1978

Hollister, 1979 Hollister and Baumann, 1978 Hollister, 1979 Hollister, 1979 Hollister, 1979 Hollister, 1979 Hollister, 1979 Hollister, 1979

Fisher, 1982 Fisher, 1982 Fisher, 1982 Fisher, 1982 Fisher, 1982



Idealized section through a molybdenum-rich variant of a porphyry copper deposit (from USGS Open File Report 83-423) #2.01



Idealized section through a gold-rich variant of a porphyry copper deposit (from USGS Open File Report 83-423)

Deposit Type: PORPHYRY MOLYBDENUM

Subtype: Climax or granite type.

Description: A stockwork containing disseminated grains and veinlets of quartz and molybdenite in a fractured or brecciated, hydrothermally altered granite porphyry intrusive.

GEOLOGIC FEATURES

Regional Setting: Confined mainly to the continental interior and in belts of Cretaceous-Tertlary tectonic activity that were disturbed by subsequent igneous activity. The deposits show some relations to rift zones that are characterized by continental thyolite-basalt associations.

Structural Peatures: Complex structural intersections probably served to locate the intrusives, but surface indications are not always present. The deposits commonly occur in clusters or groups within an area 5 to 20 miles in radius. Ore bodies tend to be dome-shaped and centered on intrusive cupolas. Simple to composite intrusives, dikes and breacia pipes localize the ore. Such features as steeply dipping radial and concentric dikes, veins, faults, and joints are indicative of forceful emplacement of magmatic cupolas. Intensive, widespread, and continued fracturing of host intrusives and enclosing country rocks is characteristic. Almost invariably one or more major faults pass through or close by the ore bodies.

<u>Stratigraphic and Lithologic Characteristics</u>: Deposits of this type are found in nigh silica, alkali-rich porphyry granites having glassy and sometimes smoky quartz phenocrysts. These granites are mid-Tertlary or younger in age. Favorable granites will have \geq 250 ppm rubidium, \leq 50 ppm strontium, \geq 50 ppm niobium, \leq 300 ppm barium, and be enriched in fluorine and tin.

MINERAL CHARACTERISTICS

Deposit Features: Many complex shapes are exhibited, but the domal, funnelshape, accuate, and annular bodies are most common. The deposits may range from a few hundred to several thousand feet in horizontal dimension and may extend to depths of several thousand feet. Commercial deposits range in size from about 50 to 900 million tons. Molybdenite as the ore mineral, forms disseminations and veinlets giving an overall grade of 0.1 to 0.3 percent molybdenum. This is accompanied by pyrite, fluorite, small amounts of tungsten, tin, and zinc minerals. Fluorite and/or topaz are especially notable by their presence in most deposits.

<u>Alteration</u>: Intense hydrothermal alteration of the host intrusive as well as of country rock is always displayed. The degree and sequence of the alteration stages varies widely among deposits. A zonal pattern is commonly displayed with intense quartz and potassium feldspar veins in the orebody followed by argillic and propylite alteration above and along the sides of the main mineralization. Greasen veins found below and sometimes extending through the ore zone are thought to represent a late barren stage.

Ore Controls: Intense fracturing of the intrusive and intruded rocks allowed hydrothermal fluids to permeate the region over small cupolas. The involved rock must be competent enough to prevent venting and escape of the ore-bearing fluids.

<u>Weathering</u>: The surface expression relates to the amount of erosion and the degree alteration present. Many deposits show striking color variation due to black manganese oxides, yellows from ferinolybdenite, jarosite, and limonite, bright reds from lepidocrosite and hematite all against a background of bleach rock. Probably the most consistent feature is the iron staining for oxidized pyrite in the ore body (if exposed).

Geochemical Expression: Anomalous amounts of molybdenum, tin, and tungsten are found especially above the deposit. A zinc and lead halo may show a half mile or more from a deposit. A low silver halo is sometimes found outside the zinc anomaly. Fluorite is found through and around the mineralization area most often in the form of pale blue or green variety. Copper is abnormally low.

<u>Geophysical Expression</u>: Gravity lows of moderate to large intensity and of small areal extent may reflect both the siliceous rocks and the alteration around mineralized cupolas, depending on the density of the host rock. There is some evidence to suggest that favorable areas are along the margins of large regional gravity lows that are presumed to reflect a buried parent mass of batholithic size. Induced polarization may detect ore zones of molybdenite and/or quartz-sericite-pyrite alteration. Generally, none of the current geophysical methods have been overly successful in detecting this type of deposit.

Associated Mineralization: Tin in the form of cassiterite and tungsten as huebnerite or wulfenite may be present in high enough quantities to be recovered as a co-product. Above the main molybdenum mineralization zone, it is not uncommon to find a broad zone of veins and veinlets of base metal sulfides, indochrosite, and fluorite.



#2.02A

Deposit Examples:

Chicago Basin (CO) Climax (CO) Hahns Peak (CO) Henderson-Urad (CO) Horseshoe Bend (CO) Leavenworth Creek (CO) Mount Emmons (CO) Mye (CO) Redwell (CO) Treasure Mountain (CO) Turquoise Lake (CO) Winfleld (CO)

Cumo (ID) Ima (ID) Little Falls (ID)

Bald Butte (MT) Big Ben (MT) Emigrant Gulch (MT)

Majuba Hill (NV) McDermitt (NV) Mount Hope (NV)

Questa-Goat Hill (NM)

Cave Peak (TX)

Marysvale (UT) Pine Grove (UT) Sand Pass (UT) References:

Schmitt and Raymond, 1977 White et al., 1981 Segerstrom and Young, 1972 Wallace, 1978 Burbank and Luedke, 1964 Taylor and King, 1967 Dings and Robinson, 1957 Thomas and Galey, 1978 Harshman, 1965 Sharp, 1978 Mutschler, 1976 Tweto, 1974 Ranta, 1974 Shannon, 1971 Rostad, 1971 Rostad, 1966

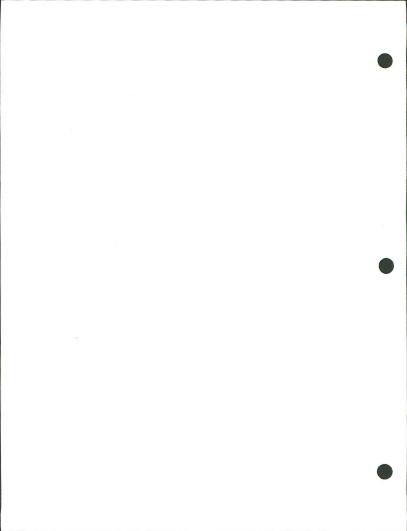
Rostad, 1969 Witkind, 1973 Kirkemo et al., 1965

MacKenzie and Bookstrom, 1976 Greene, 1976 Merriam and Anderson, 1942

Leonardson et al., 1982

Sharp, 1979

Nuella et al., 1979 Abbot and Williams, 1981 Ludington, 1982



Deposit Type: PORPHYRY MOLYBEDNUM

Subtype: Fluorine-deficient type

Synonym: Calc-alkaline molybdenum stockwork

<u>Description</u>: A stockwork containing disseminated grains and veinlets of quartz and molybdenite in fractured and brecciated, hydrothermally alterated felsic porphyry.

GEOLOGIC FEATURES

Regional Setting: Confined mostly to continental margins and areas of back-arc spreading associated with converging plate margins in intercratonic rifts.

Structural Features: The deposits generally parallel the tectonic grain of Mesozoic batholiths and are probably controlled by the overall extent of the associated igneous rocks. Many deposits are related to small intrusives that are satellites of a batholith. Location of the deposits suggests that they are more than casually related to the western U.S. porphyry copper belt.

Stratigraphic and Lithologic Characteristics: Favorable host rocks are monzonltes, quartz monzonite, quartz granodiorites, and quartz diorites. The texture should exhibit conspicuous quartz phenocrysts set in a welldeveloped, microgranular quartz-potassium feldspar ground mass. One of the diagnostic indicators of this type of deposit is abundant, closely spaced quartz veinlets throughout large volumes for rock (stockworks). Age of the intrusives range from late Cretaceous to late Tertiary. Age of the intruded rocks may range as old as Upper Cambrian. Plutons favorable to this type of deposits have 100-250 ppm rubidium, 100-200 strontium, and 20 ppm niobium.

MINERAL CHARACTERISTICS

Deposit Features: The main mineralization consists of molybdenite and pyrite disseminated and in veinlets and fractures. Minor amounts of associated scheelite and chalcopyrite may occur. Molybdenum content commonly ranges from 0.05 to 0.15 percent. Fluorine content is low (0.25 percent) and this usually absent. Pyrite content is known to increase in the upper parts of some deposits and secondary enrichment of copper may occur in the central part of the system. Most deposits are crudely circular in size and less than 1500 meters in diameter. Tonnages range from about 15 to 700 million tons.

<u>Alteration</u>: Most deposits show a potassic core, a quartz-sericite-pyrite intermediate zone, an outer and upper argillic zone, and a propylitic halo. The highest concentration of molybdenite is found at the outer edge of the potassic zone and within the inner quartz-sericite-pyrite zone. Minor amounts of chalcopyrite may surround the molybdenite mineralization zone. Ore Controls: Mineralization is controlled by the stockwork of fractures and breccia zones within the intrusive. Ore minerals may stop abruptly at the intrusive contact or gradually diminish in occurrence.

Weathering: Surface features will vary considerably depending on the degree and amount of erosion of the ore body. Oxidation of the pyrite will produce widespread iron staining. Molybdenite may oxidize into yellow ferrimolybdenite and copper into green oxide stains.

<u>Geochemical Expression</u>: Many deposits showing a zoning outward and upward from molybdenum to copper to zinc and gold to lead, gold, and silver. Productive stockworks typically show 70-300 ppm molybdenum and 20-40 ppm tungsten. Soils are found to be reliable in some areas while rock sample are better in others.

Geophysical Expression: Very few of the known deposits have documented geophysical signatures. If the molybdenite and/or pyrite venilets are concentrated enough, electrical and electromagnetic methods might be successfully employed. Also, if hydrothermal alteration has not been intense enough to destroy the primary magnetite, the resulting rocks should show a low magnetic susceptibility. If the potassic core has been exposed by erosion, gamma-ray spectrometry might be used to successfully identify this core.

Associated Mineralization: Sporadic vein deposits of chalcopyrite-enargitebornite-molybdenite, or pyrite-gold, or sphalerite-gelenargold-silver are sometimes found in the peripheral zones of productive stockworks.

#2.02B

Deposit Examples:

Burroughs Bay (AK) Quartz Hill (AK)

Thompson Creek (ID) White Cloud (ID)

Cannivan Gulch (MT) Turnley Ridge (MT)

Buckingham (NV) Hall (NV) Pine Nut (NV) UV Industries (NV)

Mt. Tolman (WA)

References:

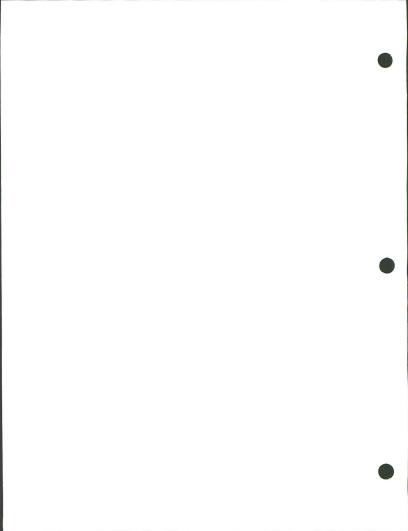
Hudson et al., 1981 Hudson et al., 1979

Schmidt, 1978 Bennett, 1978

Schmidt and Worthington, 1971 Senter, 1976

Theodore, 1982 Kirkemo et al., 1965 Schilling, 1962 Menzie and Theodore, 1983

Menzie and Theodore, 1983



Deposit Type: IRON SKARN

<u>Subtype</u>: Distinction can be made between calc-from skarns and magnessium-from skarns based on tectonic setting, associated intrusive, and related minor metals. In most cases, the outward difference in appearance of either is minor. This model will treat both subtypes as one, noting only the pertinent differences.

Synonym: Contact metamorphic, contact metasomatic, pyrometasomatic, or iron tactite.

Description: Magnetite in coarse-grain Ca-Pe-Mg-Mn silicates formed by the replacement of carbonate-bearing rocks at or near the contact with an igneous intrusive.

GEOLOGIC FEATURES

Regional Setting: Calc-iron deposits are related to island areas and rifted continental margins while magnesium-iron deposits are nearly all found along continental margins.

Structural Features: Deposits form at the contact of a younger igneous intrusive and older carbonate rocks, calcareous clastic rocks, or occasionally in volcanic sequences.

Stratigraphic and Lithologic Characteristics: The magnesium skaras are associated with rocks ranging from gabbros to granites. Calc skaras prefer intrusives of gabbro, diabase, diorite, or syenite having a medium to fine-grain texture and forming large to small stocks and/or dikes. Thinly interbedded sequences of limestone and volcanics seem to be more favorable than thick carbonate beds. Magnesium skaras are usually found in association with hypabysal stocks and dikes of granodiorite, quartz monzonite, and occasionally in granite. Dolomite is the preferred host rock over limestone. Nearly all the deposits are Mesozoic or younger in age, but could be any age.

MINERAL CHARACTERISTICS

Deposit Features: The ore bodies range from 300,000 tons to 170 million tons and have a typical grade of 40 percent iron. They form tabular to irregular mass along the intrusive contact. Mineralization consists of magnetite and minor amounts chalcopyrite, cobaltite, pyrite, and pyrchotite. All of these minerals are enclosed in massive amounts of garnet, pyroxene, and epidote that make up the skarn.

Alteration: Widespread and intense alteration of the igneous rocks is more common of the calci type than of the magnesium type. Alteration causes the formation of diopside, hedenbergite, grossularite and andradite garnets, and additional epidote. Ore Controls: The igneous contact and the attending fractures of the intrusive action are probably the major contributors. Secondary control is probably exerted by the permeability of the carbonate rocks and/or calcareous rocks.

Weathering: No particular distinctive pattern for this type of deposit. Magnetite is somewhat resistant to weathering and should crop out as float. Any pyrite and/or pyrrhotite would oxidize and form the characteristic brown stain on enclosing tocks.

<u>Geochemical Expression</u>: Anomalous amounts of copper, cobalt, and gold can be found around many deposits. In the calc-types addition anomalies of zinc, gold, nickel, molybdenum, and silver are sometimes displayed.

<u>Geophysical Expression</u>: Very little information is available on use of geophysics in exploring for this type of deposits. The most obvious choice would be aeromagnetics for coverage of large areas. Detection response would depend on the amount overburden, thickness and size of the deposit.

Associated Deposits: There are no known associated mineralizations with this type of deposit.

#2.03

Deposit Examples:

Beck (CA) Cave Canyon (CA) Eagle Mountain (CA) Iron Hat (CA) Lake Hawley (CA) Lava Bed (CA) Old Dad Mountains (CA) Shasta-California (CA) Silver Lakes (CA) Vulcan (CA)

Elkhorn Peak (MT)

Dayton (NV)

Dover (NJ)

Capitan (NM) Copper Flat (NM) Guchillo-Negro (NM) Flerro-Hanover (NM) Galilias (NM) Iron Mtn.-Colfax Co (NM) Jicarilla (NM) Jones Camp (NM) Orogrande (NM) Santa Rita (NM) Yellow Jacket (NM)

St. Lawrence (NY)

Cornwall (PA) Dillsburg (PA) French Creek (PA) Grace (PA)

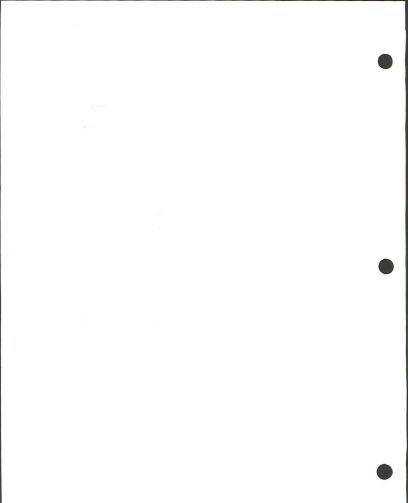
Iron Springs (UT)

References:

Wright et al. 1953 Wright et al. 1953 Dubois and Brummett, 1968 Lamey, 1948 Durrell and Proctor, 1948 Lamey, 1948 Wright et al. 1953 Lamey, 1948 Lamey, 1948 Lamey, 1948 Knopf, 1913 Reeves et al., 1958 Simms, 1958 Kelly, 1952 Kellev, 1949 Kelley, 1949 Hernon and Jones, 1968 Kelley, 1949 Kellev, 1949 Harrer and Kelly, 1963 Kelley, 1949 Kelley, 1949 Harrer and Kelly, 1963 Harrer and Kelly, 1963 Kelley, 1949 Buddington and Leonard, 1962 Lapham, 1968

Hotz, 1960 Smith, 1931 Simms, 1968

Mackin, 1968



Deposit Type: COPPER SKARN

Subtype: This is a general type that is sometimes subdivided into skarns associated with porphyry copper deposits and skarns associated with nonporphyry copper deposits. Distinction is made mainly on the basis of size and accessory minerals.

Synonym: Contact metamorphic, contact metasomatic, pyrometasomatic, or copper tactite.

<u>Description</u>: Copper minerals with Ca-Fe-Mg-Mn silicates consisting primarily of garnet and formed by the replacement of carbonate-bearing rocks at or near the contact with an igneous intrusive.

GEOLOGIC FEATURES

Regional Setting: The majority of these deposits are related to the orogenic belt at continental margins. Very few are situated in an oceanic island-arc setting.

Structural Features: Those associated with porphyry copper take on the distinctive features that are found with this type of deposit. The nonporphyry subtype generally has no recognizable structure features but is probably related to zones of weakness that enable the related intrusive to be implanted as stocks and plutons.

Stratigraphic and Lithologic Characteristics: Both subtypes of copper skarms are associated with Felsic, porphyry-textured stocks ranging in age mainly from Mesozoic to Tertiary but could be any age. The intruded rocks are carbonate and/or calcareous clastic rocks and can be of any age. All porphyry copper deposits have associated copper skarms that make up between 1 and 100 percent of the ore. Very few large copper skarms are formed when the host rock is wholly dolomitic.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: By their very nature, the skarns are very irregular in size. Porphyry copper related skarns commonly contain 25 to 385 million tons of open pit ore while nonporphyry skarns may contain only 35,000 tons to 10 million tons. Copper is found mainly in the form of chalcopyrite and lesser amounts of bornite. Grade typically ranges from 1 to 2 percent copper. Intermixed with the copper minerals are varying amounts of pyrite, hematite, magnetite, pyrrhotite, molybdenite, and sphalerite. Accessory minerals of calc skarns are andradite garnet, pyroxene, diopside, epidote, and wollastonite. In magnesium skarns, fosterite, brucite, tremolite, and serpentine commonly form. <u>Alteration</u>: In the porphyry subtype skarns, alteration can be extensive with a major potassic interzone grading out to sericite and phylite. Intruded rocks are frequently altered to epidote, pyroxene, diopside, and garnet. The nonporphyry subtypes are generally not as extensively altered but have the same mineral assemblages. Carbonate beds can be altered to coarse marble and calcareous beds to hornfels.

Ore Controls: The permeability and badding planes in the carbonate and/or calcareous rocks are of prime importance in the development of the skarn. Equally important are induced fractures as the result of the intrusive action.

Weathering: Surface erosion will produce copper carbonates and oxides while the iron minerals will frequently form a gossan.

<u>Geochemical Expression</u>: Normally, anomalous amounts of copper, lead, zinc, gold, silver, molybdenum, and possibly bismuth will be present in the general vicinity depending on the size of the deposit and the amount of surface exposure.

Geophysical Expression: Magnetic and/or electric methods could be used, but the results would depend on the amount of magnetite, pyrite, and pyrnotite in the skarn and the amount of cover if buried. Gravity methods might detect the presence of buried stocks but not whether there were associated skarns.

Associated Mineralization: Besides porphyry copper deposits, some zinc-bearing skarn may form and replacement deposits of lead and zinc may form in the outer marble zones of carbonate beds.

Deposit Examples:

Christmas (AZ) Lakeshore (AZ) Mission (AZ) Morenci (AZ) Pima (AZ) Silver Bell (AZ) Twin Buttes (AZ)

Empire (ID)

Blue Bell (MT) Elkhorn (MT)

Blue Stone (NV) Casting (NV) Copper Basin (NV) Douglas Hill (NV) Ely (NV) Ludwig (NV) Lyon (NV) Mason Valley-Malachite (NV) McConnell (NV) Victoria (NV) Western Newada (NV)

Continental (NM) McKnight and Fellows, 1978 Sant Pedro (NM) Santa Rita (NM) Snowshoe (NM)

Carr Fork (UT)

References:

Perry, 1969 South, 1972 Richard and Courtright, 1959 Moolick and Durek, 1966 Langlois, 1978 Graybeal, 1982 Barter and Kelly, 1982

Umpleby, 1917

Knopf, 1913 Klepper et al., 1957

Harris and Einaudi, 1982 Harris and Einaudi, 1982

Blake et al., 1978 Harris and Einaudi, 1982 James, 1976 Harris and Einaudi, 1982

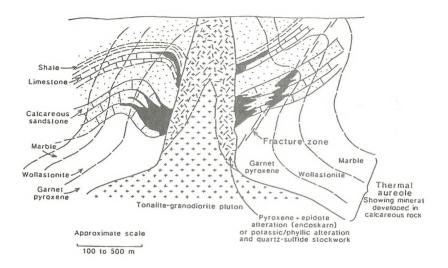
Einaudi, 1977 Harris and Einaudi, 1982 Atkinson et al., 1982 Harris and Einaudi, 1982

Einaudi, 19822Pinos Altos (NM)

Smith et al., 1945 Nielson, 1970

Atkinson and Einaudi, 1978

#2.04



Idealized section of a copper skarn (from USGS Open File Report 83-423)

Deposit Type: ZINC-LEAD SKARN

Subtype: Some distinction can be made between those formed near batholiths, near stocks, near dikes, and those distant for any igneous source. Differences are noted mainly on the basis of associated minerals and degree of alteration.

Synonym: Contact metamorphic, contact metasomatic, pyrometasomatic, or zinc-lead tactite.

Description: Sphalerite and galena in calc-silicates associated with carbonate and/or calcareous clastic rocks.

GEOLOGIC FEATURES

Regional Setting: Most deposits occur along continental margins and are related to syngenetic or late orogenic activity.

<u>Structural Features:</u> No particular structural features or orientation noted in most deposits. Those deposits associated with igneous bodies are probably related to zones of weakness that allowed emplacement, but evidence of such may or may not be discernible. One distinctive feature of this type of mineralization is that the skarns occur relatively distal to their magmatic source.

Stratigraphic and Lithologic Characteristics: A wide range of carbonate and calcareous sedimentary rocks can serve as a host for mineralization with no particular preference. Age can range from Paleozoic to Mesozoic. Intrusive rocks can vary from granites and granodiorites to diorite and syenite. Ages of these rocks can span from Paleozoic to Tertiary although most are Mesozoic. The apparent nonselectivity of intrusive and intruded rock types is another distinctive feature of this model.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Typical grades range from 3 to 14 percent zinc and usually lesser amounts of lead. The zinc/lead ratio is frequently 2:1 or less. Small amounts of copper and gold are nearly always present. Silver values are commonly in the 1-9 percent range. The ore minerals most often found are sphalerite, galena, pyrrhotite, magnetite, chalcopyrite, and arsenopyrite. The accessory minerals that make up the skarn are manganoan hedenbergite, andraditic garner, spessartine, bustamite and rhodonite. The presence of manganese-bearing minerals is distinctive to this type of skarn. Those deposits that form near contacts with batholithic intrusives tend to be smaller than other types of zinc skarn. The largest are those found near intrusive stocks. As a group, zinc-lead skarns range from 300,000 tons to 20 million tons in size.





<u>Alteration</u>: Intense but localized hydrothermal alteration of the intrusive causes development of epidote, and manganese-bearing pyroxenes and garnets similar to those found in the skarn portion. Alteration of the skarn shows up as development of manganerous actinolite, ilvaite, chlorite, and rhodochrosite.

Ore Controls: The skarn in most deposits occurs somewhat distal (300 feet or more) to the intrusive. Localization appears to occur along structural and lithologic contacts in the host rocks and/or along faults.

Weathering: A gossan may develop above the skarn zone if topography and past climatic conditions were favorable. Iron oxides and staining of the nearby rocks would be the most ovyious indicator.

<u>Geochemical Expression</u>: Above or immediately downslope from the skarn, anomalous amounts of silver, arsenic, beryllium, cobalt, copper, fluorine, lead, tin, and tungsten should be present. Because of its mobility, anomalous amounts of zinc will show up in the peripheral area of the skarn.

<u>Geophysical Expression</u>: Very little information is available on the signature this type of mineralization may have. Various electrical and/or magnetic methods might be applied but the response would depend upon the amount of sulfide minerals and magnetite present in the skarn, and the amount of cover.

Associated Mineralization: This model type is frequently associated with copper skarn. A zoning of the major silicates and sulfides minerals is frequently displayed. The deeper and central part of the skarns commonly exhibit garnet, magnetite, and chalcopyrite. Higher and laterally in the skarn this assemblage may change to pyroxene and sphalerite. The outermost zone may contain mainly galena.

Deposit Examples:

Aravaipa (AZ) Washington Camp (AZ)

Cooney Zinc (CA) Mineral King (CA)

South Mountain (ID)

Paymaster (NV)

Black Hawk (NM) Empire (NM) Ground Hog (NM) Linchburg (NM)

References:

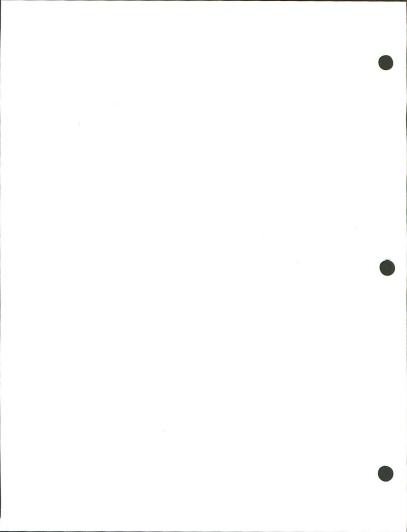
Reiter, 1981 Simon, 1972

Samson and Tucker, 1940 Knopf and Thelen, 1905

Sorenson, 1927

Gulbrandes and Gielow, 1960

Mosier, 1983d Hernon and Jones, 1968 Hernon and Jones, 1968 Titley, 1961



Deposit Type: TUNGSTEN SKARN

Subtype: It has been proposed that tungsten skarns be classified as either reduced or oxidized on the basis of host rock and depth of formation. In comparing actual deposits there is considerable overlap and the outward appearance of both subtypes is not that different.

Synonym Name: Contact metamorphic, contact metasomatic, pyrometasomatic, or tungsten tactite.

Description: Scheelite mixed with calc-silicate minerals that have formed in calcareous sediments at or near the contact with an intruding magma.

GEOLOGIC FEATURES

Regional Setting: Most deposits are related to the continental portions that were in late to postorogenic stages of activity.

Structural Features: This type of mineralization is related to stocks and batholiths that have intruded calcareous sediments. The intrusive is typical a coarse-grained porphyritic granodiorite or quartz monzonite of mid-Paleozoic to late Cretaceous age. No predominant type or orientation of faults or lineaments have been found associated with these deposits.

Stratigraphic and Lithologic Characteristics: Skarns most frequently occur in argilaceous carbonate rocks and intercalated carbonate-pelite or carbonatevolcanic sequences. It is common to find the skarn occurring in the lowest carbonate bed of a stratigraphic sequence. Age of the host rock can range from Precambrian to Triassic. The intrusive action will generally alter the sediments to hornfels and/or marble.

MINERAL CHARACTERISTICS

Deposit Features: Skarn is most often found at or near the intrusive Contact in stratiform structures or pods. Tungsten content is very erratic with most deposits ranging in grade from 0.3 to 1.5 percent tungsten oxide. Although a few de-posits may contain up to 25 million tons of ore, the majority are much smaller in size, having from 55,000 tons to 1 million tons. The main ore mineral is scheelite that always carries small amounts of molybdenum. Accessory minerals are chalcopyrite, pyrhotite, pyrite, magnetite, sphalerite, arsenopyrite and bismuthinite. Copper content may be high enough in some deposits to be considered a co-product. The skarn composition will vary somewhat depending on the host rock. Most contain mixtures of mainly pyroxene and garnet with lesser amounts of epidote, wollastonite, sphene, apatite, and actinolite.

<u>Alteration</u>: Generally, very little alteration of the intrusive takes place. <u>Nearly all alteration is evident in the intruded rocks where diopside, heden-</u> bergite, grossular and andradite garnets are formed. Outside of the skarn may lie a barren wollastonite zone and beyond that, marble. Ore Controls: The structure of the host rocks, the dip of the bedding relative to the intrusive contact, jointing and/or fractures of the intrusive and the configuration of the intrusive contact are important ore controls. Most favorable situations are where the intrusive intersects the limb of an anticline of favorable host rocks or a fold in the host rocks plunges steeply towards the intrusive contact.

Weathering: Oxidation of some of the iron-bearing minerals will result in staining of nearby rocks. The amount or degree will depend on past climatic conditions, topography, and amount of skarn exposed to weathering.

Geochemical Expression: Anomalous amounts of tungsten, copper, molybdenum, Dismuth, beryllium, Cin, fluorine, and niobium should be tested for near granitic contacts with calcareous beds.

<u>Geophysical Expression</u>: Magnetic and electrical methods could be useful, but their success would depend on how much magnetite, pyrite, and pyrrhotite is in the skarn. Regional aeromagnetics might outline buried plutons as magnetic highs which would be significant in areas where favorable carbonate rock are known to exist.

Associated Mineralization: Occasionally, tin and zinc skarn are associated with the tungsten mineralization but this is not commonly found.



Black Rock (CA) Pine Creek (CA) Strawberry (CA)

Tungsten Jim (ID) Yellow Pine (ID)

Calvert (MT) Lost Creek (MT)

Mills City (NV) Nevada-Scheelite (NV) Osgood Range (NV) Tem Piute (NV)

Iron Mountain (NM)

Milford (UT)

References:

Rinehart and Ross, 1956 Newberry, 1982 Nokleberg, 1981

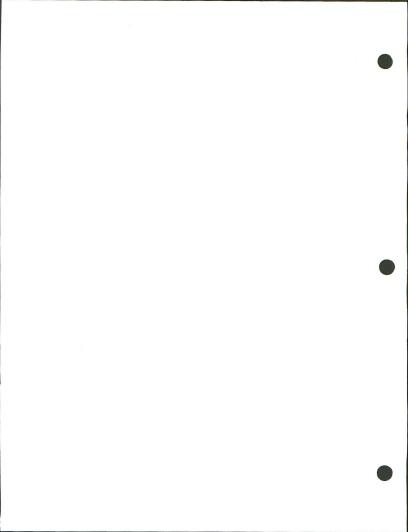
Sigurdson, 1974 Cooper, 1951

Geach, 1972 Collins, 1977

Kerr, 1934 Geehan and Trengrove, 1950 Taylor and O'Neil, 1977 Buseck, 1967

Jahns, 1944

Hobbs, 1945



Deposit Type: VOLCANOGENIC MASSIVE SULFIDES

Subtype: Cyprus

Synonym: Cupreous pyrite type

Description: Lenses of massive pyrite and copper minerals in mafic-ultramafic rocks. Name derived from the deposits on the island of Cyprus.

GEOLOGIC FEATURES

<u>Regional Setting</u>: These deposits are found along the orogenic belts of continental margins. Their position is believed to have resulted from seafloor spreading that subsequently uplifted and thrust fragments of the crust containing the deposits onto the continents during plate convergence.

Structural Features: Most of the major deposits are associated with pillow lavas or with volcanic breccias. Also, there seems to be an affinity with tensional graben-type faulting.

Stratigraphic and Lithologic Characteristics: Mineralization is most often found in the lower mafic volcanics and diabase dike complexes of ophiolites. Very few deposits that occur in clastic or pyroclastic strata are present in the volcanic sequence. Pelagic sedimentary rocks may cover the volcanic series. The age of these deposits range from late Precambrian to Mesozoic with most dated at Ordovician-Devonian or Cretaceous.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: The main mineralization is pyrite and chalcopyrite and less commonly pyrrhoitte. Minor amounts of marcasite and sphalerite may be present. Galena and bornite are rarely present. The gangue minerals are typically quartz, chlorite, carbonate, and gypsum. The sulfides occur most frequently as concordant tabular, lenticular, or saucer-shaped bodies in pillow lawas. A pipe or funnel-shaped stockwork zone of veins and copper dissemination may occur beneath the massive sulfide portion. Some deposit exhibit an iron-rich capping of ochre. Chert and manganese oxides may cap the ochre. Size of the deposits can range from a few thousand tons to 20 million tons but 2 to 5 million tons is more usual. Copper values commonly range from 0.5 percent to 4 percent, zinc from 0.1 to 3 percent, and gold and silver 0.1 to 2 ounces per ton.

<u>Alteration</u>: Mafic rocks are altered to zeolite and green-schist facies along with variable chloritization and/or silicification. Most of alteration is confined to the stockwork zone if present.

<u>Ore Controls</u>: The pillow basalt structure and brecciation of mafic volcanics appear to strongly influence the ore location. Pre-ore faulting and related fractures is important in some deposits. Weathering: Under most weathering conditions, a gossan will form above a massive sulfide deposit if exposed to weathering. The appearance of the gossan depends on the original sulfide mineralization, nature of the host rock, and the climatic conditions. In any case it will probably be some shade of tan or brown. Development of the gossan often mimics the size and size of the underlying massive sulfide bodies.

Geochemical Expression: Pillow lavas near points of mineralizations will frequently have anomalous amounts of zinc and cobalt and be depleted in copper. Lavas in the peripheral areas often have anomalous amounts of iron, sulfur, copper, zinc, and cobalt. Manganese and iron-rich cherts nearby are indicators of hydrothermal discharge and possible sulfide mineralization. Gold may be found in nearby stream sediments.

<u>Geophysical Expression</u>: Aeromagnetic surveys may be useful in locating volcanic centers that are often reflected as a negative anomaly. Massive sulfide bodies would reflect as positive anomalies, but the response would depend on the amount of pyrrhotite present, size and shape of the bodies, and amount of cover.

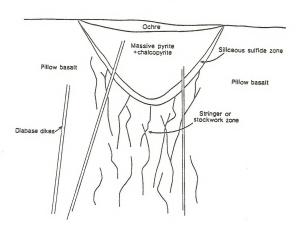
Associated Mineralization: None are known to occur with this type of mineralization, but occasionally chalcopyrite is subordinate to sphalerite. The zinc has an erratic distribution and may be only locally abundant.

Rua Cove (AK) Island Mountain (CA) Big Mike (NV) Turner-Albright (OR)

References:

Singer and Mosier, 1983 Stinson, 1957 Rye et al., 1984

Shenon, 1933



Idealized cross-section of a Cyprus type volcanogenic massive sulfide deposit (from USGS Open File Report 83-423)

Shale or phyllite and interbedded Fe-Mn-rich chert

Deposit Type: VOLCANOGENIC MASSIVE SULFIDES

Subtype: Polymetallic

Synonym: Lead-zinc-copper-silver type

Description: Lenticular to tabular bodies of massive pyrite with predominant lead, zinc, and copper minerals in felsic volcanic rocks.

GEOLOGIC FEATURES

Regional Setting: This type forms mainly in eugeosynclinal troughs bordering volcanic domes and cratonic areas. These are late stage activities believed to have resulted from seafloor spreading.

Structural Features: Usually tensional graben-type faulting is present. In most deposits, the rocks are medium to highly metamorphosed disguising somewhat the original structure and altering the character and texture of the rocks.

Stratigraphic and Lithologic Characteristics: The enclosing rocks are commonly thin volcanic sequences that grade upward from mafic, to intermediate, to felsic tholeitic basalt. Silici-alkalic lava, and pyroclastics are found associated with most deposits. Intercalated with the volcanics are sedimentary rocks, such as limestones, bedded sulfates, minor iron formations, graphic schists, and ferrugenous chert. The age of these deposits range from late Proterozoic to Tertiary. The younger deposits are generally smaller in size than the older deposits.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Mineralization is generally stratabound, irregular and lenticular to tabular in form. Pyrite is the dominant sulfide in mildly metamorphosed areas while pyrthotite occurs when highly metamorphosed. Lead occurs as galena, zinc as marmatitic sphalerite, and copper as chalcopyrite. Lead-rich ores carry important amounts of silver with lesser amounts of gold. Quartz is the most common gangue mineral, in fact mineralization generally is more intense in the more silicic host rock. Lead content commonly ranges from 1 to 3 percent, zinc 3 to 8 percent; and copper 0.5 to 2.0 percent. Silver content is usually 1 to 3 ounces per ton. These deposits are relatively small, ranging from 100,000 tons to 1.5 million tons.

<u>Alteration</u>: The more intensely altered areas are generally at the bottom or "root zone" of mineralization. They may consist of chloritization, sericitization, silicification, and tourmalinization. Areas lateral to the deposit are commonly chloritized and sericitized. Ore Controls: Mineral zoning is common with lead at the top and grading downward into xinc and copper. Mineralization contacts are frequently sharp in the upper areas of the deposit with veins and venlets developing in the lower part or bottom of the deposit. The ore minerals commonly grade laterally into the host rock resulting in an iron sulfide halo several times larger than that of the ore body. The iron sulfide content usually increases toward the ore body. The lenses of mineralization rarely occur alone but are found in clusters or in boudin-form.

Weathering: If exposed, gossan will generally develop. Most often it is lenticular in shape and yellow, red or brown in color. The appearance depends on the original sulfide mineralization, nature of the host rocks, and climatic conditions.

<u>Geochemical Expression</u>: Lead is frequently anomalously high in the gossan whereas zinc and copper values are usually not present. Detectable amounts of gold in the gossan are indicative of polymetallics beneath the capping rather than barren iron sulfides. The halo of iron sulfides surrounding many deposits is usually anomalously low in copper, lead, zinc, gold, and silver values.

Geophysical Expression: Electromagnetic methods are especially suited to detecting this type of deposit because of the high pyrite and/or pyrthotite content of the ore. Detection depends on the size and shape of the ore body, and the amount and type of cover. Aeromagnetics might be useful in locating felsic volcanic centers that would be indicated by negative anomalies. Massive sulfide deposits under favorable conditions would be indicated by positive anomalies.

Associated Mineralization: In a few deposits gold-quartz veins are known to occur beneath but at a considerable distance from massive sulfide deposits of this type.

#3.01B

Deposit Examples:

*Arctic (AK) *Beatson (AK) *Greens Creek (AK) *Orange Point (AK) *Antler (AZ) *Binghampton (AZ) *Bruce (AZ) Iron King (AZ) Afterthought (CA) Big Bend (CA) *Blue Ledge (CA) Blue Moon (CA) Bully Hill-Rising Star (CA) Copper Hill (CA) *Gray Eagle (CA) Keystone-Union (CA) Newton (CA) North Keystone (CA) Penn (CA) Quail Hill (CA) Spencerville (CA) *Chestatee (GA)

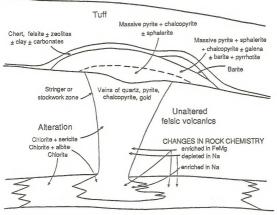
*Bald Mountain (ME) *Barrett (ME) *Big Hill (ME) *Deer Isle (ME) *Emerson (ME) *Tapley (ME) *Davis (MS)

*Pecos (NM)

*Flambeau (WI)

References: Mosier et al., 1983 Bateman, 1924 Mosier et al., 1983 Mosier et al., 1983 Anderson and Guilbert, 1979 Anderson and Guilbert, 1979 Anderson and Guilbert, 1979 Gilmore and Still, 1968 Albers and Robertson, 1961 Eric, 1948 Mosier et al., 1983 Heyl and Cox, 1948 Albers and Robertson, 1961 Mosier et al., 1983 Mosier et al., 1983 Hey1, 1948(a) Heyl and Eric, 1948 Hey1, 1948(a) Hey1 et al., 1948 Hey1, 1948(b) Mosier et al., 1983 Mosier et al., 1983 Mosier et al., 1983 Young, 1963 Young, 1963 Mosier et al., 1983 Mosier et al., 1983 Earl, 1950 Mosier et al., 1983 Mosier et al., 1983 Mosier et al., 1983

*Geologic information is insufficient to classify as deposit type, may be either #3.01B or #3.01C.



Unaltered intermediate volcanics

Felsic igneous dome

Idealized cross-section of a polymetallic type of a volcanogentic massive sulfide deposit (from USGS Open File Report 83-423)

#3.01C

Deposit Type: VOLCANOGENIC MASSIVE SULFIDES

Subtype: Zinc-copper

Synonym: Primitive or Kuroko type

Description: Lenticular deposits of zinc, copper, and massive iron sulfides associated with Precambrian and early Paleozoic marine volcanic rocks.

GEOLOGIC FEATURES

Regional Setting: These deposits are associated with very large scale eugeosynclinal volcanism occurring in mainly very early stage of a tectonic cycle.

Structural Features: No distinctive type of structural features are known to be associated with these deposits. General high-angle normal faults are nearby as expected in a synclinal environment.

Stratigraphic and Lithologic Characteristics: Most occurrences are found in a thick sequence of volcanic rocks several tens of thousand feet thick. Mafic rocks made up the bottom of the pile with a gradation to intermediate and felsic types near the top. Sedimentary iron formations, ferruginous cherts, and volcanogenic iron and silica-rich graywacks are frequently intercalated with the volcanic layers. The age of the majority of deposits is late Precambrian to Devonian, but a very few deposits are dated as late Mesozoic and early Genozic. The larger deposits are generally the oldest.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Deposits of this type are frequently in irregular lenticular or tabular shapes. Copper, zinc, and gold are the primary elements dispersed in pyrite or pyrrhotite. They take on a massive or layered appearance depending upon the sulfide content. Mineralization is frequently found associated with the more silicic volcanic rocks. Ore minerals are most often sphalerite and chalcopyrite and rarely galena. More gold is found in copper rich deposits whereas high silver is associated with zinc-rich deposits. Most of the known deposits range in size from 100,000 to 20,000,000 tons. Copper grade ranges from 0.40 to 3.50 percent; zinc, 2 to 9 percent. Gold values are commonly in the 0.003 to 0.07 ounces per ton range, and silver ranges from 0.35 to 3.00 ounces per ton.

<u>Alteration</u>: At the base of the deposits alteration is usually more intense. It consists of chloritization, sericitization, silicification, and tourmalinization. The sides of the deposit may have chlorite and sericite. Some of these products may have resulted from regional metamorphism rather than alteration.







<u>Ore Controls</u>: A predominate number of deposits are located near centers of felsic volcanism such as felsic domes, calderas, and volcanic necks. Rocks may be locally brecciated near the centers. Mineral zoning is sometimes apparent with zinc at or near the top of the deposit and grading downward toward increasing copper. The upper contact of these deposits is sharp whereas below the main mineralized, veins may occur. A halo of iron sulfides is often found around the deposit.

Weathering: If past weathering conditions were favorable, a gossan will most likely form over the deposits. The appearance of the gossan will depend on the original sulfide minerals, the host rock, and the climatic conditions. The most common colors are various shades of brown and dark red.

<u>Geochemical Expression</u>: Detectable amounts of gold in the gossan are generally indicative of base metals beneath the gossan. Copper and zinc usually are depleted in the gossan area, but sometimes are found as a halo region along with anomalous amounts of gold and silver.

<u>Geophysical Expression</u>: Electromagnetic methods would probably be the most discriminating to use, but success would depend on depth, type of cover, shape of the deposit, and amount of iron sulfide present. Aeromagnetics might locate felsic centers, often reflected as negative anomalies, whereas sulfide bodies could show up as positive anomalies if enough pyrrhotite is present.

Associated Mineralization: Gold in quartz veins is sometimes found in the root zone beneath deposits of this type.

#3.01C

Deposit Examples:

*Artic (AK) *Beatson (AK) *Greens Creek (AK) *Orange Point (AK) *Antler (AZ) *Binghampton (AZ) *Bruce (AZ) United Verde (AZ) Balaklala (CA) *Blue Ledge (CA) Early Bird (CA) *Gray Eagle (CA) Iron Mountain (CA) Keystone (CA) Mammoth (CA) Shasta King (CA) Stowell (CA) *Chestatee (GA) *Bald Mountain (ME) *Barrett (ME) *Big Hill (ME)

*Deer Isle (ME) *Emerson (ME) *Tapley (ME)

*Davis (MS)

*Pecos (NM)

*Flambeau (WI)

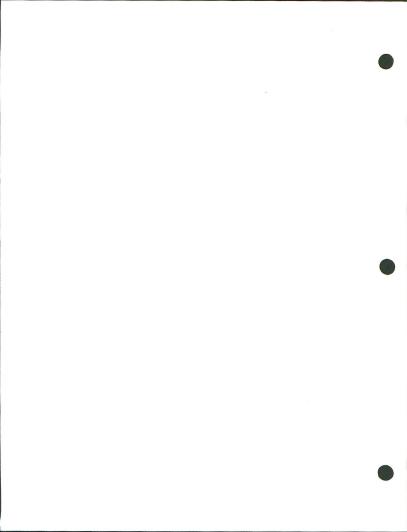
References:

Mosier et al., 1983 Mosier et al., 1983 Mosier et al., 1983 Mosier et al., 1983

Anderson and Guilbert, 1979 Anderson and Guilbert, 1979 Anderson and Guilbert, 1979 Anderson and Creasey, 1958

Kinkel et al., 1956 Mosier et al., 1983 Kinkel et al., 1956 Mosier et al., 1983 Kinkel et al., 1956 Kinkel et al., 1956 Mosier et al., 1983 Kinkel et al., 1956 Kinkel et al., 1956 Mosier et al., 1983 Mosier et al., 1983 Young, 1963 Young, 1963 Mosier et al., 1983 Mosier et al., 1983 Earl, 1950 Mosier et al., 1983 Mosier et al., 1983 Schwenk, 1977

*Geologic information is insufficient to classify as to deposit type; may be either #3.01B or #3.01C.



Deposit Type: VOLCANOGENIC MANGANESE

Synonym: Ophiolitic manganese

Description: Massive lenses of manganese oxide, carbonate, and silicate in intercalated sequences of volcanic and sedimentary rocks.

GEOLOGIC FEATURES

<u>Regional Setting</u>: Most deposits of this type are believed to have formed in an eugeosynclinal subduction complex. More explicitly, they are thought to have resulted from seafloor hot spring systems that may have formed in marginal ocean basins, at the base of island arcs, oceanic ridges, or at the base of oceanic islands.

Structural Features: Regional rocks are commonly basaltic lava that show pillow structure indicative of underwater deposition.

Stratigraphic and Lithologic Characteristics: Host rocks are usually thinbedded red radiolarian chert intercalcated with basalt, greenstones, tuff, and graywacke. Varying amounts of metamorphism is nearly always present distorting the bedding features and altering the associated minerals. Age of the enclosing rocks varies from Cambrian to Plicoene.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Typically, the ore occurs as massive black lenses of poorly crystalline or amorphous manganese oxides, carbonates, and manganiferous chert. The lenses are generally circular or elliptical in shape having a thickness from a few inches to 60 or 90 feet and a length extending 650 to 900 feet. The ore bodies can range from less than 3,000 to several million tons in size. It is not uncommon to find two, three, or more lenses stacked more or less above one another and separated by 5 to 30 feet of thin bedded chert. The enclosing chert is frequently red jasper but may range in color from white to green or brown. Beneath the cherts, basalts or greenstones are usually found. Ore minerals are frequently a mixture of rhodochrosite, pyrolusite, hausmannite, braunite, rhodonite, bementite, neotocite, and inesite. In the more massive portions of the lenses, 30 to 50 percent manganese may be found. At the edges of the lenses, the silica content increases to form manganiferous chert.

<u>Alteration</u>: The originally deposited minerals are thought to have been manganiferous opal, carbonate, and oxides. Through later supergene and metamorphic activity, these minerals were altered to the present mineral suite. Alteration of the surrounding basalts to greenstones is common along with silification and the injection of ferriferous solutions. <u>Ore Controls</u>: Faults and fractures in permeable sequences of beds in the subsea floor allowed hydrothermal solutions to Circulate. Manganese was deposited at the seafloor-seawater interface as the result of reduction and oxidation of the ascending solutions. Later supergene enrichment probably upgraded the manganese content.

Weathering: Under most weathering conditions secondary manganese oxide minerals of birnessite, pyrolusite, todorokite, and amorphous manganese oxide may form.

Geochemical Expression: Anomalous amounts of barium, lithium, rubidium, zinc, lead, and copper can be expected near deposits of this type.

Geophysical Expression: None have been documented.

Associated Mineralization: None definitely identified. It has been speculated that the mode of deposition might cause deposition of iron-rich sulfides at depth beneath the magnaces accumulations.

#3.02

Deposit Examples:

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Blue Jay (CA)
Buckeye (CA)
Commings (CA)
Fablan (CA)
Foster Mountain (CA)
Ladd (CA)
Liberty (CA)
South Thomas (CA)
Thatcher Creek (CA)
Thomas (CA)
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Black Diablo (NV)

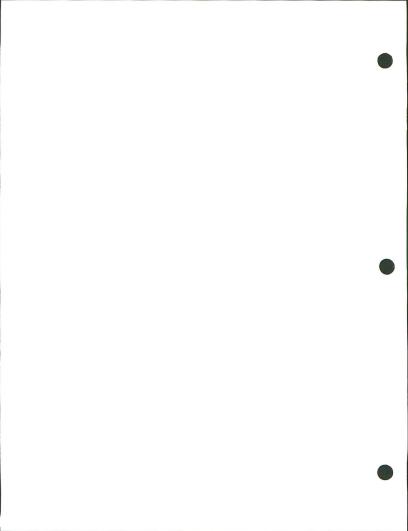
Beaver Falls (WA) Brown Mule (WA) Cresent (WA) Dosewallipo (WA) Hurricane (WA) Littleton (WA) Skunk Creek (WA)

References:

Crerar et al., 1982 Taliaferro and Hudson, 1943 Trask et al., 1943 Trask et al., 1943 Trask et al., 1943 Trask et al., 1943 Crerar et al., 1942 Trask et al., 1943 Trask et al., 1943

Needham and Trengrove, 1950

Sorem and Gunn, 1967 Park, 1942 Park, 1942 Sorem and Gunn, 1967 Park, 1942 Sorem and Gunn, 1967 Park, 1942



Deposit Type: PRECIOUS AND BASE METAL REPLACEMENT

Subtype: Limestone replacement

Synonym: Manto or metasomatic deposits

Description: A hydrothermal, epigenetic, sulfide mineral deposit that commonly replaces selective beds of limestone, dolomite, or other soluble rock.

GEOLOGIC FEATURES

Regional Setting: Most deposits have formed in a plutonic environment spatially related to batholiths. The carbonate host rocks commonly occur in broad sedimentary basins and are moderately deformed.

<u>Structural Features</u>: Thin bedded and brittle rock sequences that are folded, fractured, and faulted are especially favorable because they present channelways for ascending mineral solutions. The majority of the deposits are found near the borders of batholiths in the contact zones of satellite plutons, or in roof pendants in plutons.

Stratigraphic and Lithologic Characteristics: The associated intrusives are predominantly calc-alkaline of the granodiorite or quartz monzonite type. They frequently are equigranular and contain late dikes and quartz veining. Most, but not all, are late Mesozoic to early Cenozoic in age. The sedimentary host rocks are chiefly limestones, but can also be dolomite and shales. In some deposits only the more pure limestone is replaced while in others impure carbonates and/or siliceous clastic beds are favored. Sometimes a particular layer in an otherwise indistinguishable sequence of rock is the favorable replacement bed. The sediments can be any age but predominantly are Paleozoic.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: The deposits may occur in many forms, but are usually irregular or poditke in shape. Occasionally, ribbonlike or blanketlike deposits will form along selective layers or beds. A great variety of ore minerals can occur in this type of deposit but never more than two or three at one time. Some of these are galena, sphalerite, argentite, tetrahedrite, enargite, chalcopyrite, proustite, pyrargyrite, jamesonite, bournonite, tennantite, jordanite, stephanite, polybasite, sylvanite, calaverite, native gold, blamuthinite, and rhodochrosite. Accessory minerals can be pyrite, marcasite, barite, quartz, arsenopyrite, garnet, wollastonite, fluorite, and calcite.

Deposits of this type range widely in size from 250,000 tons to 15 million tons. For those deposits containing predominantly copper and/or gold, the grade is generally high but extremely variable. Gold frequently grades from one to two ounces/ton while copper often ranges from 5 to 20 percent. Deposits containing lead and zinc often carry minor but important amounts of gold and silver.

<u>Alteration</u>: Very little alteration is generally evident in the intrusive rock and when present is very narrow and at the contact. Alteration of carbonates is evident by bleaching, recrystallization to marble, and development of some calc-silicate minerals such as scapolite, idocrase, or tremolite. Alteration is usually gradational away from the intrusive contact.

<u>Ore Controls</u>: The configuration of the intrusive and the attitude of the host rocks are extremely important in emplacement of the deposits. One of the more favorable situations is where the intrusive intersects the limb of an anticline or fold of favorable host rocks thereby allowing mineralizing solutions to travel upward along bedding planes. Joints and fractures in the host and/ or intrusive also aids migration of the fluids and induces chemical reactions where possible.

Weathering: No unique weathering features are generally present. Iron minerals may oxidize to ocherous masses containing oxides of copper, lead, zinc, and silver.

<u>Geochemical Expression</u>: Many deposits exhibit a zonal arrangement of anomalous copper at the center followed by lead-silver, and a zinc and manganese zone at the fringes. Other elements that are frequently associated with this type of deposit are bismuth, arsenic, fluorine, and molybdenum.

<u>Geophysical Expression</u>: Generally, most geophysical methods will not detect these deposits directly because of the small target size. Regional aeromagnetics have been used successfully in detecting buried plutons in areas where favorable carbonate rocks are known to exist. These usually appear as magnetic highs.

Associated Mineralization: Some precious and base metals occur in veins in and/or near the contact zone.



Kennecott, AK

Christmas, AZ Bisbee, AZ Pima, AZ

Gilman, CO Leadville, CO Ouray, CO

Phillipsburg, MT

Eureka, NV

Homestake, SD

Bingham, UT Park City, UT Tintic, UT References:

Bateman, 1942

Willden, 1964 Hogue and Wilson, 1951 Irvin, 1959

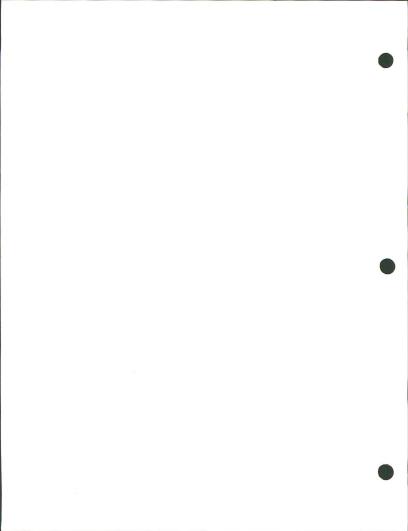
Tweto and Lovering, 1947 Emmons et al., 1927 Burbank, 1968

Prinz, 1961

Nolan and Hunt, 1968

McLaughlin, 1931

Rubright and Hart, 1968 Barnes and Simos, 1968 Morris and Lovering, 1979



Deposit Type: CARBONATE-HOSTED GOLD

Subtype: Disseminated gold

Synonym: Carlin type or colloidal gold

Description: Very fine colloidal gold and sulfides disseminated in carbonaceous carbonate rocks.

GEOLOGIC FEATURES

Regional Setting: Eugeosynclinal carbonates deposited in a somewhat reducing environment that later becomes tectonically active.

Structural Features: Moderate regional folding of the sediments is common with many high angle normal faults. Dikes and small plutons are nearly always found in the immediate vicinity.

<u>Stratigraphic and Lithologic Features:</u> Favorable host rocks are limestones, dolomites, and siltstones containing organic carbon and pyrite. Age of the host rocks does not seem to be a critical favor. The associated intrusives range from quartz diorite to quartz monzonite and are frequently porphyritic. They are mainly Tertiary in age but can range to early Cretaceous.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Microscopic and submicroscopic gold occurs as tiny seems and veinlets with quartz. It is also found filling fine fractures in pyrite and to a lesser extent as precipitations on organic carbon. The gold particles range from 10 to 0.5 micron in size and commonly grade from 0.03 to 0.30 ounces/ton. The deposits may range in size from two million to 25 million tons. Most of the ore deposits are stratiform and lie in a zonal pattern that may encircle chimney areas of silification. The ore frequently exhibits a spatial relation to the intersection of high-angle faults although the faults themselves are rarely mineralized. Most of the deposits have oxidized and unoxidized ores. Oxidized ores contain iron oxides, clays, microscopic quartz, arsenic, antimony, and mercury sulfides. Unoxidized ores are higher in pyrite, organic carbon, arsenic sulfides, and barite.

Alteration: Hydrothermal solutions usually bleach and leach large areas. Both the host rocks and the associated igneous rocks were affected, but sediments show the greatest change. Widespread replacement of carbonate by silica is the most common exhibited feature. In the unxidized ore zones, jasperoid, quartz, illite, kaolinite, and calcite are found. The oxidized ores exhibit kaolinite, montmorillonite, illite, jarosite, and alunite.



<u>Ore Controls</u>: The principal control features were tectonic folding and fracturing of brittle sediments. The hydrothermal solutions selectively replaced carbonaceous carbonate beds adjacent to high angle faults and regional thrust faults.

<u>Weathering</u>: No particular distinguishing features are found at most deposits. The widespread alteration usually found may change selective beds or layers to light reddish gray or tan. Light brown to reddish brown stained jasperoid is also usually present.

Geochemical Expression: Most deposits exhibit anomalous amounts of arsenic (400 ppm), antimony (100 ppm), mercury (25 ppm), and thallium (50 ppm). Barite is commonly present along with higher than normal amounts of molybenum, tungsten, and fluorine.

<u>Geophysical Expression</u>: No distinctive pattern has been recognized over this type of deposit. The amount of disseminated pyrite in the unoxidized ore zone is low enough that electrical methods probably could not detect the mineralization.

Associated Mineralization: Small prospects of silver, lead, zinc, copper, and mercury sometimes occur in the near vicinity.

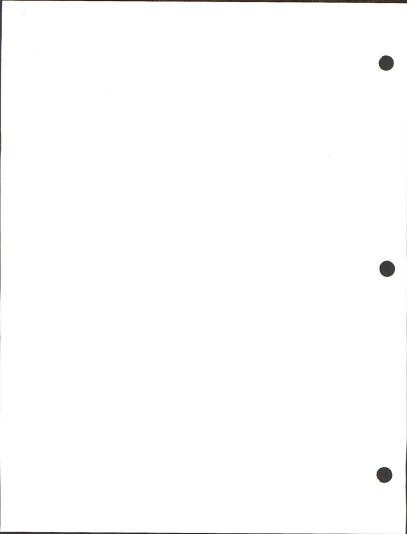
Alligator Ridge (NV) Carlin (NV) Cortez (NV) Getchell (NV) Gold Quarry (NV) Jerritt Ganyon (NV) Maggie Creek (NV) Northumberland (NV) Pinson (NV) Preble (NV) Santa Fe (NV)

Mecur (UT)

References:

Stanford, 1984 Hausen and Kerr, 1968 Wells et al., 1969 Joralemon, 1951 Hausen and Kerr, 1968 Jackson, 1982a Hausen and Kerr, 1968 Kral, 1951 Jackson, 1982b Jackson, 1982b Clark, 1922

Butler et al., 1920



Deposit Type: LOW-SULFIDE GOLD-QUARTZ VEINS

Synonym: Mother Lode type

Description: Gold in massive but discontinuous quartz veins and lodes mainly in regionally metamorphosed sediments and volcanics

GEOLOGIC FEATURES

<u>Regional Setting</u>: Deposits are found along the mobile belts of the continental margins, namely the western foothills of the central Sierra Nevada Mountains (Nother Lode region) and the southeastern front of the Appalachian Mountains (Piedmont region).

Structural Features: Orogenic stresses caused complex folding and overturning of beds with accompanying fractures and faulting. Much of the mineralization is associated with thrust faults of small displacement; normal faults are also present. Shear zones along northwest or northeast directions have created secondary tension cracks.

Stratigraphic and Lithologic Characteristics: The principal rocks are slates, schists, greenstones, quartzites, and gneisses. Very few of the original Precambrian to Tertiary sediments have escaped regional metamorphism. Age of the rocks ranges from early Paleozoic to mid-Mesozoic. Later intrusions of peridotite, diabase, gabbro, granodiorite, quartz diorite, and small dikes of various types have invaded the metamorphosed sediments. Great belts of peridotite have been altered to serpentine that is partly associated but not exclusively with the mineralization.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: The gold deposits occur either as quartz veins or as bodies of mineralized country rock. The quartz veins generally occur as systems of parallel or acutely intersecting veins. Not many can be traced for more than a few thousand feet. They cut the enclosing rocks at an acute angle in both strike and dip. The veins swell and pinch abruptly filling fissures that were formed mainly from reverse faulting. Banding or ribboning of the quartz is common when the veins traverse schist or slate.

The ore bodies of mineralized country occur either adjacent to quartz veins or in broad zones of fissuring. Many of these ore bodies formed either in the footwall or hanging wall of a large thick barren quartz vein. The gold in both types of deposits is dispersed as small blebs and/or veinlets and ranges in grade from 5 to 40 grams per ton. Ore bodies commonly vary from 15,000 to 1.5 million tons in size. Mineralization is usually accompanied by ankerite, sericite, fine-grained pyrite, galena, sphalerite, and arsenopyrite. Galena and petzite are indicators of good ore.

<u>Alteration</u>: Carbonization (or ankeritization) is the chief alteration feature and is present in all rocks regardless of type. Ankeritized serpentine takes on a pale green color due to the formation of a chromium potassium mica (mariposite). Sericite, albite, pyrite, and arsenopyrite are often formed by the alteration of the wall rocks.

<u>Ore Controls</u>: Mineralization appears to have been controlled by faulting and joints. The country rock had little effect on deposition although slates seem to be more favorable than greenstones, and serpentine is almost absent of gold mineralization. In general, the more brittle rocks are the most favorable.

<u>Weathering</u>: No distinctive features have been noted in the mineralized areas. Some of the early rich placers developed above and downslope from mineralized zones as the result of weathering of the country rock and concentrating the gold in place.

<u>Geochemical Expression</u>: Arsenic has been found to be the most persistent pathfinder for this type of mineralization. Other elements are less reliable. In the southern Piedmont area, there is a general regional zonal arrangement. Gold is found nearest the magmatic source and then grades outward to a zone characterized by chalcopyrite, pyrite, and pyrrhotite and then a zone of galena, sphalerite, and barite.

Geophysical Expression: No specific signature is known to be connected with this type of mineralization. Some success has been obtained using aeromagnetics to map ultrabasic rock units in the Mother Lode area, but no correlation was found between magnetic patterns and known mineral deposits.

Associated Mineralization: Some massive sulfide copper deposits are found in the general region. Nearly all of the deposits contain small amounts of associated copper, silver, lead, and zinc, but may be too low a grade to recover.

Dutch Bend, AL Hog Mtn., AL

Angels Camp, CA Carson Hill, CA Grass Valley, CA Jackson-Plymouth, CA Jamestown, CA Souls byville, CA

Barlow, GA Battle Branch, GA Findley Ridge, GA

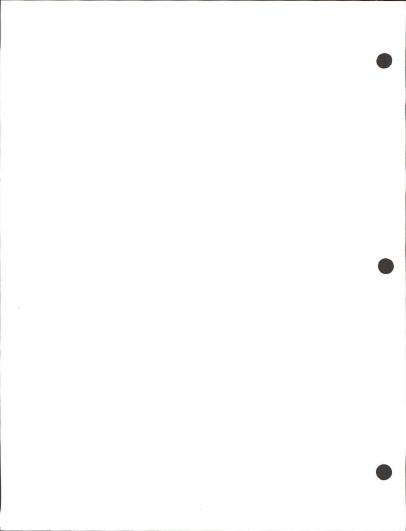
Coggins, NC Gold Hill, NC Howie, NC Lewis, NC Parker, NC Phoenix, NC Rudisil, NC

Bar Kat, SC Brewer, SC Dorn, SC Haile, SC

Franklin, VA Melville, VA References:

Pardee and Clark, 1948 Pardee and Clark, 1948 Eric et al., 1955 Knopf, 1929 Johnston, 1940 Knopf, 1929 Eric et al., 1955 Knopf, 1929 Pardee and Clark, 1948 Pardee and Clark, 1948





Deposit Type: EPITHERMAL GOLD AND SILVER

Subtype: Quartz-adularia

Synonym: Precious and base-metal veins; Comstock epithermal veins

Description: Precious metals in vuggy quartz and adularia veins with abundant pyrite and arsenopyrite and lesser amounts of galena and sphalerite.

GEOLOGIC FEATURES

Regional Setting: Most of the major deposits are associated with orogenic activities found in the subduction zone along continental margins and to a lesser extent with back-arc basins.

<u>Structural Features</u>: The most important districts occur where there are strong, persistent fracture systems such as Basin and Range-type faults, caldera ring fracture zones, caldera-related graben structures, or complex faulting or domal areas. The structures may be any age, but some of the more productive deposits are related to Tertiary age thermal events.

Stratigraphic and Lithologic Characteristics: There is generally no favorable nor preferential rock type in most deposits. Deposition is mainly related to brittleness of the related rocks when deformed. Brittle rocks form more cavities and solution paths when folded or faulted than more pliable units thereby enhancing deposition. Carbonate rocks are occasionally more favorable for mineralization depending on the character of the mineralizing solutions.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: The majority of the deposits occur as veins, stockworks, and replacements in volcanic rocks and as replacements and veins in sedimentary sequences. The ore frequently exhibits colloform banding, crustifications, open-space filling, and comb structure. Mineralization can take the form of native gold or silver; tellurides; and sulfides of arsenic, antimony, silver, mercury, and base-metals. Quartz and various amounts of calcite, fluorite, and barite always accompany these ore minerals. Dark or banded quartz are good indicators of nearby ore while white or "bull" quartz is usually found in unproductive portions of a vein or ore body. The size of the deposits can range from over one ounce per ton to a few parts per million. It is not uncommon to find ore shoots within ore bodies that are extremely rich. Silver content can range from 0.5 ounce per ton to 16 ounces; copper 0.1 to 1.0 percent; jacd 0.01 to 5.0 percent.

<u>Alteration</u>: During the mineral formation stage large amounts of quartz are involved which frequently results in a silicification or jasperoid some above the ore zone. Acid-leach alteration may occur above and below the zone in the form of argillization or quartz-tillite. Kaolinite, montmorillonite, alunite, and zeolites can also occur as irregular patches depending on the host rock







type. The alteration may be widespread and extend for considerable distances away from the productive zone. Along vein types deposits, quartz and adularia are nearly always found, but the amount may be variable.

<u>Ore Controls</u>: Deposits of this type require open spaces provided by a throughoing braided fracture system. A sudden change in attitude of a fault or an intersection with another fault will nearly always cause a change in the mineralization. In wide quartz veins, the ore may form either in the footwall or the hanging wall of the vein system. Many deposits show a gradual decrease in precious metal values downward on the vein and a general increase in base metal.

Weathering: Areas of bleached country rock may be present that may or may not be accompanied with a gossan cap. Clay material is often present in flat areas protected from erosion. Weathering of the upper part of a deposit may concentrate gold into rich residual placer deposits.

<u>Geochemical Expression</u>: The most useful and consistent elements accompanying this type of mineralization are arsenic, antimony, and mercury. Closer to the center of the deposit, anomalous amounts of gold and/or silver are sometimes exhibited depending on the level of erosion of the deposit. Anomalous amounts of lead and zinc can indicate that just the roots of the mineral system are all that are left in the erosion cycle.

<u>Geophysical Expression</u>: No distinct gravity or magnetic signature has been found associated with the ore itself, but both of these methods can be used to infer favorable associated structures.

Associated Mineralization: The erosion of high grade deposits can form rich gold placers for several miles downslope from the original deposit.

Katherine (AZ) Kofa (AZ) Oatman (AZ) Sheep Tanks (AZ)

Bodie (CA) Calistoga (CA) Monitor (CA)

Bonanza (CO) Creede (CO) Telluride (CO)

Flathead (MT)

Aurora (NV) Bellehelen (NV) Bovard (NV) Bullfrog (NV) Bruner (NV) Como (NV) Comstock (NV) Cornucopia (NV) Divide (NV) Eagle (NV) Eastgate (NV) Fairview (NV) Gold Circle (NV) Jarbidge (NV) National (NV) Rawhide (NV) Sand Springs (NV) Searchlight (NV) Seven Troughs (NV) Silver City (NV) Tonopah (NV) Tuscarora (NV) Wonder (NV)

Mogollon (NM) Nogal (NM) Steeple Rock (NM)

Blue River (OR)

Gold Mountain (UT) Gold Spring (UT) State Line (UT)



Republic (WA)

References:

Lausen, 1942 Jones, 1916 Lausen, 1942 Wilson et al., 1967

Silberman et al., 1972 Hanks and Irelan, 1887

Burbank, 1932 Stevens and Ratte, 1960(a) Burbank, 1941

Ross, 1960

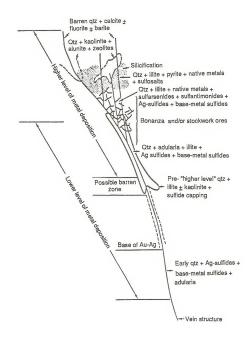
Vanderberg, 1937 Spur, 1903 Ross, 1961 Kral, 1951 Kral, 1951 Stoddard and Carpenter, 1950 Becker, 1882 Granger et al., 1957 Knopf, 1921 Hill, 1916 Vanderburg, 1940 Greenan, 1914 Rott, 1931 Schrader, 1923 Lindgren, 1915 Vanderburg, 1937 Vanderburg, 1940 Callaghan, 1939 Vanderburg, 1936 Stoddard and Carpenter, 1950 Bonham and Garside, 1979 Nolan, 1936 Vanderburg, 1940 Ferguson, 1927

Lindgren et al., 1910 Lindgren et al., 1910

Callaghan and Buddington, 1938

Callahan, 1938 Butler et al., 1920 Butler et al., 1920

Lindgren and Bancroft, 1914



Idealized cross-section of an Au-Ag-base metal deposit (from USGS Open File Report 83-423) Deposit Type: EPITHERMAL GOLD AND SILVER

Subtype: Quartz-alunite

Synonym: Enargite gold

Description: Gold, pyrite, and enargite in vuggy veins and breccia within zones of advanced argillic alteration related to felsic volcanism

GEOLOGIC FEATURES

<u>Regional Setting</u>: Although these deposits are found in a variety of tectonic settings, they are most commonly found in island arcs and back-arc spreading areas.

<u>Structural Features</u>: A throughgoing fracture system or ring fracture zones of caldera or domal structures are commonly associated with this type of mineralization. In some deposits, graben structures and/or normal faults are the primary feature.

Stratigraphic and Lithologic Features: The deposits are found associated mainly with intermediate calc-alkalic rocks. Most common are rhydacites, trachyandesites, or quartz latites that are nearly always porphyritic having large phenocrysts. Occasionally, sedimentary volcanic clastics are the mineral host, but intermediate intrusive and/or extrusive volcanics are nearly always found in the near vicinity.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Mineralization forms in fault breactas and/or fill openspace fractures as the result of breakage of advanced argillic alteration. The ore bodies are frequently irregular, crudely pipe-like, or tabular in shape. They are also often porous and vuggy. Metal values usually range from 0.04 to 0.3 ounces/ton for gold and 0.12 to 2.1 percent for copper. High grade ore often exceeds one ounce of gold and one percent copper/ton. Most deposits range in size from 50 thousand to 5 million tons. The ore minerals are usually enargite, tetrahedrite, and in some cases silver sulfogalits. Gold is probably present in ubiquitous pyrite although occasionally native gold is found. Varying amounts of bismuthinite, tellurides, chalcopyrite, galena, and sphalerite occur in some deposits. Gangue minerals are usually quartz and varying mounts of bismuthinite, and solinite, lilite, and montmorillonite. Most, if not all known deposits, are believed to be Cenozoic in age, but can be any age

<u>Alteration</u>: Pervasive argillic alteration in the mineral zone and in the surrounding host rocks is characteristic of this model type. Quartz, alunite, kaolinite, illite, and montmorillonite are common alteration assemblages that form around feeder conduits for hydrothermal solutions. Not all of these



minerals may be present in an individual deposit or district. Which ones develop depend on intensity of leaching, total sulfur concentration, and amount of alkalis in the hydrothermal solution. Propylitized rocks may be found ringing the argillic zone or as patch occurrences at the outer edges. Ore bodies are always found in or adjacent to the more advanced argillic zones.

Ore Controls: Mineralization and the accompanying alteration always is found along highly permeable structured features such as faults or stratigraphic features i.e., coarse clastic or volcaniclastic beds, flow breccias, or lithologic contacts. Original rock type appears to influence alteration patterns only through permeability, not compositional control, and has no significant influence on localization of the ore. Some deposits exhibit a zoning of base metals at the outer edges of the productive areas.

<u>Weathering</u>: Oxidation of the ubiquitous pyrite in the argillic zone produces outcrops of gaudy yellow, brown, orange, or red limonite stains. Relict textures of the original ore minerals and host rock minerals may be preserved.

Geochemical Expression: Rock samples in the argillic alteration zone have been used successfully to detect mineralized areas. Anomalous amounts of gold, silver, arsenic, antimony, lead, bismuth, mercury and possibly tellurium, selenium, molybdenum are indicative of possible mineralization. Stream sediment and soil samples are most likely to show anomalous amounts of silver, lead, and bismuth if the area is mineralized.

<u>Geophysical Expression</u>: Magnetic methods may be useful in determining the extent of argillic alteration under cover and down dip. Various electrical methods might be used, but interpretation of results may be difficult because of varying amounts of disseminated pyrite and clay alteration.

Associated Mineralization: Some deposits are found above and along the periphery of porphyry copper deposits giving some credence to the supposition that enargite-gold type deposits may be a shallow-level indicator of deeper porphyry systems.

#4.04B

Deposit Examples:

Masonic (CA) Mohave (CA) Stedman (CA)

Red Mountain (CO) Summitville (CO)

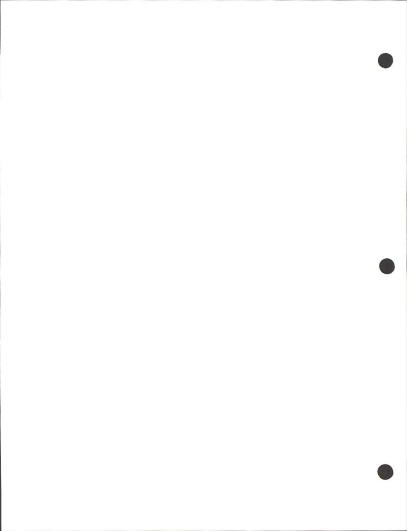
Borealis (NV) Goldfield (NV) Peavine-Wedekind (NV) Pyramid (NV)

References:

Eakle and McLaughlin, 1919 Troxel and Morton, 1962 Wright et al., 1953

Burbank, 1941 Stevens and Ratte, 1960

Mosier and Menzie, 1983 Ransome, 1909 Overton, 1947 Overton, 1947



#4.05

Deposit Type: HOT SPRINGS GOLD-SILVER

<u>Description</u>: Shallow, epithermal, disseminated and/or stockwork occurrences of hot-spring origin gold and silver found usually in felsic volcanic or volcaniclastic rocks.

GEOLOGIC FEATURES

Regional Setting: All known deposits of this type are found in regions of the Great Basin that have had significant volcanic activity resulting in felsic intrusive and extrusive rocks. Most occurrences are Tertiary and younger in age.

Structural Features: Areas of complex high-angle faulting found at caldera margins, resurgent domes, horst and grabens, or large, strike-slip faults with accompanying normal splays are the most favorable for these deposits. Hydrothermal brecciation and minor explosive volcanic activity such as breccia dikes are other features that are commonly exhibited in these deposits.

<u>Stratigraphic and Lithologic Characteristics</u>: The host rocks are nearly always volcaniclastics, lake beds, tuffaceous sandstones, or volcanic tuff breccias. Whatever the rock type, it must have a relatively high degree of permeability. This model type represents the surface or near-surface expression of a venting geothermal spring system.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: By their very nature, deposits of this kind are low grade but of large tonnage. Gold content commonly ranges from 0.05 to 0.20 ounces/ ton, and silver may range upward to 1 ounce/ton. Reserves can range up to 50 million tons. The bulk of the mineralization is disseminated into a favorable rock type or zone. Veins and stringers that generally accompany these deposits are usually enriched, particularly at the surface due to weathering of the more soluble components. Pyrite is common, often imparting a gray or bluich color to the quartz. Frequently, the gold and silver mineralization is accompanied by stibuite, arsenopyrite, fluorite, and sulfur. Various types of hydrothermal quartz are always present along with adularia, sericite, kaolinite, and alunite. Besides gold and silver, there may be enough mercury and alunite present to warrant recovery. <u>Alteration</u>: There is evidence of zoning at some deposits, but it is highly variable. A central area of massive silicification followed by an argillic zone and/or phyllic alteration is sometimes found. The alteration pattern seems to vary with the rock types present and the composition and intensity of the geothermal system.

<u>Ore Controls</u>: A high degree of permeability of the host rocks and a throughgoing fracture system is necessary to localize the ore solutions. Some fractures may be induced through hydrothermal brecciation and minor explosive volcanic activity.

<u>Weathering</u>: The high pyrite content of the host rock will generally oxidize into light browns, tans, reds, and white along with general bleaching of the rocks by geothermal action. Coloring is due to mixtures of limonite, hematite, goethite, jarosite, and alunite.

<u>Geochemical Expression</u>: Trace elements are present in the altered zones but are highly variable. Arsenic, mercury, antimony, and thallium are commonly found in anomalous amounts especially in the upper portions of the deposit and decreasing with depth.

<u>Geophysical Expression</u>: Little is known of any diagnostic signature above deposits of this type. It has been suggested that enhanced infrared photography might delimeate argillic areas surrounding these deposits.

Associated Mineralization: Epithermal precious metal quartz veins and base metal sulfides may lie at depth.



McLaughlin (CA)

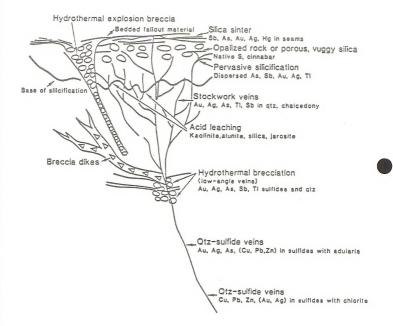
DeLamar (ID)

Hasbrouck Peak (NV) Round Mountain (NV) Sulphur (NV) References:

Averitt, 1945

Lindgren, 1900

Bonham and Garside, 1979 Tingley and Berger, 1980 Wallace, 1980



Idealized section through a hot-springs gold-silver type deposit (from USGS Open File Report 83-423)

Deposit Type: DISSEMINATED MERCURY

Synonym: Volcanic mercury

Description: Stratabound disseminated mercury minerals in volcaniclastic and associated sedimentary rocks.

GEOLOGIC FEATURES

Regional Setting: Nearly all deposits of this type have been found within and around the borders of volcanic centers located along major deep-seated fault zones.

Structural Features: Some sort of fracture system is required to give access to the ascending mineral solutions, but most any type of fault, fracture, or jointing seems to suffice for a channel way.

Stratigraphic and Lithologic Characteristics: These deposits are typically found in andesite lava flows and tuffs, andesite dikes, and volcanic vent breccia of Tertiary age. A few are found in shale, graywacke, and calcareous graywacke. The pore space in a rock unit is usually the critical factor in localizing the mineral solutions.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Size and shape of the various known deposits vary greatly. The mineralization is disseminated into favorable rock units cut by fractures and faults. Mercury content diminishes outward from the fracture feeder zones. Where permeable rock units are not available, mineralization is severely limited to small fractures and veinlets in discontinuous zones. Depth of these deposits is usually only 100 to 300 feet. Material considered "ore" varies with the current market price and the mining method employed but usually ranges from 3 to 12 pounds of mercury per ton. Cinnabar is usually the only mercury mineral present and is accompanied by abundant pyrite or marcesite and occasionally small quantifies of chalcedony, quartz, and calcite.

<u>Alteration</u>: The conspicuous lack of alteration of wall rock to clay is one of the characteristics of this type of deposit.

Ore Controls: Localization of cinnabar is controlled mainly by steeply dipping fracture and faults (that have had only small movement) cutting permeable rock units. Weathering: The chemical weathering of the pyrite or marcasite in the ore zone usually produces jarosite and gypsum in the outcrop area. Some clay minerals may also be present depending on the types of enclosing rocks, the degree of weathering, and the topography.

Geochemical Expression: Besides anomalous amounts of mercury in the general ore zone, arsenic and antimony are commonly found.

<u>Geophysical Expression</u>: Very little information is available on the geophysical signature of this type of deposit. The pyrite or marcasite in the ore should lend itself to detection by electro methods if the weathering is not too deep.

Associated Mineralization: Stibnite veins are known to occur occasionally in the general vicinity of these deposits.

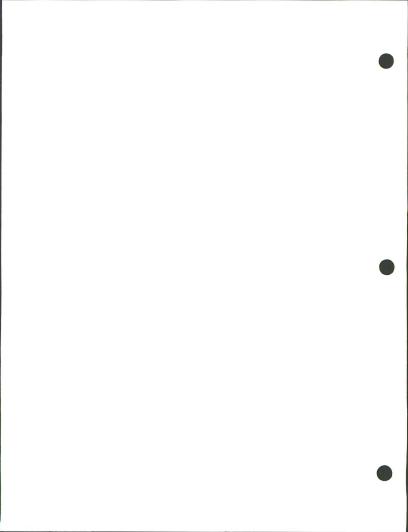
A&B (NV) Berry Creek (NV) Castle Rock (NV) Castle Rock (NV) Cinnabar Hill (NV) Diamonfield (NV) Hillside Mercury (NV) Nevada Cinnabar (NV) Poinsetta (NV) Red Bird (NV) Red Cloud (NV) San Pedro (NV) Washington (NV)

Black Butte (OR) Blue Boy (OR) Lucky Strike (OR) Red King (OR) Sesena Creek (OR)

References:

Bailey and Phoenix, 1944 Bailey and Phoenix, 1944

Wells and Waters, 1934 Ross, 1942 Ross, 1942 Ross, 1942 Ross, 1942 #4.06



Deposit Type: SILICA-CARBONATE MERCURY

Synonym: New Almaden type

Description: Cinnabar replacement of silica-carbonate rocks resulting from hydrothermal alteration of serpentine intruded into graywackes and siltstones.

GEOLOGIC FEATURES

<u>Regional Setting</u>: Nearly all the deposits are believed to have occurred in accreted bodies of rock in the upper plate of thrust faults in major subduction zones. All known deposits are of Tertiary age.

Structural Features: Usually found near sills and dikes of serpentine intruded into sedimentary rocks.

Stratigraphic and Lithologic Characteristics: The most important deposits have formed in shattered zones of Franciscan graywacke, siltstone, chert, and limestone above sloping serpentine intrusives.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Mineralization occurs as veins, veinlets, and impregnation in fractured rocks near and within fault zones. Minerals are cinnabar, metacinnabar, and minor amounts of pyrite, stibnite, chalcopyrite, sphalerite, galena, and bornite. Grade of ore commonly ranges from 0.20 to 0.75 percent mercury/ton. Size of the deposits are nearer the lower range.

<u>Alteration</u>: These deposits have formed as the result of hydrothermal action that has commonly replaced the serpentine at the contact with quartz and dolomite to form the slitac-actionate rock that is the host rock.

Ore Controls: Especially favorable situations are where the ore solutions have permeated along a serpentine-siltstone contact. At the contact with other rock types, the serpentine may be mineralized but to a lesser degree.

Weathering: No particular diagnostic features other than the formation of the white silica-carbonate rock. This may weather faster than the surrounding rocks and be covered on undisturbed outcrops.

<u>Geochemical Expression</u>: Little is known about the geochemical anomalies above deposits of this type. Probably high amounts of mercury and antimony would be found and possible copper and zinc.

Geophysical Expression: No published investigations are known.

Associated Mineralization: A few deposits have small stibuite veins nearby and within the mercury mineralization. In most cases they are of minor importance.

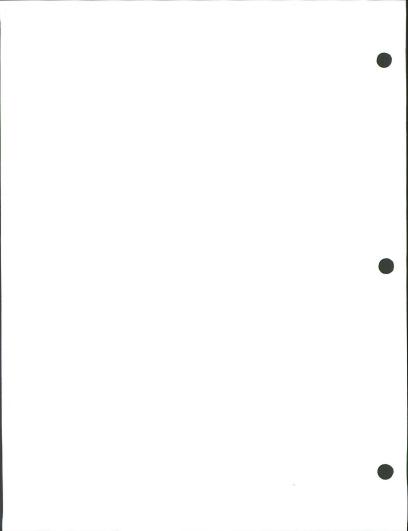


Abbott (CA) Aetna (CA) Bella Oak (CA) Chicago (CA) Contact (CA) Corona (CA) Culver-Bear (CA) Dewey's (CA) Esperanza (CA) Great Eastern-Mt. Jackson (CA) Harrison (CA) Helen (CA) Keystone (CA) Knoxville (CA) La Joya (CA) La Libertad (CA) Lion Den (CA) Manhatten (CA) Mirabel (CA) Mt. Diablo (CA) New Almaden (CA) Patriquin (CA) Polar Star (CA) Red Elephant (CA) Red Rick (CA) Reed (CA) Socrates (CA) Soda Springs (CA) Twin Peaks (CA) Wall Street (CA)

References:

Becker, 1888 Becker, 1888 Ransome and Kellog, 1939 Ransome and Kellog, 1939 Bradley, 1918 Ransome and Kellog, 1939 Ransome and Kellog, 1939 Bradley, 1918 Ransome and Kellog, 1939 Becker, 1888 Averitt, 1945 Ransome and Kellog, 1939 Becker, 1888 Averitt, 1945 Ransome and Kellog, 1939 Ransome and Kellog, 1939 Ransome and Kellog, 1939 Averitt, 1945 Ransome and Kellog, 1939 Becker, 1888 Bailey and Everhart, 1964 Ransome and Kellog, 1939 Becker, 1888 Averitt, 1945 Bradley, 1918 Averitt, 1945 Ransome and Kellog, 1939 Averitt, 1945 Ransome and Kellog, 1939 Becker, 1888





Deposit Type: HOT SPRINGS MERCURY

Synonym: Sulphur Bank type

<u>Description</u>: Dissemination of cinnabar and pyrite in predominantly volcanic sediments, intrusives, and extrusives associated with recent or paleo hot springs system.

GEOLOGIC FEATURES

Regional Setting: Generally centered around areas of Tertiary intermediate to mafic volcanic activity accompanied by normal faulting.

Structural Features: Favorable conditions are found in areas of fossil hot springs and at the water table level that existed during hot springs activity.

<u>Stratigraphic and Lithologic Characteristics:</u> Host rocks are most commonly extrusive flows ranging from andesite to basalt, andesite tuffs and tuff breccias, and diabase dikes. The critical feature is that the rock types must be either permeable or brittle enough to readily develop fractures for mineral solution access.

MINERAL CHARACTERISTICS

<u>Deposit Features:</u> Mineralization consists principally of cinnabar, native mercury, metacinnabar, and cordierite finely disseminated in microscopic crystals and aggregates in permeable rocks. It can occury intergranular spaces in the groundmass, coat fragments, line vugs and gas cavities, and is commonly intergrown with euhedral quartz. Size of the deposits is generally small, ranging from 250 to 490,000 tons. Grade varies over a narrow range of 0.20 to 0.60 percent/ton. Associated with the mercury mineralization is stibnite, pyrite, marcasite, dolomite, calcite, opal, cristobalite, and anatase.

<u>Alteration:</u> Hydrothermal alteration is pervasive with the formation of kaolinite, halloysite, sericite, jarosite, and montmorillonite. Alteration will vary with the rock types involved and the intensity of the solutions. In the upper parts of the hot springs system, native sulfur may be present.

<u>Ore Controls:</u> High angle faults and attending fractuces that intersect the paleo ground-water level within a hot spring system are the favorable conditions necessary for mineral deposition.

<u>Weathering</u>: Alteration will soften the enclosing rocks and along with weathering action will generally cause rapid erosion. Any quartz deposited in the form of quartz veinlets, sinter, and opalite as the result of the hot spring action will naturally be more resistant and may form a more or less resistant outcrop.

<u>Geochemical Expression:</u> Anomalous amounts of mercury, argenic, and antimony can be expected in the nearby vicinity, also possible anomalous amounts of gold. Some studies suggest present thermal waters very high in carbon dioxide, boron, ammonia, sodium, and iodine, and low in silica and potassium may be indicative of mercury deposits of this type.

<u>Geophysical Expression:</u> No characteristic geophysical signature has been documented around deposits of this type.

<u>Associated Mineralization:</u> Although antimony and arsenic are frequently associated with the mercury mineralization, they are not usually present in recoverable amounts. It has been suggested that gold may be present at depth or nearby since the chemical and physical conditions are similar for a deposit of this type.

Sulphur Bank (CA Walibu (CA)

Idaho Almaden (ID)

Baldwin (NV) B and B (NV) Butte (NV) Coleman (NV) Corders (NV) F and L (NV) Goldbanks (NV) Goldbanks (NV) Governor (NV) McDermitt (NV) Nevada Sulphur (NV) Steamboat Springs (NV)

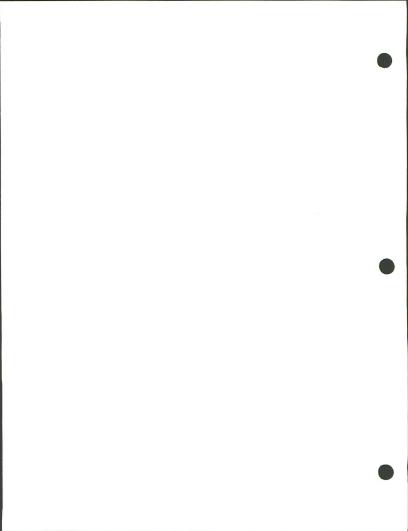
Giass Butte (OR) Opalite (OR) References:

White and Roberson, 1962 Averitt, 1945

Ross, 1956

Bailey and Phoenix, 1944 Bailey and Phoenix, 1944

Ross, 1942 Yates, 1942



Deposit Type: VOLCANOGENIC URANIUM

Description: Uranium mineralization in an epithermal assemblage of quartz, fluorite, and iron, arsenic, and molybdenum sulfides.

GEOLOGIC FEATURES

Regional Setting: Generally found in continental rift areas that have associated calderas. Age can range from Pre-Cambrian to Tertiary.

<u>Structural Features:</u> Usually associated with subaerial to subaqueous volcanic complexes of high silica alkali rhyolites and potash trachytes. Igneous rocks can be either shallow intrusives or vesicular flows.

Stratigraphic and Lithologic Characteristics: Peralkaline and peraluminous rhyolite are favored host rocks. Texture of the volcanics can range from porphyritic to microcrystalline.

MINERAL CHARACTERISTICS

<u>Deposit Features:</u> A typical deposit can have a grade of 0.02 to 0.3 percent uranium oxide and a size ranging from 20,000 tons to 6 million tons. Ore minerais are commonly coffinite, uraninite, and brannerite accompanied by pyrite, realgar and orpiment, jordisite, leucoxene, fluorite, quartz, adularia, and barite. Minor amounts of gold are found in some deposits, and it is not uncommon to find bastnaesite associated with alkaline complexes.

<u>Alteration:</u> Kaolinite, montmorillonite, and alunite are commonly formed in the ore zone. Also, silicification and the development of adularia are often found in the wall rock and mixed in with the ore minerals.

Ore Controls: A pervasive system of fractures and breccia along the edges and margins of intrusives direct the ore solutions. Where extrusive rocks were exposed to ore solutions, permeable units are important.

Weathering: Shallow weathering will produce a variety of uranium oxide minerals. Supergene secondary enrichment does not significantly develop.

<u>Geochemical Expression:</u> Anomalous amounts of arsenic, antimony, fluorine, and molybdenum are usually found in the ore zone. Rare earth elements may also be present (bastnaesite). Mercury and lithium are usually found at the edges of the mineralization and molybdenum near the bottom of deposition.

<u>Geophysical Expression:</u> No characteristic signature is known for deposits of this type.

Associated Mineralization: Roll-front type uranium deposits may develop in associated volcaniclastic sediments.



Anderson (AZ)

McDermitt (NV)

Aurora (OR) Lakeview (OR)

Marysville (UT) Spor Mountain (UT) Thomas Range (UT)

References:

Sherborne, et al., 1979

Rytuba and Glauzman, 1979

Roper and Wallace, 1981 Cohenour, 1960

Kerr, et al., 1957 Lindsey, 1977 Staatz and Carr, 1964

Deposit Type: SUBAERIAL VOLCANOGENIC MANGANESE

Synonym: Epithermal manganese

Description: Epithermal veins of manganese mineralization filling faults and fractures in subaerial volcanic rocks.

GEOLOGIC FEATURES

Regional Setting: All known deposits of this type are associated with Tertiary age volcanic centers.

Structural Features: Ring and radial faulting typically associated with volcanic centers are important in localizing ore solution. Also, the development of breccias along major faults are favorable loci.

Stratigraphic and Lithologic Characteristics: Most deposits occur in flows, tuffs, breccias, and agglomerates. Rock types can range in composition from rhyolites, dacites, and andesites to basalts.

MINERAL CHARACTERISTICS

Deposit Features: The ore minerals are usually psilomelane, pyrolusite, braunite, wad, manganite, and thodochrosite. Accessory minerals are commonly various iron oxides, manganocalcite, calcite, quartz chalcedony, barite, and zeolites. These deposits are found in veins that may form into bunches and pockets or as stringers, nodular masses, or as disseminations. Individual occurrences are generally small, ranging from 2,600 to 290,000 tons. The grade usually ranges from 20 to 42 percent manganese.

Alteration: The main and most significant alteration product is kaolinite.

<u>Ore Controls</u>: It is important to have a throughgoing system of faults and fractures and/or brecciated volcanic rocks.

Weathering: The formation of abundant manganese and iron oxide minerals mixed in with kaolinite are the most apparent products at the surface of these deposits.

<u>Geochemical Expression</u>: Anomalous amounts of manganese are distributed over a wide area. Patches of anomalous amounts of lead, silver, gold, and copper may occur but not consistently at every deposit.

Geophysical Expression: No characteristic signature is known.

Associated Mineralization: Epithermal gold and silver deposits are known to occur in the general vicinity of some deposits of this type.

Armour Group (AZ) California Group (AZ) Hatton (AZ) J. M. Meadows Group (AZ) Manganese Development (AZ) Thurston and Hardy (AZ) Topock (AZ) U.S. Group (AZ)

Black Crow-San Juan (NM) Cliff Roy (NM) Goryana (NM) Griffith (NM) JVB Claim (NM) M and M Group (NM) Manganese Chief (NM) Niggerhead (NM) Phillips Lease (NM) Red Hill (NM)

References:

Jones and Ransome, 1919 Jones and Ransome, 1919 Farnham and Stewart, 1958 Jones and Ransome, 1919 Jones and Ransome, 1919 Jones and Ransome, 1919 Farnham and Stewart, 1958

Farnham, 1961 Farnham, 1961 Farnham, 1961 Mosier, 1983a Mosier, 1983a Farnham, 1961 Farnham, 1961 Mosier, 1983a Farnham, 1963

Deposit Type: CARBONATE-HOSTED MANGANESE REPLACEMENT

Description: Epigenetic veins and cavity fillings of manganese mineralization in limestone, dolomite, or marble.

GEOLOGIC FEATURES

Regional Setting: Found along continental margins that have been subjected to folding and major deformation and accompanied igneous intrusions. Although the deposits can be any age, most are Palezozic or Mesozoic.

Structural Features: Favorable regions are miogeosyncline sequences of sediments that were later intruded by small plutons or other intrusive complexes.

Stratigraphic and Lithologic Characteristics: The predominant sedimentary rocks are the calcareous types such as limestones, dolomites, and marble. The intrusives are commonly granite and granodiorite.

MINERAL CHARACTERISTICS

Deposit Features: The most common ore minerals are psilomelane, pyrolusite, rhodochrosite, wad, manganite, and rhodonite. Accessory minerals are usually calcite, quartz, barite, fluorite, jasper, and manganocalcite. Minor amounts of sulfides, such as pyrite, chalcopyrite, galena, and sphalerite, may occur in the ore zone. This mineral assemblage is found in the form of tabular veins, irregular open space fillings, lenticular pods, chimneys, or pipes. Grade of the deposits can range from 15 to 45 percent manganese. Size can vary from 1,000 to 600,000 tons, but the average deposit is 25,000 tons or less.

Alteration: Because this type of deposit frequently develops at a limeigneous contact, skarn mineralization is commonly found in the limy members.

<u>Ore Controls</u>: Open space filling in porous carbonite rocks is important in emplacement of the ore minerals. Also, the igneous intrusive contacts are generally loci of mineralization.

Weathering: Limonite and kaolinite generally develop at the outcrop along with blackening of the surrounding rocks by the relatively soluble manganese oxide minerals. <u>Geochemical Expression</u>: Anomalous amounts of manganese are distributed over a wide area. High amounts of copper, lead, and zinc are also possible, but these elements are not always present in the ore.

Geophysical Expression: No documented characteristic signature is known other Than that associated with igneous intrusives that may or may not be mineralized.

Associated Mineralization: Silver-bearing replacement veins are known to occur in the mear vicinity of some deposits. Copper, lead, and zinc are sometimes found where skarn develops but not often in commercial amounts.

Atlas (AZ) Crown King (AZ) Danville-Hanchette (AZ) Golden Gate (AZ) Hendricks-Twilight (AZ) Mamoth (AZ) No. 4 - Summit (AZ) Oregon (AZ) Waterloo (AZ)

Philipsburg (MT)

Essex and Steptoe (NV)

Bear Mtn. (NM) Birchfield (NM) Chloride Flat (NM) Kingston (NM) Lake Valley (NM) Lone Mtn. (NM)

Detroit (UT)

References:

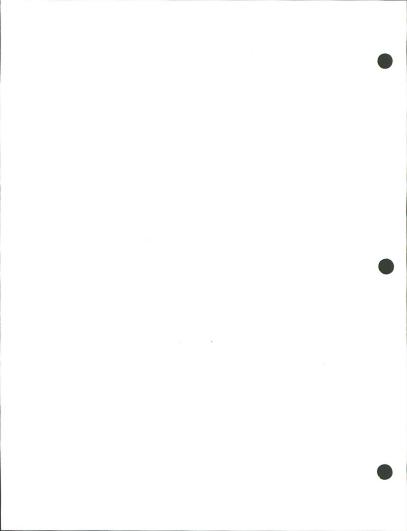
Jones and Ransome, 1919 Jones and Ransome, 1919 Farnham et al., 1961 Farnham et al., 1961

Goddard, 1940

Crittenden, Jr., 1964

Farnham, 1961 Farnham, 1961 Farnham, 1961 Mosier, 1983b Farnham, 1961 Mosier, 1983b

Mosier, 1983b



Deposit Type: RED-BED COPPER

Synonym: Copper shales; sediment-hosted copper

Description: Stratabound, disseminated copper sulfides in a sedimentary sequence that have accumulated under reducing conditions.

GEOLOGIC FEATURES

Regional Setting: Deposits are restricted to intracratonic rift basins and the shallow-marine environment that developed over the filled rifts.

Structural Features: Block faulting and gentle folds are the main tectonic features present. Some reverse faults of small magnitude are present in a few deposits.

Stratigraphic and Lithologic Characteristics: The rifts are commonly filled with thick sequences of sediments, evaporites, and sometimes hypelyses1 and volcanic rocks of mostly basaltic nature. The alluvium is generally oxidized red but may be gray, green, or white in color. Beds of sandstone, conglomerate, limestone, and organic-rich shales are frequently intercalated. Evaporites, when present, commonly cap this sequence. Nearly all deposits are either late Precambrian or Permian to early Mesozoic in age.

MINERAL CHARACTERISTICS

Deposit Features: Copper mineralization in any one area is restricted to one or more sedimentary zones probably dictated on original permeability and/or porosity. Sandstones and shales are the most preferred host rocks, especially units containing abundant fossil wood, algal material, and biogenic sulfur. The usual ore minerals are chalcocite and lesser amounts chalcopyrite and bornite. Pyrite is nearly always present and some deposits exhibit small amounts of native silver. Dissemination of the minerals within the host rock is frequently zoned with centers of chalcocite and bornite, a rim of chalcopyrite, and minor amounts of galena and sphalerite in the peripheral areas. Native copper may be the primary ore mineral in some deposits rather than chalcocite. Grade of the deposits can range from 0.5 to 4.0 percent copper but most are one percent or less. Values are generally concentrated in thin layers several inches thick but may range on occasion to as much as 25 feet thick. A few deposits are as large as 1 million tons, but most are much smaller. Size of the deposits ranges widely from 1.65 million to 360 million tons.

<u>Alteration</u>: Varying degrees of alteration may reduce the red-beds to gray, green, or white in color. Beds subjected to regional metamorphism may be changed to purple color. <u>Ore Controls</u>: An oxidation-reduction interface in a low pH environment is necessary to mobilize the copper and to allow it to be redeposited and concentrated. Pyritic sediments and carbonaceous matter apparently enhances concentration of the copper.

<u>Weathering</u>: The effect may vary considerably depending on the paleo-climate. Copper staining and red spots from oxidized pyrite in dark limestones and shales are the most obvious features. Outcrop may be completely leached of copper and redeposited down dip.

<u>Geochemical Expression</u>: Anomalous concentrations of copper, lead, zinc, vanadium, and molybdenium are the elements most likely sought as pathfinders. Anomalous amounts of silver and uranium may also be present in potential mineralized area.

Geophysical Expression: None are presently known.

Associated Mineralization: Evaporites commonly overlay the red-beds and may contain extractable amounts of halite, sylvite, and gypsum.



Milan (KS) Runnymede (KS)

White Pine (MI)

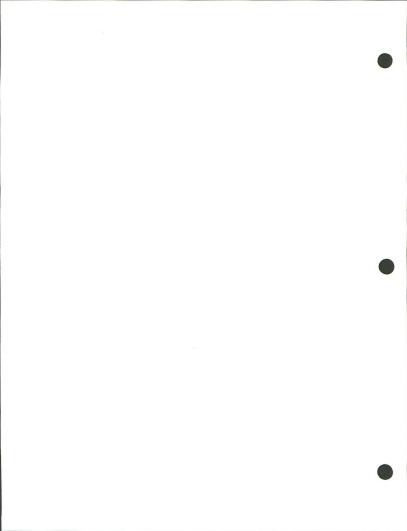
Eureka (NM) High Rolls (NM) Nacimiento (NM) Pintada (NM) Rayo (NM) Stauber (NM) Zuni (NM)

Byars (OK) Creta (OK) Magnum (OK) Teepee (OK)

Wyoming Co. (PA)

Buzzard Peak (TX) Copper Breaks (TX) Crowell (TX) Medicine Mound (TX) Old Giory (TX)

References: Waugh and Brady, 1976 Waugh and Brady, 1976 Ensign et al., 1968 LaPoint, 1976 Stroud et al., 1970 Dingess, 1976 Johnson, 1976 Stroud et al., 1970 Bulter, 1938 Smith, 1976 Smith, 1976 Smith, 1976 Smith, 1976 Smith, 1976



Deposit Type: SANDSTONE URANIUM

Synonym: Epigenetic roll-front; epigenetic carbonaceous uranium

Description: Deposits, chiefly in sandstones, of uranium oxides minerals that mainly fill the pores of the host rock and replace fossil plants but also partly replace the sand grains and the cementing minerals of that rock.

GEOLOGIC FEATURES

Regional Setting: Most deposits are restricted to stable platforms or forelands, coastal plains of the continental shelf, and intermontane basins of a mobile belt within pre-foreland areas. Geosynclinal and oceanic crustal settings are unfavorable.

Structural Features: Wide continental basins of 25 to 150 miles are most favorable for regional depositions. On a smaller scale, fluvial, lacustrine, deltaic, and strand-plain features are favored over other parts of a basin. Many deposits are localized in the mid-fan facies of alluvial fans.

Stratigraphic and Lithologic Characteristics: The most favorable host rock is a medium to coarse grained sandstone composed of devitrified volcaniclastics or feldspathic material containing intercalated carbonized plant matter and/or pyrite. When units of this type are interbedded with mudstones, a situation of permeable beds restricted by impermeable beds can cause increased concentration of uranium minerals. Age of favorable host rock may range from Silurian to Tertiary.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: Deposits may occur as tabular bodies that lie nearly parallel to the bedding of the host rock or as roll-front structures occupying part or all of a permeable rock unit. Ore minerals are most often carnotite, coffinite, or pitchbledde but can contain varying amounts of 40 or so other uranium minerals. Pyrite is almost invariably present in and around the mineralized area. Also, small amounts of selenium, molybdenum, and vanadium are commonly present in the mineralized area. Ore bodies may range from a few hundred tons to several million tons. Overall grade of ore bodies generally averages 0.2 percent Vq0.

<u>Alteration</u>: Because of the reduction-oxidation environment caused by the pyrite and carbonized plant matter, favorable sandstones are frequently altered to light gray or white but may be shades of pink, yellow, or brown depending on minerals present. Mudstones are often altered to green or gray. In roll-front deposits, the oxidized sandstone is frequently pinkish and the reduced sandstone is gray. Ore Controls: Permeable rock units with low dips containing carbonaceous material and a sulfur source are the primary host requirements. There must be a uranium source either from granite-core mountain range or tuffaceous formation. Leaching of these sources into the favorable host unit is required for deposition and concentration of the uranium.

<u>Weathering</u>: May produce colorful oxidized uranium minerals of yellow, green, motiled brown, bluish and greenish black. Sandstones may have bleached appearances. Any enclosed pyrite will be oxidized forming red to brown splotches in the sandstone.

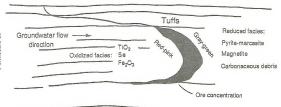
Geochemical Expression: Anomalous amounts of selenium, molybdenum, vanadium, copper, and uranium are the usual indicator elements exhibited.

Geophysical Expression: Anomalous radioactivity of the order of 5 to 10 times background count as measured by gamma-ray spectrometers is generally considered significant. Low order seismic methods might be applicable to determine thickness of favor beds and pinchouts or to detect regional unconformities that could signal favorable depositional conditions.

Associated Mineralization: Because of similar environment, red-bed copper deposits may occur in the same general area. Occasionally, the vanadium content of some uranium deposits is great enough to warrant recovery as a by-product.

Nearly 4600 deposits in sandstone are known in the United States. Most of the major deposits occur within the Colorado Plateau and the intermontane basins of Wyoming. References:

Finch, 1976 Finch, 1982 Nash et al., 1981



Idealized section across a roll-front sandstone uranium deposit (from USGS Open File Report 83-423)



Deposit Type: BEDDED BARITE

Synonym: Stratiform barite

Description: Fine-grained barite disseminated in stratiform bodies of layered sequences of siliceous and/or carbonate rocks.

GEOLOGIC FEATURES

Regional Setting: The rocks containing bedded barite were deposited in a wide variaty of environments, but most show evidence of being deposited in eugeosynclinal areas within orogenic beits accreted to continental margins.

Structural Features: Some deposits are associated with hinge faults that are found in sedimentary basins.

Stratigraphic and Lithologic Characteristics: Almost all deposits are in well-bedded siliceous and/or limestone rocks. The deposits are frequently found intercalated with chert, shale, mudstone, argillite, and sandstone. The majority of the deposits are in Devonian rocks, but are known to occur in Cambrian, Ordovician, and lower Mississippian rocks as well.

MINERAL CHARACTERISTICS

Deposit Features: The barite may occur as very fine-grained but distinct disseminated mineral grains or as cementing material between rock grains. The overall appearance of bedded concentrations is dark gray to black layers or beds of spotty or shaley limestone. The rock frequently gives off an odor of hydrogen sulfide when freshly broken and commonly sparkles to a slight extent in sunlight. The beds are often lense-shaped or tabular and may be from a few inches to 50 feet thick. Laterally, the beds can extend over many acres but may also pinch, swell, and become discontinuous because of folding and faulting. The barite is generally found as small grains as small as 0.1 millimeter but occasionally occurs in larger masses aro softees and nodules in the enclosing rock. Where concentrated, the beds may contain 50 to over 90 percent barite. The deposits can range from 130,000 tons to 30 million tons in size.

<u>Alteration</u>: Heat generated during burial or from external sources will cause <u>originally</u> dark colored beds to become light gray or white. This feature depends on the amount of organic carbon that is present.

<u>Ore Controls</u>: Depositional basins formed morphologic traps for location of barium solutions in select sedimentary units. Where barite forms a cement in sandstone and siltstone beds, it indicates migration and precipitation of barium solutions during diagenesis. Weathering: Usually no distinct features. Generally resembles ordinary limestone or dolomite. Occasionally barite rosettes or nodules will weather out along the surface exposure.

<u>Geochemical Expression</u>: A high barium content in the near vicinity of a deposit along with an extremely low calcium and magnesium content is common. Associated strontium values may show as high as 7000 ppm. Where barite is found peripheral to sedimentary-hosted zinc-lead, there may be lateral zoning of copper, lead, zinc, barite, and/or regional manganese halos.

Geophysical Expression: No distinct nor diagnostic pattern is known to occur over bedded barite deposits.

Associated Mineralization: Occasionally, marine exhalant zinc-lead type deposits are found in the near vicinity but are not thought to be genetically related to the barite.

Deposit Examples:

Magnet Cove (AR)

Castella (CA)

Sun Valley (ID)

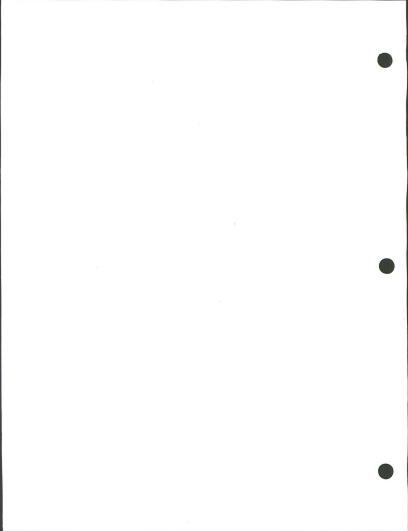
Argenta (NV) Bateman Canyon (NV) Greystone (NV) Mountain Springs (NV) Northumberland (NV) Slaven Canyon (NV)

Uribe (WA)

References:

Scull, 1958
Weber and Mathews, 1967
Brobst, 1958
Ketner, 1963
Ketner, 1963
Ketner, 1963
Shawe et al., 1969
Ketner, 1963

Orris, 1983



Deposit Type: PLACER GOLD

<u>Description</u>: Grains of elemental gold and sometimes platinum-group alloys deposited and concentrated by flowing water in gravel, sand, silt, clay, and their consolidated equivalent.

GEOLOGIC FEATURES

Regional Setting: Found in areas that were tectonically active during the Cenozoic, particularly the western coast ranges of the United States.

Structural Features: Near major fault zone areas where erosion has proceeded long enough to have repeatedly reworked the sediments.

Stratigraphic and Lithologic Characteristics: These deposits typically form where a high energy stream gradient flattens and water velocity lessens as at the inside of stream meanders, or below rapids and falls. They will also form beneath boulders, in vegetation mats, and along past strandlines lines.

MINERAL CHARACTERISTICS

Deposit Features: The gold is most often found as small flattened flaky particles that can range downward in size to fine powder. More rare are equidimensional nugget size pieces. Accessory minerals are varying amounts of magnetite, chromite, ilmenite, zircon, garnet, and rutile. None of these heavy minerals are usually present in economic quantities. Size and grade of placers vary greatly depending on the amount of original gold in the source rock, the degree of erosion to free the gold, and the amount of mechanical concentration of the freed particles.

Alteration: Not found in this type of deposit.

Ore Controls: Highly sensitive to the depositional environment. Natural impediments of water flow carrying gold particles can create traps and enhance concentration. Such features are natural riffles, transverse fractures, dikes, and joints. Particularly favorable are upturned beds of slate, schist, and phyllite that lie oblique to the stream flow.

Weathering: Because of the high specific gravity, weathering processes tend to move the gold particles through the alluvium to bedrock. Local in-place concentration may take place at the outcrop if the surface is not too steep. <u>Geochemical Expression</u>: Anomalous amounts of silver, arsenic, mercury, antimony, copper, sulfur, and iron are sometimes found with these deposits. Not all elements are always present. Also, the heavy-mineral group will be present in abnormal amounts.

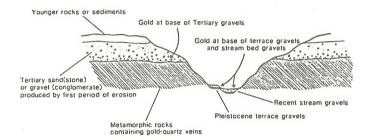
<u>Geophysical Expression</u>: No characteristic signature other than that associated with alluvium. Placers may be cemented and give an indication of consolidated sediments.

Associated Mineralization: In areas of ultramafic complexes and serpentine, platinum group minerals may be associated with the gold in the placer deposit. Epithermal vein deposits that contributed the placer gold may be nearby.

Deposit Examples:*	References:**
АК	Brooks, 1913
CA	Lindgren, 1911
CO	Henderson, 1926
ID	Savage, 1961
MT	Koschmann and Bergendahl, 1968
OR	Brooks and Ramp, 1968

* Individual deposits are too numerous to list. **These are general references that describe the many deposits that occur in each of the main placer-occurring states.

#6.01



Idealized cross-section through a Tertiary Sierra Nevada type gold placer (from USGS Open File Report 83-901)

Deposit Type: MARINE PHOSPHATE

Subtype: Upwelling

<u>Description</u>: A major stratigraphic unit of phosphorite sediments within a sequence of marine beds that were formed in basins having easy access to the open sea and upwelling currents.

GEOLOGIC FEATURES

Regional Setting: Deposits are limited to former shelf and platform areas and/or areas of miogeosynclines and eugeosynclinal deformation. Age may range widely from Precambrian through Miocene.

<u>Structural Features</u>: Marine sedimentary basins with shoaling bottoms were the favored environment. The thickest accumulations formed in areas of geosynclinal subsidence.

Stratigraphic and Lithologic Characteristics: Where favorable marine basins have formed, a typical lateral sequence of rocks are found in a seaward direction. These consist of dark carbonaceous shale followed by phosphatic shale, phosphorite, dolomite, chert or diatomite, several facies of carbonate rock, saline beds, and red or light-colored sandstone or shale. These rocks grade laterally into each other and are also found nearly in the same order vertically or in reverse sequence.

MINERAL CHARACTERISTICS

<u>Deposit Features</u>: The principal primary mineral is apatite, but secondary processes include diagenetic phosphatization of calcium carbonate and interstitial precipitation, reworking by waves and currents, and weathering. More often, this changes the material into a series of undistinguishable secondary minerals and collectively called phosphorite that occurs as pellets, nodules, phosphatized shell, and bone materials. Accessory minerals are dolomite, calcite, quartz, clays, and sometimes variable amounts of iron oxide, gypsum, halite, and pyrite. Occasionally, small range up to 150 feet thick, but the mineable beds are usually only a few feet thick. Grade generally ranges from 15 to 33 percent phosphorus oxide, but the riccher portions of a deposit are commonly local in extent and highly lenticular. Size of the deposits can be very large (a range from 25 million toons to as much as 3,500 million tons.

Alteration: No alteration accompanies these deposits.

<u>Ore Controls</u>: Stratigraphic basins or parts of basins that were favorable for the accumulation of organic-rich sediments and subsequent conditions that allowed the evolutionary processes to form phosphorites are the controlling features.

Weathering: Phosphorite breaks down easily during weathering and rarely forms natural outcrops. The presence of concealed phosphorite beds can often be detected by looking for tan to black chips of phosphorite in the soil. Highly weathered pieces have a characteristic dull-bluish-white film or bloom.

Geochemical Expression: Anomalous amounts of phosphate, nitrogen, fluorine, carbon, and uranium are usually present.

<u>Geophysical Expression</u>: Because nearly all marine phosphorites contain from 0.005 to 0.02 percent uranium, gamma-ray well logging can be used for identification. Under favorable circumstances, aerial gamma-ray spectrometry can be used.

Associated Mineralization: The depositional environment is conducive to possible sedimentary manganese to form in the same basins but not necessarily along with the phosphorite.

Deposit Examples:

Brooks Range (AK)

New Cuyama (CA)

Conda (ID) Henry (ID) Mabie Canyon (ID) Wooley Valley (ID)

Warm Springs (MT)

Lee Creek (NC)

Uinta Mtns. (UT) Vernal (UT)

Wyoming

References:

Patton and Matzko, 1959

Gower and Madsen, 1964

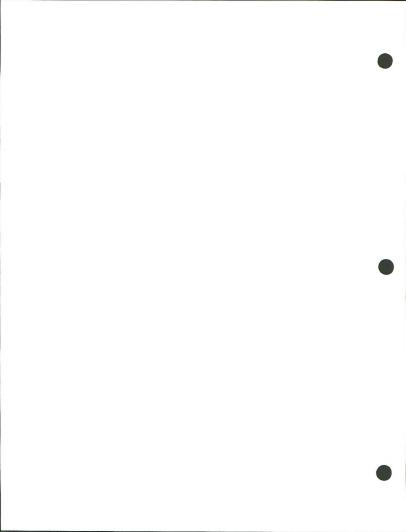
Gulbrandsen and Krier, 1980 Gulbrandsen and Krier, 1980 Gulbrandsen and Krier, 1980 Emigh, 1958

Popoff and Service, 1965

Mosier, 1983c

Cheney, 1957 Cheney, 1957

Sheldon, 1963



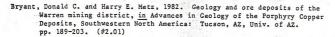
REFERENCES

(Numbers at the end of each reference denote specific deposit types)

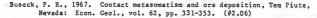
- Abbott, J. T. and S. A. Williams, 1981. The Pine Grove Molybdenum system, southern Wah Wah Mountains, Beaver County, Utah: Am. Inst. Mining, Metall., and Petroleum Engineers 101st Annual Meeting, p. 21. (#2.02A)
- Albers, John P. and Jacques F. Robertson, 1961. Geology and ore deposits of East Shasta copper-zinc district, Shasta County, California: U.S. Geol. Sur. Prof. Paper 338, 107pn. (#3.018)
- Anderson, C. A., E. A. Scholz, and J. D. Strobell, Jr., 1956. Geology and ore deposits of the Bagdad area, Yavapai County, Arizona: U.S. Geol. Sur. Prof. Paper 278, 103pp. (#2.01)
- Anderson, C. A. and S. C. Creasey, 1958. Geology and Ore deposits of the Jerome area, Yavapai County, Arizona: U.S. Geol. Sur. Prof. Paper 308, 185pp. (#3.01C)
- Anderson, Philip and John M. Guilbert, 1979. The Precambrian massive sulfide deposits of Arizona - a distinct metallogenic epoch and province, <u>in</u>: Papers on Mineral Deposits of Western North America: NV Bur. Mines Geol. Report 33, pp. 34-48. (#3.01B, 3.01C)
- Atkinson, W. W., Jr. and M. T. Einaudi, 1978. Skarn formations and mineralization in the contact aureole at Carr Fork, Bingham, Utah: Econ. Geol., vol. 73, pp. 1326-1355. (#2.04)
- Atkinson, W. W. Jr., J. H. Kaczmarowski, and J. Erickson, Jr., 1982. Geology of the skarn-breccia orebody at the Victoria Mine, Elko County, Nevada: Econ. Geol., vol. 77, no. 4, pp. 899-918. (#2.04)
- Averitt, Paul, 1945. Quicksilver deposits of the Knoxville district, Napa, Yolo, and Lake Counties, California: CA. Jour. Mines and Geol., vol. 41, no. 2, pp. 65-89. (\$4.05, \$4.07, \$4.08)
- Bailey, Edgar and David A. Phoenix, 1944. Quicksilver deposits in Nevada: Univ. NV Bull. vol. 38, no. 5, Geol. and Mining Series no. 41, 197pp. (#4.06, #4.08)
- Bailey, Edgar and Donald S. Everhart, 1964. Geology and quicksilver deposits of the New Almaden district, California: U.S. Geol. Sur. Prof. Paper 360, pp. 1-206. (#4.07)
- Barnes, M. P. and J. S. Simos, 1968. Ore deposits of the Park City district, Utah, in Ore Deposits of the United States, 1933-1967, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 1102-1126. (#4.01)



- Barter, Charles F. and James L. Kelly, 1982. Geology of the Twin Buttes mineral deposit, Pima County, Arizona, in Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 407-432 (#2.01, #2.04)
- Bateman, Alan M., 1924. Geology of the Beatson Copper Mine: Econ. Geol., vol. 19, pp. 338-368 (#3.01B)
- Bateman, Alan M., 1942. The ore deposits of Kennecott, Alaska, in Ore Deposits as Related to Structural Features: Princeton, NJ, Frinceton University Press, pp. 188-193. (#4.01)
- Bauer, H. L., Jr., R. A. Breitrick, J. J. Cooper, and J. A. Anderson, 1966. Porphycy copper deposits in the Robinson mining district, Nevada, <u>in</u> Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 233-244. (#2.01)
- Becker, G. F., 1882. Geology of the Comstock Lode and the Washoe district: U.S. Geol. Sur. Monograph 3, with Atlas. 442pp. (#4.04A)
- Becker, George F., 1888. Geology of the quicksilver deposits of the Pacific Slope: U.S. Geol. Sur. Monograph vol. 13, 485pp. (#4.07)
- Bennett, E. H., 1937. Petrology and trace element distribution of the White Cloud stock: unpub. Ph.D. thesis, Univ. ID, Moscow, ID. 172pp. (#2.4)
- Blake, D. W., Ted G. Theodore, and Edward L. Kretschmer, 1978. Alteration and distribution of sulfide mineralization at Copper Canyon, Lander County, Nevada: AZ Geol. Soc. Digest, vol. 11, pp. 67-78. (#2.04)
- Blake, S. W., 1971. Geology, alteration, and mineralization of the San Juan Mine area, Graham County, Arizona: unpub. M.S. thesis, Univ. of AZ, Tucson, AZ. 85pp. (#2.01)
- Bonham, H. F. and L. J. Garside, 1979. Geology of the Tonopah, Lone Mtn., Klondike, and Northern Mud Lake quadrangles, Nevada: NV Bur. Mines and Geol. Bull. 92, 142pp. (#4.04A, #4.05)
- Bradley, Walter W., 1918. Quicksilver resources of California: CA Mng. Bur. Bull. 78, 389pp. (#4.07)
- Brobst, Donald A., 1958. Barite resources of the United States: U.S. Geol. Sur. Bull. 1072-B, pp. 92-97. (#5.03)
- Brooks, Alfred H., 1913. The mineral deposits of Alaska: U.S. Geol. Sur. Bull. 592. (#6.01)
- Brooks, H. C. and Len Ramp, 1968. Gold and silver in Oregon: OR Dept. Geol. and Mineral Ind. Bull. 61, 338pp. (#6.01)



- Buddington, A. F. and B. F. Leonard, 1962. Regional geology of the St. Lawrence County Magnetite district, northwest Adirondacks, New York: U.S. Geol. Sur. Prof. Paper 376, 145pp. (#2,03)
- Burbank, W. S., 1932. Geology and ore deposits of the Bonanza mining district, Colorado: U.S. Geol. Sur. Prof. Paper 169, 166pp. (#4.04A)
- Burbank, W. S., 1941. The structural control of ore deposits in the Red Mountain, Sneffels, and Telluride districts of the San Juan Mountains, Colorado: CO Sci. Soc. Proc., vol. 14, no. 5, pp. 141-261. (#4.04A, #4.04B)
- Burbank, W. S. and Robert G. Luedke, 1964. Geology of the Ironton quadrangle, Colorado: U.S. Geol. Sur. Quad Map GQ-291. (#2.02A)
- Burbank, W. S. and Robert G. Luedke, 1968. Geology and ore deposits of the western San Juan Mountains, Colorado, in Ore Deposits of the United States, 1933-1967, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 714-733. (\$4.01)



Butler, B. S., G. F. Loughlin, V. C. Heikes, and others, 1920. The ore deposits of Utah: U.S. Geol. Sur. Prof. Paper 111, pp. 391-395. (#4.02, 4.04A)

Butler, Robert D., 1938. A "Red Bed" type copper occurrence, Wyoming County, Pennsylvania: Econ. Geol., vol. 33, pp. 625-634. (#5.01)

Callaghan, Eugene, 1938. Preliminary report on the alunite deposits of the Marysvale region, Utah: U.S. Geol. Sur. Bull. 886-D, pp. 98-100 (\$4.04A)

Callaghan, Eugene, 1939. Geology of Searchlight district, Clark County, Nevada: U.S. Geol. Sur. Bull. 906-D, pp. 135-188. (#4.04A)

Callaghan, Eugene and A. F. Buddington, 1938. Metalliferous mineral deposits of the Cascade Range in Oregon: U.S. Geol. Sur. Bull. 893, 141pp. (\$4.04A)

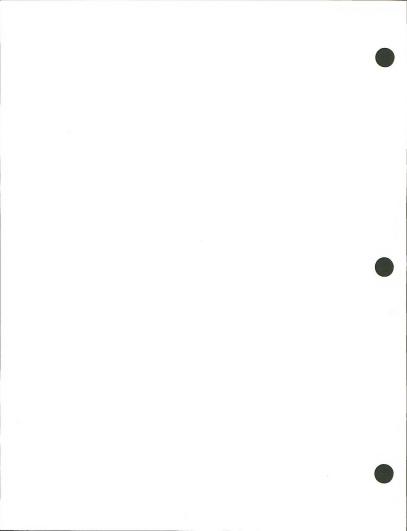
Chaffee, Maurice, 1977. Geochemical exploration techniques based on distribution of selected elements in rocks, soils, and plants, Vekal porphyry copper deposit, Pimal County, Arizona: U.S. Geol. Sur. Bull. 1278-E, 78pp. (42.01)

Castiel a serial and it is the

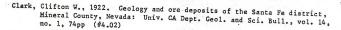


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Cheney, Thomas W., 1957. Phosphate in Utah and an analysis of the stratigraphy of the Park City and the Phosphoria Formations: UT Gev1. and Mineralogy Sur. Bull. 59, 54pp. (#6.02)

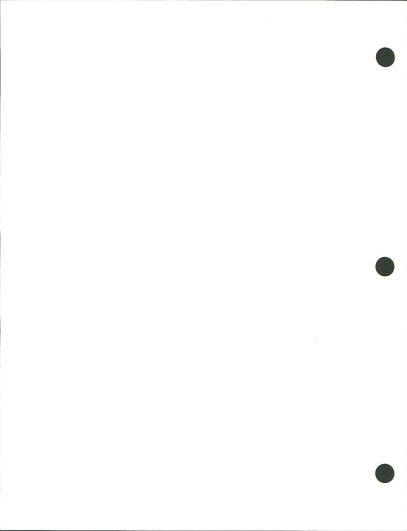


- Cohenour, R. E., 1960. Geology and uranium occurrences near Lakeview, Oregon: U.S. Atomic Energy Comm., RME 2070, 33pp. (#4.09)
- Collins, B. I., 1977. Formation of scheelite-bearing and scheelite-barren skarns of Lost Creek, Pioneer Mountain, Montana: Econ. Geol., vol. 72, pp. 1505-1523. (#2.06)
- Cooper, J. R., 1951. Geology of the tungsten, antimony, and gold deposits near Stibnite, Idaho: U.S. Geol. Sur. Bull. 969-F, pp. 151-197. (#2.06)
- Corn, Russell M., 1975. Alteration-mineralization zoning, Red Mountain, Arizona: Econ. Geol. vol. 70, pp. 1437-1447. (#2.01)
- Cox, Dennis P. and Singer, Donald A., Editors, 1986 Mineral Deposit Models, U.S. Geological Survey Bulletin 1693
- Creasey, S. C. and G. L. Quick, 1955. Copper deposits of part of Helvetia mining district, Pima County, Arizona: U.S. Geol. Sur. Bull. 1027-F, pp. 301-323. (#2.01)
- Crerar, David A., Jay Namson, Michael So Chyl, Loretta William, and Mark D. Feignson, 1982. Manganiferous cherts of the Franciscan Assemblage: I. General Geology, Ancient and Modern Analogues, and Implications for Hydrothermal Convection at Oceanic Spreading Centers: Econ. Geol. vol. 77, pp. 519-540. (#3.02)
- Crittenden, M. D., Jr., 1964. Manganese, in Mineral and Water Resources of Nevada: Mineral and Water Resources of NV, Senate Doc. no. 87 (88th Congress) pp. 113-119. (#4.11)
- Cummings, Robert B., 1982. Geology of the Sacaton porphyry copper deposit, Pinal County, Ariz., in Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 507-521. (\$2.01)
- Dingess, Faul D., 1976. Geology and mining operations at the Creta copper deposit of Eagle-Pitcher Industries, Inc., in Stratiform Copper Deposits of the Midcontinent Region, A Symposium: OK Geol. Sur. Circ. 77, pp. 15-24. (#5.01)
- Dings, M. G., 1952. The Wallapai mining district, Cerbat Mountain, Mohave County, Ariz.: U. S. Geol. Sur. Bull. 978-A, pp. 123-163. (#2.01)

Dings, M. G. and C. S. Robinson, 1957. Geology and ore deposits of the Garfield quadrangle, Colorado: U.S. Geol. Sur. Prof. paper 289, 110pp. (#2.02A)







- Dixon, D. W. 1966. Geology of the New Cornelia Mine, Ajo, Arizona, in Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp 123-132. (#2.01)
- Dubois, R. L. and R. W. Brummett, 1968. Geology of the Eagle Mountain mine area, in Ore Deposits of the United States, 1933-1967 (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 1592-1606. (#2.03)
- Dunn, Peter G., 1982. Geology of the Copper Flat porphyry copper deposit, Hillsboro, Sierra County, New Mexico, <u>in</u> Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 313-325. (#2.01)
- Durrell, C. and P. D. Proctor, 1948. Iron ore deposits near Lake Hawley and Spencer Lakes, Sterra County, California: CA Div. Mines Bull. 129, pp. 67-192. (#2.03)
- Eakle, A. S. and R. P. McLaughlin, 1919. Mono County, Masonic district: 15th Annual Rept. of the CA State Mineral., Pt. 1, pp. 160-165. (#4.048)
- Earl, K. M. 1950. Investigation of the Tapley copper deposit, Hancock County, Maine: U.S. Bur. Mines Rept. Inv. 4691, 7pp. (#3.01B, 3.01C)
- Einaudi, M. T., 1977. Petrogenesis of the copper-bearing skarn at the Mason Valley mine, Yerington district, Nevada: Econ. Geol., vol. 72, pp. 769-795. (#2.04)
- Einaudi, M. T., 1982. Skarns associated with porphyry plutons, <u>in</u> Advances in Geology of the Porphyry Copper Deposits of Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 139-183. (#2.04)
- Emigh, G. D., 1958. Petrography, mineralogy, and origin of phosphate pellets in the Phosphoria Formation: ID Bur. Mines and Geol. Pamph. 114, 60pp. (#6.02)
- Emmons, S. F., J. D. Irving, and G. F. Loughlin, 1927. Geology and ore deposits of the Leadville mining district, Colorado: U.S. Geol. Sur. Prof. Paper 148, 368p. (#4.01)
- Ensign, W. O., W. S. White, J. C. Wright, J. L. Patrick, R. J. Leone, D. J. Hathaway, J. N. Trammell, J. J. Fritts, and T. L. Wright, 1968. Copper deposits in the Nonesuch Shale, White Pine, Michigan, <u>in</u> Ore Deposits of the United States, 1933-1967, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 460-488. (#5.01)
- Eric, John H., 1948. Zinc-copper deposits of the Big Bend Mine, Butte County, California: CA Div. Mines Bull. 144, pp. 31-42. (#3.01B)

- Eric, John H., A. Stromquist, and C. Melvin Swinney, 1955. Geology and mineral deposits of the Angel Camp and Sonora quadrangles, Calaveras and Tuolumne Counties, California: CA Div. of Mines Special Rept. 41, 55pp. (#4.03)
- Farnham, L. L., 1961. Manganese deposits of New Mexico: U.S. Bur. Mines Inf. Circ. 8030, 176pp. (#4.10, #4.11)
- Farnham, L. L. and L. A. Stewart, 1958. Manganese deposits of western Arizona: U.S. Bur. Mines, Inf. Circ. 7843, 87pp. (#4.10)
- Farnham, L. L., L. A. Stewart, and C. W. DeLong, 1961. Manganese deposits of eastern Arizona: U.S. Bur. Mines Inf. Circ. 7990, pp. 93-184. (#4.11)
- Ferguson, H. G., 1927. Geology and ore deposits of the Mogollon mining district, New Mexico: U.S. Geol. Sur. Bull. 787, 100pp. (#5.4)
- Finch, Warren I., 1967. Geology of uranium deposits in sandstone in the United States: U.S. Geol. Sur. Prof. Paper 538, 121pp. (#5.02)
- Finch, Warren I., 1982. Preliminary concepts of a simple existence model for large sandstone uranium deposits, in Characteristics of Mineral Deposit Occurrences: U.S. Geol. Sur. Open File Rept. 82-795, pp. 201-4. (#5.02)
- Fisher, Frederick S., 1982. Porphyty copper deposits associated with the Needle Creek Igneous Center, Southern Absaroka Mountains, Wyoming, in the Genesis of Rocky Mountain Ore Deposits: Changes with Time and Tectonics: Proceedings of the Denver Region Exploration Geologists Society Symposium, Nov. 4-5, Denver, Golo. pp. 87-93. (#2.01)
- Geach, R. D., 1972. Mines and mineral deposits, Beaverhead County, Montana: MT Bur. Mines and Geol. Bull. 85, pp. 147-148. (#2.06)
- Geehan, R. W. and R. R. Trengrove, 1950. Investigation of Nevada Scheeltte, Inc. deposit, Mineral County, Nevada: U.S. Bur. Mines Rept. Inv. 4681. (#2.06)
- Gilmore, Paul and A. R. Still, 1968. The geology of the Iron King Mine, in Ore Deposits in the United States (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, vol. 2, pp. 1238-1257. (#3.01B)
- Goddard, E. N., 1940. Manganese deposits of Philipsburg, Granite County, Montana, a preliminary report: U.S. Geol. Sur. Bull. 922-G, pp. 157-204. (#4.11)
- Gower, H. D. and B. M. Madsen, 1964. The occurrence of phosphate rock in California: U.S. Geol. Sur. Prof. Paper 501-D, pp. D79-D85. (#6.02)
- Granger, H. E., M. M. Bell, G. C. Simmons, and Florence Lee, 1957. Geology and mineral resources of Elko County, Nevada: NV Bur. Mines Bull. 54, 190pp. (#4.04A)

- Graybeal, F. T., 1982. Geology of the El Tiro ore deposit, Silver Bell mining district, Arizona, in Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 487-505. (#2.01, #2.04)
- Greenan, J. 0., 1914. Geology of Fairview, Nevada: Eng. and Mng. Jour., vol. 97, pp. 791-793. (#4.04A)
- Greene, R. C., 1976. Volcanic rocks of the McDermott caldera, Nevada-Oregon: U.S. Geol. Sur. Open File Rept. 76-753, 80pp. (#2.02A)
- Gulbrandes, R. A. and D. G. Gielow, 1960. Mineral assemblage of a pyrometasomatic deposit near Tonopah, Nevada: U.S. Geol. Sur. Prof. Paper. 4008, pp. B20-B21. (#2.05)
- Gulbrandsen, R. A. and Donathan J. Krier, 1980. Large and rich phosphorus resources in the Phosphoria Formations in the Soda Springs area, Idaho: U.S. Geol. Sur. Buil. 1496, 25pp. (#6.02)
- Hallof, P. G. and E. Winniski, 1971. A geophysical case history of the Lakeshore ore body: Geophysics, vol. 36, pp. 1232-1249. (#2.01)
- Hanks, Henry G. and William Irelan, 1887. Calistoga silver mines: 6th Annual Report of the CA State Mineral., pp. 76-77. (#4.04A)
- Harrer, C. M. and F. J. Kelly, 1963. Reconnaissance of iron resources in New Mexico: U.S. Bur. Mines Inf. Circ. 8190, 112pp. (#2.03)
- Harris, N. B. and M. T. Einaudi, 1982. Skarn deposits in the Yerington district, Nevada, metasomatic skarn evolution near Lugwig: Econ. Geol., vol. 77, pp. 877-898. (#2.04)
- Harshman, E. N., 1965. Wilma claim and others, in Investigations of Molybdenum Deposits in Conterminous United States, 1942-60: U.S. Geol. Sur. Buil. 1182-E, 90pp. (#2.02A)
- Hausen, Donald M. and Paul F. Kerr, 1968. Fine gold occurrence at Carlin, Nevada, in Ore Deposits of the United States, 1933-1967, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 910-940. (#4.02)
- Henderson, C. W., 1926. Mining in Colorado; a history of discovery, development, and production: U.S. Geol. Sur. Prof. Paper 138, 263pp. (#6.01)
- Hernon, R. M. and W. R. Jones, 1968. Ore deposits of the Central mining district, New Mexico, in Ore Deposits of the United States, 1933-1967, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 1211-1238. (#2.03, #2.05)

- Heyl, George R., 1948(a). Ore deposits of Copperopolis, Calaveras County, California: CA Div. Mines Bull. 144, pp. 93-110. (#3.01B)
- Heyl, George R., 1948(b). The zinc-copper mines of the Quail Hill area, Calaveres County, California: CA Div. Mines Bull. 144, pp. 111-126. (#3.01B)
- Heyl, George R. and Manning W. Cox, 1948. Zinc deposits of the American Eagle-Blue Moon area, Mariposa County, California: CA Div. Mines Bull. 144, pp. 133-150. (#3.01B)
- Heyl, George R., Manning W. Cox, and John H. Eric, 1948. Fenn Zinc-copper Mine, Calaveras County, California: CA Div. Mines Bull. 144, pp. 61-84. (#3.01B)
- Heyl, George R. and John H. Eric, 1948. Newton Copper Mine, Amador County, California: CA Div. Mines Bull. 144, pp. 49-60. (#3.01B)
- Hill, J. M., 1916. Notes on some mining districts in eastern Nevada: U.S. Geol. Sur. Bull. 648, 214pp. (#4.04A)
- Hobbs, S. W., 1945. Tungsten deposits of Beaver County, Utah: U.S. Geol. Sur. Bull. 945-D, pp. 81-111. (#2.06)
- Hogue, W. G. and E. D. Wilson, 1951. Bisbee or Warren district, in Arizona Zinc and Lead Deposits, pt. 1: AZ Bur. Mines Geol. Ser. no. 18, Bull. no. 156, pp. 17-19. (#4.01)
- Hollister, V. F., 1979. Porphyry copper-type deposits of the Cascade volcanic arc., Washington: Mineral Science Eng., vol. 11, pp. 22-36. (#2.01)
- Hollister, V. F. and F. W. Baumann, 1978. The North Fork porphyry copper deposits of the Washington Cascades: Mineralium Deposita, vol. 13, pp. 191-199. (#2.01)
- Hotz, F. E., 1960. Diamond-drill exploration of the Dillsburg magnetite deposit, York County, Pennsylvania: U.S. Geol. Sur. Bull. 969-A, 25pp. (#2.03)
- Hudson, Travis, J. G. Smith, and R. L. Elliott, 1979. Petrology, composition and age of intrusive rocks associated with the Quartz Hill molybdenite deposit, Southeastern Alaska: Canadian Jour. of Earth Sciences, vol. 16, no. 9, pp. 1805-1822. (#2.028)
- Hudson, Travis, J. C. Arth, and K. C. Muth, 1981. Geochemistry of intrusive rocks associated with molybdenite deposits, Ketchikan quadrangle, Southesstern Alaska: Econ. Geol., vol. 76, pp. 1225-1232. (#2.028)
- Irvin, G. W., 1959. Pyrometasomatic deposits at Pima Mine: AZ Geol. Soc. Guidebook II, Southern Arizona, pp. 198-199. (#4.01)
- Jackson, Dan, 1982(a). Jerritt Canyon project: Eng. and Mng. Jour., July, pp. 54-58. (#4.02)

Jackson, Dan, 1982(b). Pinson gold: Eng. and Mng. Jour., Aug., pp. 64-68. (#4.02)

- Jahns, R. J., 1944. Beryllium and tungsten deposits of the Iron Springs district, Sierra and Socorro Counties, New Mexico: U.S. Geol. Sur. Bull. 945-C, pp. 45-79. (#2.06)
- James, L. P., 1976. Zoned alteration in limestone at porphyry copper deposits, Ely, Nevada: Econ. Geol. vol. 71, pp. 488-512. (#2.04)

James, L. P., 1978. The Bingham copper deposits, Utah, as an exploration target: history and pre-excavation geology: Econ. Geol. vol. 73, pp. 1218-1227. (#2.01)

- Johnson, Kenneth S., 1976. Permian copper deposits of southwestern Oklahoma, in Stratiform Copper Deposits of the Midcontinent Region: A Symposium, OK Geol. Sur. Circ. 77, pp. 3-14. (#5.01)
- Johnston, W. D., Jr., 1940. The gold quartz veins of Grass Valley, California: U.S. Geol. Surv. Prof. Paper 194, 101pp. (#4.03)

Johnston, W. P. and J. D. Lowell, 1961. Geology and origin of mineralized breccia pipes, Copper Basin, Arizona: Econ. Geol., vol. 56, pp. 916-940. (#2.01)

- Jones, E. L., Jr., 1916. A reconnaissance of the Kofa Mountains: U.S. Geol. Sur. Bull. 620-H, pp. 151-164. (#4.04A)
- Jones, E. L. and F. L. Ransome, 1919. Deposits of managanese ore in Arizona: U.S. Geol. Sur. 710-D, pp. 93-184. (#4.10, #4.11)
- Joralemon, P., 1951. The occurrence of gold at the Getchell Mine, Nevada: Econ. Geol., vol. 46, pp. 267-309. (#4.02)
- Kelley, V. C., 1949. Geology and economics of New Mexico iron-ore deposits: NM Univ. Publ., Geol. Ser. no. 2, 246pp. (#2.03)
- Kelley, V. C., 1952. Origin and pyrometasometic zoning of the Capitan iron deposit, Lincoln County, New Mexico: Econ. Geol. vol. 47, pp. 64-83. (#2.03)
- Kerr, P. F., 1934. Geology of the tungsten deposits near Mill City, Nevada: NV Univ. Bull. vol. 28, no. 2, 46pp. (#2.06)
- Kerr, P. F., H. M. Dahl, J. Green, and L. E. Woolard, 1957. Marysvale, Utah uranium area, geology, volcanic relations, and hydrothermal alteration: Geol. Soc. of Am. Spec. Paper 64, 212pp. (#A.09)
- Ketner, Keith B., 1963. Bedded barite deposits of the Shoshone Range, Nevada, in Geol. Sur. Research 1963: U.S. Geol. Sur. Prof. Paper 475-B, Art. 11, pp. B38-41. (#5.03)



- King, John R., 1982. Geology of the San Xavier North porphyry copper deposit, Pima mining district, Pima County, Arizona, in Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 475-483. (#2.01)
- Kinkel, A. R., W. E. Hall, and J. P. Albers, 1956. Geology and base metal deposits of the West Shasta copper-zinc district, Shasta County, California: U.S. Geol. Sur. Prof. Paper 285, 156pp. (#3,01C)
- Kinnison, John E., 1966. The Mission copper deposit, Arizona, <u>in</u> Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 281-287. (#2.01)
- Kirkemo, Harold, C. A. Anderson, and S. C. Creasey, 1965. Investigation of molybdenum deposits in the conterminous United States, 1942-60; U.S. Geol. Sur. Bull. 1182-E, pp. 1-90. (#2.02A, #2.02B)
- Klepper, M. R., R. A. Weeks, and E. T. Ruppel, 1957. Geology of the southern Elkhorn Mountains, Jefferson and Broadwater Counties, Montana: U.S. Geol. Sur. Prof. Paper 292, 82pp. (#2.04)
- Knopf, A., 1913. Ore deposits of the Helena mining region: U.S. Geol. Sur. Bull. 527, 143pp. (#2.03, #2.04)
- Knopf, Adolph, 1921. The Divide silver district, Nevada: U.S. Geol. Bull. 715-K, pp. 147-170. (#4.04A)
- Knopf, Adolph, 1929. The Mother Lode system of California: U.S. Geol. Sur. Prof. Paper 157, 88pp. (#4.03)
- Knopf, A. and P. Thelen, 1905. Sketch of the geology of Mineral King, California: Univ. CA Dept. Geol. and Sci. Bull. 4, pp. 227-262. (#2.05)
- Kolessar, Joseph, 1982. The Tyrone copper deposit, Grant County, New Mexico, in Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 327-333. (#2.01)
- Koschmann, A. H. and M. H. Bergendahl, 1968. Principal gold-producing districts of the United States: U. S. Geol. Sur. Prof. Paper 610, pp. 142-171. (#6.01)
- Koski, R. A. and D. S. Cook, 1982. Geology of the Christmas porphyry copper deposit, Gila County, Arizona, <u>in</u> Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 353-374. (#2.01)
- Kral, V. E., 1951. Mineral resources of Nye County, Nevada: NV Univ. Bull. 50, 223pp. (#4.02, #4.04A)
- Lamey, C. A., 1948. Shasta and California iron ore deposits, Shasta County, California: CA Div. Mines Bull. 129, 164pp. (#2.03)

Langlois, J. D., 1978. Geology of the Cyprus Pima Mine, Pima County, Arizona: AZ Geol. Soc. Digest, vol. 9, pp. 103-113. (#2.04)

- Langon, J. M. and S. N. Williams, 1982. Structural, petrological, and mineralogical controls for the Dos Pobres ore body, <u>in</u> Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 335-352. (#2.01)
- Lapham, D. M., 1968. Triassic magnetite and diabase of Cornwall, Pennsylvania, in Ore Deposits of the United States, 1933-1967 (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 72-94. (#2.03)
- La Point, Dennis, 1976. A comparison of selected sandstone copper deposits in New Mexico, in Stratiform Copper Deposits of the Midcontinent Region, A Symposium: OK Geol. Sur. Circ. 77, pp. 80-96. (#5.01)
- Lausen, Carl, 1942. The Oatman and Katherine districts, Arizona, <u>in</u> Ore Deposits as Related to Structural Features: Princeton, NJ, Princeton Univ. Press, pp. 226-229. (#4.04A)
- Leonardson, R. W., 1982. Preliminary geology and molybdenum deposits at Questa, New Mexico, in the Genesis of Rocky Mountain Ore Deposits, Changes with Time and Tectonics: Denver Region Exploration Geologists Society Symposium, Nov. 4-5, 1982, pp. 151-155. (#2.02A)
- Lindgren, Waldemar, 1900. The gold and silver veins of Silver City, DeLamar and other mining districts in Idaho: U.S. Geol. Sur. 20th Annual Rept. 1898-99 (part III), pp. 67-256. (#4.05)
- Lindgren, Waldemar, 1911. Tertiary gravels of the Sierra Nevada of California: U.S. Geol. Sur. Prof. Paper 73, 226pp. (#6.01)
- Lindgren, Waldemar, 1915. Geology and mineral deposits of the National mining district, Nevada: U.S. Geol. Sur. Bull. 601, 58pp. (#4.04Å)
- Lindgren, Waldemar, and Howland Bancroft, 1914. Republic (Eureka) district, in The Ore Deposits of Northeastern Washington: U.S. Geol. Sur. Bull. 550, pp. 133-210. (#4.04A)
- Lindgren, Waldemar, L. C. Graton, and C. H. Gordon, 1910. The ore deposits of New Mexico: U.S. Geol. Sur. Prof. Paper 68, 361pp. (#4.04A)
- Lindsey, D. A., 1977. Epithermal beryllium deposits in water-laid tuff in western Utah: Econ. Geol. vol. 72, pp. 219-232. (#4.09)
- Lowell, J. David, 1968. Geology of the Kalamazoo orebody, San Manuel district, Arizona: Econ. Geol., vol. 63, pp. 650-653. (#2.01)
- Ludington, Steve, 1982. Granite molybdenum systems: U.S. Geol. Sur. Open File Rept. 82-795, pp. 43-46. (#2.02A)

- Lynch, Dean W., 1966. The economic geology of the Esperanza Mine and vicinity, in Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ, pp. 267-279. (#2.01)
- MacKenzie, W. B. and A. A. Bookstrom, 1976. Geology of the Majuba Hill area, Pershing County, Nevada: NV Bur. Mines and Geol. Bull. 86, 23pp. (#2.02A)
- Mackin, J. H., 1968. Iron ore deposits of the Iron Springs deposit, southwestern Utah, in Ore Deposits of the United States, 1933-1967 (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 992-1019. (#2.03)
- McKnight, J. F. and M. L. Fellows, 1978. Silicate mineral assemblages and their relationship to sulfide mineralization, Pinos Altos mineral deposit, New Mexico: AZ Geol. Soc. Digest, vol. 11, pp. 1-8. (#2.04)
- McLaughlin, D. H., 1931. Geology of the Homestake Mine, Black Hills, South Dakota: Eng. and Min. Jour. 132, pp. 324-329. (#4.01)
- Menzie, W. D. and T. G. Theodore, 1983. Molybdenum porphyry (low F type), mineral deposit grade-tonnage models: U.S. Geol. Sur. Open File Rept. 83-623, p. 31. (#2.028)
- Merriam, C. W. and C. A. Anderson, 1942. Reconnaissance survey of the Roberts Mountains, Nevada: Geol. Soc. Am. Bull., vol. 53, pp. 1675-1728. (#2.02A)
- Metz, Robert A. and Arthur W. Rose, 1966. Geology of the Ray copper deposit, Arizona, in Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 177-188. (#2.01)
- Meyer, Charles, Edward P. Shea, and Charles C. Goddard, Jr., 1968. Ore deposits at Butte, Montana, <u>in</u> Ore Deposits of the United States, 1933-1967, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 1373-1416. (#2.01)
- Moolick, R. T. and J. J. Durek, 1966. The Morenci district, <u>in</u> Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 221-232. (#2.01, #2.04)
- Morris, H. T. and T. S. Lovering, 1979. General geology and mines of the East Tintic mining district, Utah and Jaub Counties, Utah: U.S. Geol. Sur. Prof. Paper 1024, 203pp. (#4.01)
- Mosier, D. L., 1983a. Subaerial volcanogenic manganese, mineral deposit grade-tonnage models II: U.S. Geol. Sur. Open File Rept. 83-902, p. 65. (#4.10)
- Mosier, D. L., 1983b. Carbonate-hosted manganese replacement, mineral deposit grade-tonnage models II: U.S. Geol. Sur. Open File Rept. 83-902, p. 68. (#4.11)

- Mosier, D. L., 1983c. Marine phosphate upwelling-type, mineral deposit grade-tonnage models II: U.S. Geol. Sur. Open File Rept. 83-902, p. 85. (#6.02)
- Mosier, D. L., 1983d, Zinc-lead skarn, mineral deposit grade-tonnage models II: U.S. Geol. Sur. Open File Rept. 83-902, p. 26. (#2.05)
- Mosier, D. L. and W. D. Menzie, 1983. Epithermal gold, quartz-alunite type, mineral deposit grade-tonnage models: U.S. Geol. Sur. Open File Rept. 83-662, p. 90. (#4.048)
- Mosier, D. L., D. A. Singer, and B. B. Salem, 1983. Geologic and grade-tonnage information on volcanic-hosted copper-zinc-lead massive sulfide deposits: U.S. Geol. Sur. Open File Rept. 83-89, 78pp. (#3.01B, #3.01C)
- Mutschler, F. E., 1976. Crystalization of a soda granite, Treasure Mountain Dome, Colorado and the genesis of stockwork molybdenite deposits: <u>in NM</u> Geol. Soc. Spec. Paper 6, pp. 199-205. (#2.02A)
- Nash, J. T. and T. G. Theodore, 1971. Ore fluids in the porphyry copper deposit at Copper Canyon, Nevada: Econ. Geol., vol. 66, p 385. (#2.01)
- Nash, J. T., H. C. Granger, and S. S. Adams, 1981. Geology and concepts of genesis of important types of uranium deposits, <u>in</u>: Econ. Geol. (75th Annual Volume) pp. 63-116. (#5.02)
- Nason, Philip W., Allen V. Shaw, and Kent D. Aveson, 1982. Geology of the Poston Buttes porphyry copper deposit, Pinal County, Arizona, <u>in</u> Advances in Geology of the Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 375-385. (#2.01)
- Needham, A. B. and R. R. Trengrove, 1950. Investigation of Black Diablo, Black Eagle, and Black Rock manganese deposits, Persning and Lander Counties, Nevada: U.S. Bur. Mines, Rept. Inv. 471, 37pp. (#3.02)
- Newberry, R. J., 1982. Tungsten-bearing skarns of the Sierra Nevada, I. The Pine Creek Mine, California: Econ. Geol. vol. 77, pp. 823-844. (#2.06)
- Nielson, R. L., 1970. Mineralization and alteration in calcareous rocks near the Santa Rita Stock, New Mexico: NM Geol. Soc. Guidebook, 21st Field Cong., pp. 133-139. (#2.04)
- Nokleberg, W. J., 1981. Geologic setting, petrology, and geochemistry of zoned tungsten skarns at the Strawberry Mine, central Sierra Nevada, California: Econ. Geol., vol. 76, pp. 111-133. (#2.06)
- Nolan, T. B., 1936. The Tuscarora mining district, Elko County, Nevada: NV Univ. Bull. 25, 49pp. (#4.04A)

- Nolan, T. B. and R. N. Hunt, 1968. The Eureka mining district, Nevada, in Ore Deposits of the United States 1933-1966, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 961-966. (#4.01)
- Nuelle, L. M., P. D. Proctor, and S. K. Grant, 1979. Volcanology and mineralization, Ohio and Mt. Baldy mining district, Marysvale, Piute County, Utah: Geol. Soc. Am. Abstracts with Program, vol. 11, no. 6, pp. 272-273. (#2.02A)
- Olmstead, Hugh W. and David W. Johnson, 1966. Inspiration geology, <u>in</u> Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of Az. pp. 143-150. (#2.01)
- Orris, G., 1983. Bedded barite, mineral deposit grade-tonnage models II: U.S. Geol. Sur. Open File Rept. 83-902, p. 51.
- Overton, T. D., 1947. The mineral resources of Douglas, Ormsby, and Washoe Counties, Nevada: NV Univ. Bull. 46, 91pp. (#4.04B)
- Pardee, J. T. and C. F. Clark, Jr., 1948. Gold deposits of the southern Piedmont: U.S. Geol. Sur. Prof. Paper 213, 156pp. (#4.03)
- Park, Charles F., 1942. Manganese resources of the Olympic Penninsula, Washington: U.S. Geol. Sur. Bull. 931-R, p. 435-457. (#3.02)
- Patton, W. W., Jr. and J. J. Matzko, 1959. Phosphate deposits of northern Alaska: U.S. Geol. Sur. Prof. Paper 302-A, pp. 1-17. (#6.02)
- Perry, V. D., 1969. Skarn genesis at the Christmas Mine, Gila County, Arizona: Econ. Geol., vol. 64, pp. 225-270. (#2.04)
- Peterson, N. P., 1962. Geology and ore deposits of the Globe-Miami district, Arizona: U.S. Geol. Sur. Prof. Paper 342, 88pp. (#2.01)
- Popoff, C. C. and A. L. Service, 1965. An evaluation of the western phosphate industry and its resources, part 2, Montana: U.S. Bur. Mines Rept. Inv. 6611, 146pp. (#6.02)
- Prinz, W. C., 1961. Geology and ore deposits of the Phillipsburg district, Granite County, Montana: U.S. Geol. Sur. Bull. 1237, 66pp. (#4.01)
- Ransome, Alfred L. and John L. Kellogg, 1939. Quicksilver resources of California: CA Jour. Mines and Geol. vol. 35, no. 4, pp. 353-486. (#4.07)
- Ransome, F. L., 1909. The geology and ore deposits of Goldfield, Nevada: U.S. Geol. Sur. Prof. Paper 66, 258pp. (#4.04B)
- Ranta, D. E., 1974. Geology, alteration, and mineralization of the Winfield (LaPlata) district, Chaffee County, Colorado: unpub. Ph.D. thesis, CO Sch. Mines, Colden, CO. 261pp. (#2.02A)

Reeves, R. G., F. R. Shawe, and V. F. Kral, 1958. Iron ore deposits of westcentral Nevada: NV Bur. Mines Bull. 53. (#2.03)

- Reiter, B. E., 1981. Controls on lead-zinc mineralization, Iron Cap mine area, Aravaipa district, Graham County, Arizona: Geol. Soc. Am. Bull. vol. 13, p. 103. (#2.05)
- Richard, K. and J. H. Courtright, 1959. Some geologic features of the Mission copper district: AZ Geol. Soc. Guidebook 2, pp. 201-204. (#2.04)
- Rinehart, C. D. and D. C. Ross, 1956. Economic geology of the Casa Diablo Mountain quadrangle, California: CA Div. Mines Spec. Rept. 48, 17 pp. (#2.06)
- Robinson, R. F. and Annan Cook, 1966. The Safford copper deposit, Lone Star mining district, Graham County, Arizona, in Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 251-266. (#2.01)
- Roper, M. W. and A. B. Wallace, 1981. Geology of the Aurora uranium prospect, in Uranium in Volcanic and Volcaniclastic Rocks: Am. Assoc. of Pet. Geologists, Studies in Geol., no. 13, pp. 81-88. (#4.09)
- Rose, Arthur W. and Will W. Baltosser, 1966. The porphyry copper deposits at Santa Rita, New Mexico, in Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ. pp. 205-220. (#2.01)
- Ross, Clyde P., 1941. Some quicksilver prospects in adjacent parts of Nevada, California, and Oregon: U.S. Geol. Sur. Bull. 931-B, pp. 33-36. (#4.08)
- Ross, Clyde P., 1942. Quicksilver deposits in the Steen and Pueblo Mountains, southern Oregon: U.S. Geol. Sur. Bull. 931-J, pp. 227-258. (#4.06, #4.08)
- Ross, C. P., 1956. Quicksilver deposits near Weiser, Washington County, Idaho: U.S. Geol. Sur. Bull. 1042-D, pp. 79-104. (#4.08)
- Ross, C. P., 1960. Geology of Glacier National Park and the Flathead region, northwestern Montana: U.S. Geol. Sur. Prof. Paper 296, 125pp. (#4.04A)
- Ross, C. P., 1961. Geology and mineral deposits of Mineral County, Nevada: Nv. Bur. Mines Buil. 58, 98pp. (#4.04A)
- Rostad, O. H., 1966. Geochemical case history at the Little Falls molybdenite deposit, Boise County, Idano: Canada Geol. Sur. Paper 66-54, pp. 248-252. (#2.02A)
- Rostad, O. H., 1969. The use of geochemistry at the Bald Butte molybdenite prospect, Lewis and Clark County, Montana: CO Sch. Mines Quart., vol. 64, no. 1, pp. 437-449. (#2.02A)

- Rostad, O. H., 1971. Offset geochemical anomalies at the Ima Mine, Lemhi County, Idaho: Canadian Inst. Mining, Metall. Spec. vol. 11, pp. 241-246. (#2.02A)
- Rott, E. H., Jr., 1931. Ore deposits of the Gold Circle mining district, Elko, Nevada: NV Bur. Mines Bull. 58, 98pp. (#4.04A)
- Rubright, R. D. and O. J. Hart, 1968. Non-porphyry ores of the Bingham district, Utah, in Ore Deposits of the United States 1933-1967, (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 886-907. (#4.01)
- Rye, Robert O., Ralph J. Roberts, Walter S. Snyder, G. Larry Lahusen, and John E. Motica, 1984. Textural and stable isotope studies of the Big Mike cupriferous volcanogenic massive sulfide deposit, Pershing County, Nevada: Econ. Geol. vol. 79, pp. 124-140. (#3.01A)
- Rytuba, J. J. and R. K. Glauzman, 1979. Relation of mercury, uranium, and lithium deposits to the McDermitt caldera complex, Nevada-Oregon <u>in</u> Papers on Mineral Deposits of Western North America: NV Bur. Mines and Geol. Rept. 33, pp. 109-117. (#4.09)
- Samson, R. J. and W. B. Tucker, 1940. Mineral resources of Mono County, California: CA Jour. Mines Geol., vol. 56, pp. 117-156. (#2.05)
- Savage, C. N., 1961. Economic geology of central Idaho blacksand placers: ID Bur. Mines and Geol. Bull. 17, 160pp. (#6.01)
- Schenk, C. G., 1977. Discovery of the Flambeau deposit, Rusk County, Wisconsin: Geoscience WI, Univ. WI. Ext., Geol. Nat. Hist. Sur. vol. 1, pp. 27-42. (#3.01C)
- Schilling, J. H., 1962. An inventory of molybdenum occurrences in Nevada: NV Bur. Mines Rept. 2, 48pp. (#2.02B)
- Schmidt, E. A., 1978. Geology of the Thompson Creek molybdenum deposit, Custer County, Idaho: Preprint Paper presented at Northwest Mining Assoc. Meeting, Nov. 30 - Dec. 2, Spokane, WA. 2pp. (#2.02B)
- Schmidt, H. and J. Worthington, 1977. Molybdenum mineralization at Cannivan Gulch: Geol. Assoc. Canada Abs., vol. 2. p. 46. (#2.02B)
- Schmitt, L. J, and W. H. Raymond, 1977. Geology and mineral deposits of the Needle Mountains district, southwestern Colorado: U.S. Geol. Sur. Bull. 1434, 40 pp. (#2.02A)
- Schrader, F. C., 1923. The Jarbidge mining district, Nevada: U.S. Geol. Sur. Bull. 741, 86pp. (#4.04A)
- Scull, B. J., 1958. Origin and occurrence of barite in Arkansas: AR Geol. and Conserv. Commission Inf. Circ. 18, 101pp. (#5,03)

Segerstrom, K. and E. J. Young, 1972. General geology of the Hahns Peak and Farwell Mountain quadrangle: U.S. Geol. Sur. Bull. 1349, 63pp. (#2.02A)

- Senter, L. E., 1976. Geology and porphyry molybdenum mineralization of the Turnley Ridge Stock in the Elkhorn mining district, Jefferson County, Montana: unpub. M.S. Thesis, Cheney Eastern Washington Univ., 76pp. (#2.02B)
- Shannon, S. S. Jr., 1971. Evaluation of copper and molybdenum geochemical anomalies at the Cumo prospect, Boise County, Idaho: Canadian Inst. Mining, Metall. Spec. vol. 11, pp. 247-250. (%2.02A)
- Sharp, James E. 1978. A molybdenum mineralized breccia pipe complex, Redwell Basin: Econ Geol., vol. 73, pp. 369-382. (#2.02A)
- Sharp, James E. 1979. Cane Peak, a molybdenum-mineralized breccia pipe complex in Culberson County, Texas: TX Econ. Geol., vol 74, pp. 517-534. (#2.02A)
- Shawe, D. R., F. G. Poole, and D. A. Probst, 1969. Newly discovered bedded barite deposits in East Northumland Canyon, Nye County, Nevada: Econ. Geol. vol. 64, po. 245-54. (#5.03)
- Sheldon, R. P., 1963. Physical stratigraphy and mineral resources of Permian rocks in western Wyoming: U.S. Geol. Sur. Prof. Paper 313-B, pp. 49-273. (#6.02)
- Shenon, Philip J., 1933. Geology and ore deposits of the Takilma-Waldo district, Oregon, including the Blue Creek district: U.S. Geol. Sur. Bull. 866, pp. 14-194. (#5.01A)
- Sherborne, J. E., Jr., W. A. Buckovic, D. B. Dewitt, T. S. Hellinger, and S. J. Pavlak, 1979. Major uranium discovery in volcaniclastic sediments, Basin and Range Providence, Yavapai County, Arizona: Am. Assoc. Pet. Geologists Bull., vol. 63, pp. 621-646. (#4.09)
- Sigurdson, D. R., 1974. Mineral paragenesis and fluid inclusion themometry at four Western U.S. tungsten deposits: unpub. Ph.D. thesis, Univ. Ca., Riverside, CA. 214pp. (#2.06)
- Silberman, M. L., C. W. Chesterman, F. J. Kleinhampl, and C. H. Gray, 1972. K-Ar ages of volcanic rocks and gold-bearing quartz-adularia veins in the Bodie mining district, Mono County, California: Econ. Geol. vol. 67, pp. 597-604. (#4.04A)
- Simon, F. S., 1972. Mesozoic stratigraphy of the Patagonia Mountains and adjoining areas, Santa Cruz County, Arizona: U.S. Geol. Sur. Prof. Paper 658-E, pp. El-E23. (#2.05)
- Simmon, W. W. and J. E. Fowells, 1966. Geology of the Copper Cities Mine, in Geology of Porphyry Copper Deposits, Southwestern North America: Tucson, AZ, Univ. of AZ, pp. 151-156 (#2.01)

Simms, P. K., 1958. Geology and magnetite deposits of Dover district, Morris County, New Jersey: U.S. Geol. Sur. Prof. Paper 287, 162pp. (#2.03)

- Simms, S. J., 1968. The Grace Mine magnetite deposits, Beck County, Pennsylvania, in Ore Deposits of United States, 1933-1967 (Graton-Sales Volume): New York, Am. Inst. Mining, Metall., and Petroleum Engineers, pp. 108-124. (#2.5)
- Singer, D. A. and D. L. Mosier, 1983. Porphyry copper mineral deposit grade-tonnage models: U.S. Geol. Sur. Open File Rept. 83-623, pp. 21-22, 52. (#2.01, #3.01A)
- Smith, Gary E., 1976. Sabkla and tidal-flat facies control of stratiform copper deposits in north Texas, <u>in</u> Stratiform Copper Deposits of the Midcontinent Region, A Symposium: OK Geol. Sur. Circ. 77, pp. 25-39. (#5.01)
- Smith, J. F., Jr., A. H. Wadsworth, Jr., and J. R. Cooper, 1945. San Pedro and Carnahan Mines, New Placer mining district, Santa Fe County, New Mexico: U.S. Geol. Sur. Open File Rept. (#2.04)
- Smith, L. L., 1931. Magnetite deposits of French Creek, Pennsylvania: PA Geol. Sur. 4th Ser. Bull. M14, 52pp. (#2.03)
- Sorem, Ronald K. and Donald W. Gunn, 1967. Mineralogy of manganese deposits, Olympic Peninsula, Washington: Econ. Geol. vol. 62, pp. 22-56. (#3.02)
- Sorenson, R. E., 1927. The geology and ore deposits of the South Mountain mining district, Owyhee County, Idaho: ID Bur. Mines Geol. Pamph. 22, 47 pp. (#2.05)
- South, D. L., 1972. Sulfide zoning at the Lakeshore copper deposit, Pinal County, Arizona: unpub. M.S. thesis, Univ. AZ, Tucson, AZ, 79pp. (#2.04)
- Spurr, J. E., 1908. Descriptive geology of Nevada south of the 40th parallel and adjacent portions of California: U.S. Geol. Sur. Bull. 208, p. 181. (#4.04A)
- Staatz, M. H. and W. L. Carr, 1964. Geology and mineral deposits of the Thomas and Dugway Ranges, Juab and Toole Counties, Utah. U.S. Geol. Sur. Prof. Paper 415, 188pp. (#4,09)
- Stanford, Warren D., 1984. Alligator Ridge: From a lone prospector's discovery to an operating gold mine: Mng. Eng., July, pp. 593-598. (#4.02)
- Stevens, T. A. and J. C. Ratte, 1960(a). Relations of mineralization to caldera subsidence in the Creede district, San Juan Mountains, Colorado, in Geological Survey Research in 1960: U.S. Geol. Sur. Prof. Paper 400-B, pp. 14-7 (#4.06A)

- Stevens, T. A. and J. C. Ratte, 1960(b). Geology and ore deposits of the Summitville district, San Juan Mountains, Colorado: U.S. Geol. Sur. Prof. Paper 343, 70pp. (#4.048)
- Stinson, M. C., 1957. Geology of the Island Mountain Copper Mine, Trinity County, California: CA Jour. Mines and Geology, vol. 53, no. 1, pp. 9-33. (#3.01A)
- Stoddard, Carl and J. A. Carpenter, 1950. Mineral resources of Storey and Lyon Counties, Nevada: NV Univ. Bull. 49, 115pp. (#4.04A)
- Storey, Lester O., 1978. Geology and mineralization of the Lights Creek Stock, Plumas County, California: AZ Geol. Soc. Digest, vol. XI, pp. 49-58. (#2.01)
- Stroud, R. B., A. B. McMahan, R. K. Stroup, and M. H. Hibpshman, 1970. Production potential of copper deposits associated with Permian red bed formations in Texas, Oklahoma, and Texas: U.S. Bur. Mines, Rept. Inv. 7422, 103pp. (#5.01)
- Taliaferro, N. L. and F. S. Hudson, 1943. Genesis of the manganese deposits of the coast ranges of California, in Manganese in California: CA Div. of Mines Bull. 125, pp. 217-275. (#3.02)
- Taylor, B. E. and J. R. O'Neil, 1977. Stable isotope studies of metasomatic Ca-Fe-Al-Si skarns and associated metamorphic and igneous rocks, Osgood Mountains, Nevada: Contrib. to Mineralogy and Petrology, vol. 63, pp. 1-49. (#2.06)
- Taylor, R. B. and R. V. King, 1967. Preliminary report on mid-Tertiary rhyolite vents and associated mineralization south of Georgetown, Colorado: U.S. Geol. Sur. Open File Rept., 15p. (#2.02A)
- Theodore, Ted G., 1982. Preliminary model outline for fluorine-deficient porphyry molybdenum deposits, <u>in</u> Characteristics of Mineral Deposit Occurrences: U.S. Geol. Sur. Open File Rept. 82-745, pp. 37-42. (#2.02B)
- Thomas, J. A. and J. T. Galey, Jr., 1978. Mt. Emmons, Colorado, molybdenum deposit (abs.): Am. Inst. Mining, Metall., and Petroleum Engineers, 107th Annual Meeting, p. 44. (#2.02A)
- Tingley, Joseph V. and Byron R. Berger, 1985. Lode gold deposits of Round Mountain, Nevada: NV Bur. Mines and Geology Bull. 100, 32pp. (#4.05)
- Titley, S. R., 1961. Genesis and control of the Linchburg ore body. Socorro County, New Mexico: Econ. Geol., vol. 56, pp. 695-722. (#2.05)
- Trask, P. D., Ivan F. Wilson, and Frank S. Simmons, 1943. Manganese deposits of California, a summary report, in Manganese in California: CA Div. of Mines, Bull. 125, pp. 51-215. (#3.02)

- Troxel, B. W. and P. K. Morton, 1962. Mojave mining district: CA Div. Mines County Rept. 1, pp. 43-45. (#4.04B)
- Tweto, Odgen, 1974. Geologic map and sections of the Holy Cross quadrangle, Eagle, Lake, Pitkin, and Summit Counties, Colorado: U.S. Geol. Sur. Misc. Geol. Inv. Map 1-830. (#2.02A)
- Tweto, Ogden and T. S. Lovering, 1947. The Gilman district, Eagle County, Colorado: in Mineral resources of Colorado, Denver, CO, Mineral Resources Board, pp. 378-387. (#4.01)
- Umpleby, J. B., 1917. Geology and ore deposits of the Mackay region, Idaho: U.S. Geol. Sur. Prof. Paper 97, 129pp. (#2.04)
- Vanderburg, W. O., 1936. Reconnaissance of mining districts in Pershing County, Nevada: U.S. Bur. Mines Inf. Cir. 6902, pp. 37-9. (#5.4)
- Vanderburg, W. O., 1937. Reconnaissance of mining districts in Mineral County, Nevada: U.S. Bur. Mines Inf. Cir. 6941, pp. 13-16. (#5.4)
- Vanderburg, W. O., 1940. Reconnaissance of mining districts in Churchill County, Nevada: U.S. Bur. Mines Inf. Circ. 7093, pp. 40-1. (#4.04A)
- Wallace, A. B., 1980. Geology of the Sulphur district, southwestern Humboldt County, Nevada: Soc. of Econ. Geol. Field Conf. (1980) Road Log and Articles, pp. 80-91. (#4.05)
- Wallace, S. R., 1978. Geology of the Urad and Henderson molybdenum deposits. Clear Creek County, Colorado: Econ. Geol., vol. 73, pp. 324-368. (#2.02A)
- Waugh, Truman S. and Laurence L. Brady, 1976. Copper occurrences associated with Permian rocks in south-central Kansas, <u>in</u> Stratiform Copper Deposits of the Midcontinent, a Symposium: OK Geol. Sur., Circ. 77, pp. 76-79. (#5.01)
- Weber, H. F., Jr. and R. A. Mathews, 1967. Prospecting for barite in northern Shasta County: GA Div. Mines and Geology Mineral Inf. Service, vol. 20, no. 9, pp. 107-114. (#5.03)
- Wells, F. G. and A. C. Waters, 1934. Quicksilver deposits of southwestern Oregon: U.S. Geol. Sur. Bull. 850, 58pp. (#4.06)
- Weils, J. D., L. D. Stoiser, and J. E. Elliott, 1969. Geology and geochemistry of the Cortez gold deposit, Nevada: Econ. Geol., vol. 64, pp. 526-537. (#4.02)
- West, Richard J. and Daniel M. Aiken, 1982. Geology of the Sierrita-Esperenza deposit, Pima mining district, Pima County, Arizona, <u>in</u> Advances in Geology of the Porphyry Copper Deposits, SouthWestern North America: Tucson, AZ, Univ. of AZ. pp. 433-465. (#2.01)

White, D. and C. E. Roberson, 1962. Sulfur Bank, California, a major hot-spring quicksilver deposit: Geol. Soc. Am. (Buddington Volume), pp. 397-428. (#4.08)

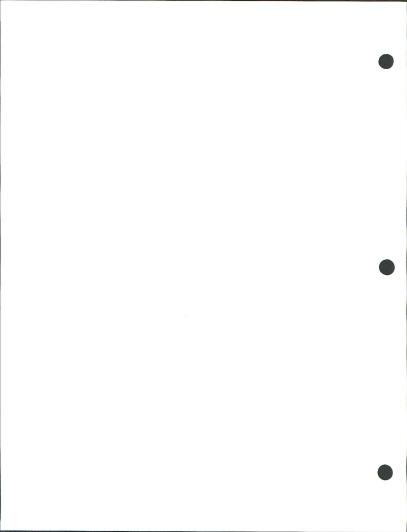
- White, W. H., A. A. Bookstrom, R. J. Kamilli, M. W. Ganster, R. P. Smith, D. E. Ranata, and R. C. Steininger, 1981. Character and origin of Climax-type molybdenum deposits: Econ. Geol. (75th Anniv. Volume), pp. 270-316. (#2.02A)
- Willden, R., 1964. Geology of the Christmas quadrangle, Gila and Pinal Counties, Arizona: U.S. Geol. Sur. Bull. 1161-E, pp. E1-E64. (#4.01)
- Wilson, Eldred D., J. B. Cunningham, and G. M. Butler, 1967. Arizona lode gold mines and gold mining; AZ Bur. Mines Bull. 137, pp. 143-147. (#4,06A)
- Wilson, J. R., 1963. Geology of the Yerington Mine: Mng. Cong. Jour., vol. 49, no. 6, pp. 30-34. (#2.01)

Witkind, I. J., 1973. Igneous rocks and related mineral deposits of the Baker quadrangle, Little Belt Mountains, Montana: U.S. Geol. Sur. Prof. Paper 752, 58pp. (#2.02A)

Wright, L. A., Richard M. Stewart, Thomas E. Gay, Jr., and George C. Hazenbush, 1953. Mines and mineral deposits of San Bernardino County, California: CA Jour. Mines and Geol., vol. 49, pp. 1 & 2, pp. 71-72. (#2.03, #4.048)

Yates, Robert G., 1942. Quicksilver deposits of the Opalite district, Malheur County, Oregon and Humboldt County, Nevada: U.S. Geol. Sur. Bull. 931-N, pp. 319-348. (#4.08)

Young, R. R., 1963. Prospect evaluations, Washington County, Maine: ME Geol. Sur. Spec. Econ. Ser. no. 3, 86 pp. (#3.01B, #3.01C)



GLOSSARY OF TERMS

ANKERITIZATION. describes the introduction or replacement by ankerite. Sometimes the term ferroan dolomite is used instead.

ARGILLIC ALTERATION. the formation of new clay minerals in silicate rocks subjected to predominant acid hydrothermal solutions. The new minerals are usually kaolinite, dickite, and under some conditions pyrophyllite.

BACK-ARC BASIN. a term that describes areas lying between the active mountain building at the continental margins and the relatively stable cratonic regions in the continental interior.

CRATON. the relatively stable central portion of a continent that has been little deformed for prolonged periods.

EPIGENETIC MINERAL DEPOSIT. a deposit that has formed later than the enclosing rocks.

EUGEOSYNCLINE. the volcanic part of an orthogeosyncline located away from the craton.

GOSSAN. An iron-bearing weathered product formed by the oxidization of sulfides and overlying a sulfide deposit.

GREISEN VEINS. veins that occur in pneumatolytically altered granitic rocks and are composed mainly of quartz, muscovite or lepidolite, and topaz. Accessory minerals may include tourmaline, fluorite, rutile, cassiterite, and wolframite.

ISLAND ARC. a curved chain of islands rising from the deep-sea floor and nearer to the continents. Its curve is generally convex toward the open ocean.

INTERCRATONIC TROUGH. a graben-like structure that forms within a craton area.

LIME SILICATE ALTERATION. the replacement of carbonate sediment at an intrusive contact by lime-silicate minerals. Usually composed of various amounts of garnet, iron-rich pyroxene, epidote, wollastonite, and scapolite. The resulting rock is called skarn or tactite.

MAGMATIC CUPOLA. an upward projection of magma into the overlying rocks.

MIGGEOSYNCLINE. the nonvolcanic part of an orthogeosyncline located near the craton. Volcanism is not associated with the sediments.

OPHIOLITE. a group of mafic and ultramafic igneous rocks ranging from spillte and baselt to gabbro and peridotite, including rocks rich in serpentine, chloride, epidote, and albite derived from them by latter metamorphism, whose origin is associated with an early phase of the development of a geosyncline. ORE DEPOSITS. A general term applied to rocks containing one or more minerals of economic value in such amounts and of such grade that they can be profitably exploited. This is different from a "mineral deposit" in that it is not yet known whether it can be profitably exploited.

ORE MINERAL. A mineral that contains a valuable or desired constituent (metallic or nonmetallic).

ORTHOGEOSYNCLINE. a regional geologic structure that contains both volcanic and nonvolcanic belts that form between continental cratons and ocean basins.

POTASSIC ALTERATION. a process that occurs as a result of greater or lesser potassium metasomatism and may be accompanied by more or less leaching of calcium and sodium in rocks containing original aluminosilicate minerals. Sometimes called K-silicate alteration.

PROPULITIC ALTERATION. a process that is characterized chiefly by development of new calcium and magnesium minerals in igneous rocks by rearrangement of original rock-forming components by the hydrothermal process. The assemblage of minerals includes the formation of propylite, carbonates, epidate, quartz, and chlorite.

QUARTZ SERICITE ALTERATION. a process caused by the leaching of sodium, calcium, and magnesium from alumino-silicate-bearing rocks, whereas potassium may be introduced or derived from original rock-forming feldspar. Sometimes called phyllic alteration.

RESURGENT DOME. a dome that has formed by uplift following subsidence of a cauldron block.

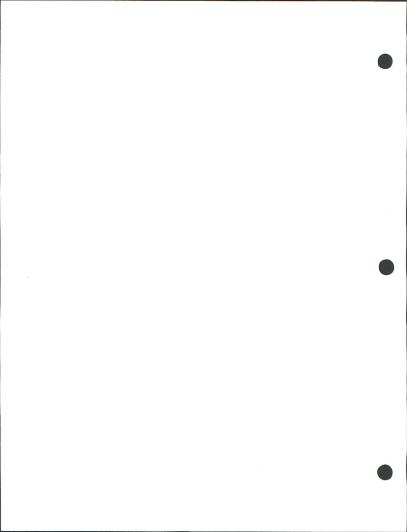
SUPERGENE ENRICHMENT. the process where near-surface oxidation produces acidic solutions that leach metals, carry them downward, and reprecipitate them, thus enriching sulfide minerals already present.

SYNGENETIC DEPOSITS. a mineral deposit formed contemporaneously with, and by essentially the same process as, the enclosing rocks.

VOLCANICLASTIC. a clastic rock containing volcanic material in whatever proposition and without regard to its origin or environment.

Appendix

The following table is a quick reference guide to assist the user in tentatively identifying mineral deposit types.



	MINERAL D	EPOSTI TIPES	
	#2.01	#2.02A	#2.02B
Regional Setting	Continental margins hav- ing abundant igneous intrusives. Mesozoic to Tertiary in age.	Continental interior in belts of Cretaceous- Tertiary tectonic activity	Mainly on continental margins and areas of back-arc spreading
Structural Features	Most spatially or genet- lcally related to intru- sives that form stocks, plugs, sills, or dikes. Breccia and brecciation pipes common.	Commonly occur in clus- ters. Tend to be dome- shaped and centered on Intrusive cupolas.	Generally related to the tectonic grain of Meso- zoic batholiths
Stratigraphic and Lithographic Characteristics	Wide range of intrusive rocks. Most favorable are qtz. monzonite, granodiorite, and tona- lite. Intruded carbonate rocks particularly fa- vorable.	High silica, alkali-rich porphyry granites favor- ed especially those hav- ing >250FPM Rb, >50PPM Nb, 5 0PPM Sr, <300PPM Ba, and enriched in flu- orine.	Closely spaced qtz. vein- lets in monzonites, qtz. monzonites, qtz. gran- odiorites, and qtz. dlo- rites are favorable con- ditions.
Deposit Features	No predominant shape. Size can vary from 20 to 1100x MT. Most contain 0.6% copper. Primary minerals are pyrite, chalcopyrite, and molyb- denite.	Commonly in domai, arcuate, funnel-shape, or annular bodies. Range from 50 to 900 MT. Over- all grade from 0.1-0.3%; molybdenlte is primary mineral.	Most are circular in size. Range in size from 15 to 700 MT. Main min- erals are molybdenite and pyrite. Grade can range from 0.05 to 0.15%.
Alteration	One type always found. Five types are possible. (propylltic, argillic, potassic, quartz-seri- cite, and lime-silicate.	Intense in intrusive as well as in host rock. Argillic and propylltic alteration common on top and sides of deposit.	Most have potassic core followed by quartz- sericite, arglilic, and propylitic alteration.
Ore Controis	Closely spaced fractures and faults as the result of folding, and doming of host rock.	Intense fracturing of Intrusive and intruded rocks.	Stockwork of fractures and breccla zones within intrusive.
Weathering	Widespread iron oxide staining depending on pyrite content. Leached capping may form.	Most show Iron staining but may also show black, yellow, and red varia- tions.	lron stalning usually present but may also have yellow and green stains.
Geochemical Expression	Anomalous amounts of Cu, Mo, Au, Ag, W In central part. Pb, Zn, Au, Ag, Mo, As, Sb, Te, Mn, and Rb In peripheral areas.	Anomalous amounts of Mo, Sn, W, and Zn. Cu is ab- normally low.	Many show anomalous zon- Ing outward from Mo to Cu to Au to Ag to As.
Geophysical Expression	Some deposits situated on flanks of magnetic highs.	Possible favorable areas along margins of gravity lows.	None reported.
Associated Mineralization	Some Pb, Zn, Ag, velns found at outer fringes of system.	Sn and W may be present in enough quantities to be recovered.	Base metals in sporadic veins sometimes present in peripheral zone.

	#2.03	#2.04	#2.05
Regional Setting	Calc-iron type related to island arcs and rifted continental margins. Mg-iron type related to continental margins.	Mostly in orogenic belts at continental margins. back-arc spreading	Mainly on continental margins,
Structural Features	Deposits form mainly at Igneous-calcareous contact.	Takes on porphyry copper features in some cases. In others, no consistent recognizable features.	No particular struc- ture. Nearby igneous - bodies probably related to zones of weakness.
Stratigraphic and Lithographic Characteristics	Associated with Igneous rocks ranging from gran- lte to gabbro. Dolomite favored over limestone. Mostly Mesozoic or younger.	Porphyry felsic stocks are favored that are Mesozoic or younger.	Wide range of calcareous sediments favorable from Paleozoic to Mesozoic. Intrusives can range from granite to syenite.
Deposit Features	Tabular to irregular along intrusive contact. Massive amounts of garnet, pyroxene, and epidote in skarn.	Irregular in size, commonly contain a mixture of chalcopyrite bornite along with pyrite, hematite, magnetite, pyrrhotite, and sphalerite.	Sphalerite and galena with small amounts of Cu, Au, and Ag. Distinc- tive feature is Mn min- erals in skarn.
Alteration	Widespread and intense alteration of igneous rocks most often found with calc type.	Variable amounts in an interzone of potassic alteration. Hornfels and marble may form in calcareous beds.	Intense but local altera- tion of intrusive pro- duces Mn accessory min- erais.
Ore Controls	Fractures attending intrusive emplacement main contributor, permeability of sedimen- tary rocks also important.	Permeablilty and bedding planes of sediments important as well as intrusive induced fractures.	Skarn usually develops away from igneous contact. Lithologic contacts and faults in host rock important.
Weathering	No distinctive pattern. May show iron oxide staining and residual magnetite float.	Oxides and carbonates of copper may form and/or a gossan.	Gossan capping may form and/or iron staining.
Geochemical Expression	Anomalous amounts of Cu, Co, Au around some de- posits. Calc types sometimes show Zn, Au, Ni, Mo, and Ag.	Anomalous Cu, Pb, Zn, Au, Ag, Mo, and possibly Bi.	Anomalous Ag, As, Be, Co, F, Pb, Sn, and W found In near vicinity. Zn found more distal.
Geophysical Expression	Very little known. Aeromagnetics might be used under favorable conditions.	Magnetic or electrical methods depending on mineral sulte.	Little is known of any signature. Electrical or magnetic might be used under favorable conditions.
Associated Mineralization	No known associated mineralization.	Sometimes porphyry copper, zinc skarn, or Zn-Pb replacements in outer carbonate beds.	Frequently associated with copper skarns.

	#2,06	#3.01A	#3.01B
Regional Setting	Majority of deposits related to continental areas subjected to late to post orogenic activity.	Situated along orogenic belts of continental margins.	From mainly in eugeosynclinal troughs bording volcanic domes.
Structural Features	Granitic stocks and batholiths intrudeded into calcareous sediment.	Associated with pillow lava and graben-type faulting.	Usually has graben-type faulting and display medium to high metamorphism.
Stratigraphic and Lithographic Characteristics	Argillaceous to inter- calated carbonate- volcanic sequences most favorable host.	Occur mainly in lower mafic volcanics and diabase dike complexes of ophiolites that range from Precambrian to Mesozoic.	Most deposits found in volcanic sequences that grade upward from mafic to intermediate to felsic tholeiltic basalt.
Deposit Features	Forms near intrusive contact in stratiform structures and pods. Scheelite and verying amounts of Mo form ore. Skarn made up mainly of garnet and pyroxeme.	Occur as tabular, lenticular, or saucer-shaped bodies in pillow lava. Consist mainly of pyrite and chalcopyrite.	Lenticular to tabular bodles of pyrite and/or pyrhotite, galena, sphalerite, and chalcopyrite.
Alteration	Nearly all alteration Is in sediments where skarn forms.	Mafic rocks altered to zeolite and green-schist facies. Mn and Fe-rich cherts may be present.	Most alteration in the root zone. Chloridization and sericitization found at the sides.
Ore Controls	Dip of the host in relation to intrusive as well igneous induced fracture and faults important.	Pillow basalt structure and brecclation of mafic volcanics of prime importance. Also faults and fractures.	Zoning common with Pb at top grading downward to Zn and Cu. Large amounts of iron suifide also commonly forms around ore body.
Weathering	Oxidation of skarn will cause iron staining.	Tan to brown gossan may form if conditions are favorable.	A lenticular gossan may form varying in color upon the minerals present.
Geochemical Expression	Above normal amounts of Wo, Cu, Mo, Bi, Be, Sn, F, and Cb usually near Intrusive contact.	Anomalous amounts of Zn and Co found near center. Cu, Zn, and Co commonly found in peripheral areas.	Anomalous Pb and Au usually found in gossan. Iron sulfide halo is usually barren of Pb, Zn, Cu, Au, and Ag values.
Geophysical Expression	Magnetic and electrical methods may be useful depending on mineral sulte.	Aeromatics frequently used but success depends on mineral suite.	Electromagnetics are most often used and indicated by positive anomalies.
Associated Mineralization	Occasionally Zn or Sn skarn associated with this type.	None known to occur.	Gold-quartz veins are sometimes found beneath but at considerable distance from this mineral type.

	#3.01C	#3.02	#4.01
Regional Setting	Associated with very extensive eugeosynclinal volcanism occurring in early tectonic cycle.	Most are thought to have formed in an eugeosynclinal subduction complex that formed at the base of island arcs, oceanic ridges, and basin margins.	Spatially related to plutons.
Structural Features	No distinctive structure other than nearby high- angle faults.	Regional rocks are commonly basalt pillow lava.	Thin bedded, brittle sediments subject to fracturing and faulting from intrusive action. Favorable areas are intrusive contacts and roof pendants.
Stratigraphic and Lithographic Characteristics	Occur in thick sequences of volcanics grading upward from mafic to felsic. Most deposits are late Precambrian to Devonian.	Host is usually radiolarian chert intercalated with basait, greenstones, tuff, and graywacke.	Intrusives are generally granodiorite or qtz. monzonite. Solubie- type sediments, such as limestones and dolomites, are usually favored.
Deposit Features	Irregular lenticular to tabular In shape. Massive pyrite and pyrrhotite with sphalerite and chalcopyrite.	Typically occurs as massive black lenses of crystalline or amorphous Mn oxide, carbonate, and manganiferous chert.	Usually irregular or podlike in shape. A wide variety of ore minerals can be present, but galena, sphalerite, and chalcopyrite are nearly always present.
Alteration	Usually found at base in the form of chioritization, sericitization, silicification, and tourmalinization.	Supergene enriched altered Mn minerals. Basalts altered to greenstones along with silicification and injection of ferriferous solutions.	Alteration limited mainly to bleaching and recrystallization of the carbonates, if present.
Ore Controls	Many located near felsic domes, calderas, and volcanic necks. Mineral zoning common with zinc at the top and increasing copper near bottom.	Faults and fractures as well as permeable beds controlled emplacement of ore.	Configuration of the intrusive and bed attitude of the host as well as joints and fractures important.
Weathering	Dark red or brown gossan found on some deposits.	Secondary Mn oxides are usually evident.	No unique weathering feature normally present. Oxidation of iron minerals may form ocher.
Geochemical Expression	Anomalous Au in gosan but no Cu or Zn except in halo area around deposit.	Anomalous amounts of Ba, Li, Rb, Zn, and Cu can be expected nearby.	Zonal arrangement of anomalous Cu at the center followed by Pb-Ag, Zn, and Mn at the outer edges.
Geophysical Expression	Electromagnetics have been used successfully. Aeromagnetics may detect sulfide bodies if enough pyrrhotite present.	None known.	Buried plutons may show as positive highs using regional aeromagnetics. Other methods not overly successful.
Associated Mineralization	Au-qtz, veins sometimes found in root zone beneath these deposits.	None definitely identified.	Some precious and base metal veins may occur in the contact zone.

	#4.02	#4.03	#4.04A
Regional Setting	Carbonate beds deposited In an eugeosynclinal environment that later becomes tectonically active.	Found mainly along mobile belts of the continen- tal margins, namely Central Sierra Nevada Mountains and Southeast Appalachian Mountains.	Most are located along continental margins or back-arc basins subjected to orogenic activities.
Structural Features	Moderate regional folding with many high angle faults. Dikes and plutons generally found in the near vicinity.	Orogenic folding and overthrusting causing thrust faults and secondary fractures.	Strong fracture system present such as Basin and Range faulting, caldera ring fractures, caldera-related graben structures, or complex faulting in domal areas.
Stratigraphic and Lithographic Characteristic	Carbonaceous cabonate rock are highly favored, age not critical. S Associated intrusives generally range from qtz. diorite to qtz. monzonite.	Regional metamorphism has produced slates, schist, greenstones, and gneisses. Later intruded with rocks ranging from peridotites to qtz. diorite.	Brittle rocks are most favored because of the open spaces created when deformed.
Deposit Features	Micron size Au in stratiform zonal patterns encircling areas of silification. Show a spatial relation to high angle faults.	Occur as quartz veins or mineralized country rock. Au usually accompanied with pyrite, galena, sphalerite, and arsenopyrite.	Occur as veins, stock- works, and replacements In volcanic rocks and sediment hosts. Ore can be native Au-Ag, tellurides, sulfides of As, 5b, Ag, Hg, and base metals.
Alteration	Bleeching and leaching of large areas usually exhibited. Replacoment of carrenets by silica and development of clays most prominent features.	Carbonization or ankeritization chief alteration feature. Wall rocks sometimes altered to sericite, pyrite, and arsenopyrite.	Silicification or jasperoid zone above ore is common. Acid leach may occur above and below the ore. Clay alteration in irregular patches.
Ore Controls	Tactonic folding of brittle host rocks creating fractures and faults.	Joints and faults in brittle host rock; slates are particularly favorable.	A throughgoing braided fracture system needed. Some may show gradual zoning from precious metals outward to base metals.
Weathering	Very little distinctive changes. Some beds may lighten in color.	Mechanical concentration can produce rich near surface and downslope placers.	Areas of bleached country rock that may have gossan cap. Clay minerals present where protected from erosion.
Geochemical Expression	Anomalous amounts of As, Sb, Hg, Tl usually found. Also higher than normal Ba, Mo, W, F frequently present.	Near mineralization, As is only mineral found in persistent amounts.	Anomalous As, Sb, Hg, Au, Ag. High amounts of Pb and Zn may indicate only roots are left.
Geophysical Expression	No distinctive pattern recognized or detectable by any present method.	have been used to map associated ultrabasic	No distinctive signature. Gravity and magnetic methods might be used to delinate favorable structure.
Associated Mineralization	Small prospects of Ag, Pb, Zn, Cu, and Hg some- times found in near vicinity.		Erosion forms rich residual Au placers.



	#4.04B	#4.05	#4.06
Regional Setting	Commonly found in Island arcs and back-arc spreading areas.	All deposits are found in areas of the Great Basin having felsic intrusive and extrusive rocks of Tertiary Age.	Nearly all deposits found within or around the borders of volcanic centers.
Structural Features	A throughgoing fracture system or ring fracture zones associated with calderas and domai structures.	Areas of complex high- angle faulting found at caldera margins, resur- gent domes, horst and and grabens, or large strike-slip faults with accompanying normal splays.	A deep seated fault zone is needed to produce a fractured system.
StratIgraphic and Lithographic Characteristics	Most deposits found associated with Intermediate calc- alkalic rocks that are nearly always porphyritic.	Host rocks are usually volcaniclastics, lake beds, tuffaceous sandstones, or volcanic tuff brecclasall having high permeability.	Most are found in andesite lava flows and tuffs, andesite dikes, and volcanic vent breccla of Tertiary age.
Depos1+	Forms in irregular, crude pipe-like or tabular ore bodies. Au occurs mainly within pyrite along with enargite and tetrahedrite.	Ore occurs mainly as dissemination in favorable beds or zones. Small veins and stringers may also occur and be enriched, particularly near the surface.	Size and shape of deposits vary greatly. Forms disseminations in favorable rock units. Cinnabar and pyrite always present.
Alteration	Argillic alteration in ore zone. Propylitized rocks form farther out as a ring or as patches.	A central area of silicification followed by argillic and/or phyllic alteration; but the pattern can be very irregular.	Conspicuous lack of alteration.
Ore Controls	Ore always found in the more permeable rocks, and fractured regions. Zoning may occur with base-metals at the outer fringes.	Permeability of the rocks and a well developed system of fractures are needed.	Mineralization controlled by steeply dipping fractures cutting permeable units.
Weathering	Argillic zone produces outcrops of yellow, red, orange, and brown if conditions permit.	Oxidation of the pyrite and bleaching by the alteration generally produce outcrops of tan, browns, reds, or white.	lron staining with the possible formation of jarosite and gypsum.
Geochemical Expression	Anomalous amounts of Au, Ag, As, Sb, Pb, Bi, Hg, and possibly Te, Se, or Mo.	Anomalous amounts of As, Hg, Sb, Tl are sometimes found in the altered area, but the amount is highly variable.	Anomalous amounts of Hg, As, and Sb generally present.
Geophysical Expression	Magnetics have been used to determine extent of arglilic alteration when covered. Electrical method has also been used, but results are difficult to Interpret.	No particular diagnostic signature known. Infrared photos might delineate argillic areas.	Little is known of distinctive signature. Electrical methods might detect pyrite if weathering not too deep.
Associated Mineralization	Some deposits found above and along periphery of copper porphyries.	Epithermal precious metal qtz. veins and base metal sulfides may ile at depth.	Stibnite veins are known to occur occasionally in the general vicinity.

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	#4.10	#4.11	#5.01	
Regional Setting	All deposits are associated with volcanic centers of Tertiary Age.	Found along continental margins subjected to major deformation and associated igneous intrusions.	Found in intercratonic rift basins and the shallow-marine environments that developed over filled rifts.	
Structural Features	Ring and radial faults and breccias developed along major faulting are important.	Miogeosyncline sequences of sediment are favor- able that have later been intruded by small plutons or other intrusive complexes.	Main tectonic features are block faulting and gentle folds.	
Stratigraphic and Lithographic Characteristics	Most deposits occur in flows, tuffs, brecclas, and agglomerates and range from rhyolites to basalts in composition.	Calcareous sediments such as limestones, dolomites, and marble are favored. The intrusive rocks are commonly granite or granodiorite.	Thick sequence of sediments usually oxidized to red, gray, green, or white in color. May be capped with evaporites.	
Deposit	Usually found in veins that may form into bunches and pockets, nodular masses, or disseminations. Psilomelane, pyrolusite, Draunite, wad, and manganite are ore minerals.	Forms in tabular veins, irregular open spaces, lenticular pods, chimneys, or pipes. Ore minerais are psilomelane, pyrolusife, rhdochrosite, wad, menganife, and rhodonife.	Dissominated minerals in one or more zones of sandstone or shale, especially those containing organic material.	
Alteration	Kaolinite is the most significant alteration product.	Skarn may form in the calcareous host at the igneous contact.	Varying degree of alteration and regional metamorphism may cause changes in color.	
Ore Controls	A throughgoing fracture system and/or volcanic breccia are needed.	Porosity of the calcareous rocks and attitude of the igneous contact are important.	An oxidation-reduction interface and low pH needed along with carbonaceous material to enhance concentration of Cu.	
Weathering	Usually a mixture of Iron and manganese oxide, and kaolin forms at the surface.	Limonite and kaolinite may develop at the outcrop along with blackening of the surrounding rocks from the soluble Mn.	Copper and iron staining can occur under ideal conditions. Cu may be completely leached and deposited down dip.	
Geochemical Expression	Anomalous amounts of Mn over a wide area. Anomalous patches of Pb, Ag, Au, and Cu may occasionally be present.	Anomalous amounts of Mn over a wide area and possibly high amounts of Cu, Pb, and Zn.	Anomalous amounts of Cu, Pb, Zn, Y, and Mo most often found. Ag and U may also be present.	
Geophysical Expression	No characteristic signature is known.	No characteristic signature known.	No signature known.	
Associated Mineralization	Epithermal Au and Ag deposits are known to occur sometimes in the general vicinity.	Silver-bearing replace- ment veins are known to occur sometimes in the near vicinity.	Evaporites in economic quantities may overile the red-beds.	

	#5.02	#5.03	#6.01
Regional Setting	Most deposits restricted to stable platforms or forelands, coastal plains, of the continental shelf, and intermontane basins.	Found in many environ- ments but most are in eugeosynclinal areas within orogenic belts accreted to continental margins.	Commonly found in the western coast ranges that were tectonically active during the Cenozoic.
Structural Features	Wide continental basins having fluvial, lacus- trine, deltaic, and strand-plain features are favorable.	Some are associated with hinge faults that develop in sedimentary basins.	Near major fault zones.
Stratigraphic and Lithographic Characteristics	Medium to coarse grained sendstones composed of devitrified volcaniclas- tics, or feldspathic material with carbonaceous material are especially favorable.	Nearly all deposits are in well-bedded siliceous and/or limestones. Many are Devonian age but may be older or younger also.	Form where high energy streams gradients flatten and water velocity lessens.
Deposit Features	Generally occur as tabular bodies or as roll-front structures occupying part or all of a permeable rock unit. Pyrite nearly always present with U minerals.	Forms in disseminations in stratiform bodies that are often lense- shaped or tabular. Barite is the ore mineral.	Gold particles usually are small flattened flaky particles ranging downward in size to powder. Accessory minerals are magnetite, chromite, garnet, and rutile.
Alteration	Sandstone altered to pale gray, brown, pink, or yellow.	Normally dark colored beds may be altered to light gray or white.	Not found.
Ore Controls	Permeability of a host rock at low dips and containing carbonaceous material is of prime importance.	Basin structure needed to trap migrating barium solutions.	Any impediment of water flow carrying gold particles can create traps and enhance concentration.
Weathering	May produce U oxides of yellow, green, brown, and black. Pyrite will produce brown patches.	No distinctive features are generally formed. Occasionally barite rosettes or nodules develop at the outcrop.	Tends to move gold particles through alluvium to bedrock.
Geochemical Expression	Anomalous amounts of Se, Mo, V, Cu, and U are indicator elements.	High Ba and Sr with very low Ca and Mg is the most common signature.	Anomalous amounts of Ag, As, Hg, Sb, Cu, S, Fe may be found but not always as a group.
Geophysical Expression	Anomalous radioactivity of 5 to 10 times back- ground count by gamma-ray spectrometers is generally significant.	No diagnostic pattern is known.	No characteristic signature.
Associated Mineralization	Red-bed copper deposits may occur in the same general area.	Occasionally marine exhalant type Pb-Zn deposits occur in the general vicinity.	Epithermal gold veins may be nearby. In ultrabasic areas platinum may be associated with gold in placers.





#6.02

Regional Setting	Limited to former shelf and platform areas and/or areas of miogeosynclines and eugeosynclinal deformation.
Structural Features	Favored environment is marine sedimentary basins with shoaling bottoms.
Stratigraphic and Lithographic Characteristics	Lateral sequences of carbonaceous shale, dolomite, chert or diatomite, saline beds and red or light-colored sandstone.
Deposit Features	Phosphorite occurs as pellets, nodules, shell and bone material in thick bads. May giso include calcite, quartz, clays, gypsum, halite, and pyrite.
Alteration	No alteration develops in these deposits.
Ore Controls	Favorable structures are stratigraphic basins or parts of basins.
Weathering	Phosphorite weathers to a dull bluish-white film.
Geochemical Expression	Anomalous amounts of P, N, F, C, and U are usually present.
Geophysical Expression	Gamma-ray well logging can usually detect the U content. Sometimes aerial gamma-ray spectrometry has been used.
Associated Mineralization	Sedimentary Mn may be in the same general area because of similar geologic environment.

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